U.S. DEPARTMENT OF THE INTERIOR U.S. GEOLOGICAL SURVEY

DATA REPORT FOR THE 1991 BAY AREA SEISMIC IMAGING EXPERIMENT (BASIX)

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INTRODUCTION

Seismicity in the San Francisco Bay Area has increased dramatically during the past decade following a long period of quiescence associated with the 1906 San This increased seismicity is best illustrated by the Francisco earthquake. October 17, 1989, 7.1 M earthquake along the Loma Prieta segment of the San Future earthquake concerns focus on two possibilities: 1) Andreas fault (SAF). that the Loma Prieta earthquake has loaded the Peninsula segment of the SAF. and 2) that the historical pattern of large paired quakes -- involving faults along the Peninsula and the East Bay -- may be about to repeat itself. Fault occurrences of paired earthquake activity between the Peninsula and East Bay have been documented in the last century and in the early 1900's (Table 1). Although these periods of paired seismicity may be fortuitous, it is possible that the network of faults that bound the San Francisco Bay Area are linked structurally and are all part of an evolving fault system, as proposed by Furlong and others (1989; Fig. 1).

		TABLE 1 PAIRED_EARTHOUAKES
1936 1938	M6.8 M7.0	Hayward Fault, northern segment San Andreas Fault, south San Francisco
1865	M6.5	San Andreas Fault, southern Santa Cruz Mts. Hayward Fault southern segment
1906	M8.3	San Andreas Fault
1911	M6.5	Calaveras Fault
?	?	San Andreas Fault, Santa Cruz Mountains ?

To test the Furlong and others (1989) model and to obtain fundamental information on the crustal structure and fault geometries that underlie the San Francisco Bay Area, the U.S. Geological Survey (USGS), in conjunction with the University of California, Stanford University, Pennsylvania State University, and Lawrence Berkeley Laboratory, conducted a major seismic reflection investigation in the fall of 1991 utilizing the marine waterway system that dissects the Bay Area (Fig. 2). This study, known as the Bay Area Seismic Imaging eXperiment, or BASIX, consisted of three complementary seismic reflection and refraction profiling methods: high-resolution seismic reflection profiling, wide-angle reflection/refraction profiling, and multichannel seismic profiling. The high-resolution, shallow-penetration data were acquired to constrain near-surface faulting, while the wide-angle reflection/refraction data were acquired to constrain the deep-crustal velocity structure of the crust. The focus of this report are the multichannel, seismic reflection data which were acquired to image structures throughout the crust down to the Moho.



Fig. 1. Schematic drawing of the lithospheric structure and fault c onfigurations in the vicinity of San Francisco. The surface plate boundar y occupies the approximate position of the previous western edge of North American plate, but the deeper shear zone boundar y is offset to the east (~40 km). Newly forming faults sit above the deeper eastern segment. Stippled region is asthenospheric mantle that has cooled and accreted to adjac ent lithosphere. Cross section is approximately 200 km (width) by 60 km (depth) in extent. Figure adapted from Furlong and others (1989).



FIGURE 4. Index map showing location of BASIX seismic study. Individual lines are numbered.

INSTRUMENTATION

The Seismic Source

The seismic source consisted of a 12-element, 2000 psi, 5828 cu. in. (95.6 liters), airgun array towed from the U.S. Geological Survey's research vessel, the S.P. Lee (Fig. 3). Airgun sizes were selected to maximize the peak-to-peak amplitude (85.3 bar-m) and the peak-to-bubble ratio (12.7) while maintaining maximum volume. Because of the shallow-water environment for most of the experiment, airguns were towed at an average depth of 7.6 m (25 ft); depths were maintained by buoying each gun with a separate Norwegian buoy. The airguns were fired at 50-m intervals in the inland and Golden Gate profiles (lines 101-113, OBS1, and 201-202) and at 75 m for the offshore transit and continental margin profile (TR1 and OBS2). Shooting was conducted during night-time hours when ship Acquisition was timed so that the traffic and ambient noise were a minimum. ship moved against the tide as much as possible; this resulted in minimum speed over the ground (approximately 3.5 knots) as well as maximum control and maneuverability of the vessel. A total of more than 11,634 airgun shots were fired during the fourteen-day experiment (Appendix A).

The Receiver Array

The seismic energy was recorded on hydrophones moored to the edge of the dredged shipping channel in the bay. The hydrophones and associated telemetry receivers, known as Telseis (Fig. 4), were leased from Fairfield The Telseis units were equiped with an antenna and a transmitter. Industries. which radioed the analog seismic data to a 120-channel, DFS-V recording system on the S.P. Lee. Separate radio frequencies had to be licensed for each of the 120 Telseis units (72-79 MHz band). Sixty to 118 of these units were used on any given night; battery limitations often required keeping half of the Telseis units onboard the Lee for recharging in preparation for the following night's In addition to the Telseis receivers, a short two-channel streamer was shooting. also deployed off the fantail of the Lee; this streamer provided a near-offset image of the crust. All data were recorded to 16 s two-way traveltime at a 0.004 s A listing of the number of Telseis deployed each night, the group sample rate. spacing, and the number of shots is provided in Appendix A.

Navigation

Navigation was based on two independent systems. The land-based Del Norte radio-navigation system was the primary method used for positioning and resulted in station locations accurate to ± 1 m. Up to 9 transponders were positioned in the area of operation so as to provide sufficient control for triangulation. These stations were moved daily as the survey moved from east to west. The Del Norte system proved successful throughout most of the study, but there were some areas where reception was limited and station coverage was not sufficient to provide an accurate fix. In these regions locations were determined from a single Global Positioning System (GPS) receiver. Because this GPS system did not operate in differential mode, absolute station accuracy was approximately ± 100 m -- relative positions from station to station, however,



configuration. Individual chamber sizes are listed in cubic inches. The chamber sizes were selected to maximize the peak-to-peak amplitude (85.3 bar-m) and the peak-to-bubble ratio (12.7) while maintaining maximum volume.



Figure 4. Schematic sketch of Fairfield Technologies Telseis I seismic telemetry exploration system.

were accurate to within a few meters. Navigation for lines 202 and OBS2 seaward of the San Andreas fault was accomplished using a combination of GPS operated in a selected availability mode and rho-rho Loran-C; positions are considered accurate to within 150 to 200 m.

METHODOLOGY

The BASIX investigation was designed to take advantage of the extensive marine waterway system that dissects the San Francisco Bay Area. A total of 140 km of multichannel reflection data were acquired across this region (Fig. 2) during 14 nights of profiling. Although each night of profiling is assigned a different line number, several of the lines overlap and they were thus processed together as merged profiles. All totaled, there are four distinct segments to the The first segment begins in the Sacramento River near Rio Vista experiment. and proceeds west through Honker, Suisun, and San Pablo Bays (lines 101 through 110; Figs. 5, 6, and 7). The next two segments run north-south from Richmond to Angel Island (line 113; Figs. 8a and 9), and from the Bay Bridge to just north of the San Mateo Bridge (line 111; Figs. 8b and 10). The fourth and final segment begins in the vicinity of Alcatraz Island and extends west underneath the Golden Gate bridge, 6 km into the open waters of the Pacific (lines 201-202; Figs. 11 and 12).

The shallow water depths and abundant ship traffic in the San Francisco Bay region precluded using a standard towed-streamer approach to seismic data acquisition. Instead, 60 to 118 independent Telseis receivers were moored to the edge of the dredged shipping channel at 100 to 200 m intervals for a fixed spread of 6 to 12 km in length. Each day the receiver array was repositioned, and each night the S.P. Lee steamed along the hydrophone array, with its airgun array firing.

The day-to-day small-boat operations began first with the retrieval of the Telseis instruments deployed the previous day. This process utilized a fleet of 4-5 small boats that ranged in size from 6 to 15 m. Once retrieved, the instruments were brought back to the S.P. Lee for recharging. The R.V. David Johnston trailed behind these small vessels and recovered the associated buoys and anchors. At this time the David Johnston also recorded navigational fixes, which were required in order to check for possible drift of the instrument sites during the night; although anchors were used to maintain instrument position, currents were strong enough in a few areas to move the anchors between deployment and retrieval.

Once the instrument retrieval was complete, the David Johnston would begin laying out the next line segment. A Global Positioning System (GPS) and Del Norte triangulation system were used to measure the 100- or 200-m distance separating instrument sites. The receiver line was positioned so as to minimize bends in the profile and to keep the buoys out of the active shipping channel. At each site, anchors were deployed and navigation fixes were taken; the anchored sites were marked by orange buoys. The smaller boats followed behind and clipped the recharged Telseis units with cabled hydrophones to these buoys. The hydrophones were weighted so as to minimize movement on the seafloor.



Figure 5. CDP locations for lines 101 through 110.













Figure 8. CDP locations for (A) Line 113, and (B) Line 111.

B.

Α.











÷.

Figure 11. CDP locations for lines 201-202.





The final phase of operation was airgun profiling. Profiling was conducted during night-time hours so as to take advantage of reduced ship traffic and lower levels of ambient noise. Each night the S.P. Lee acquired approximately 24 km of reflection data by looping through the receiver array twice. Along 10 of the 13 multichannel lines (Appendix A) the S.P. Lee fired the airgun array while steaming first from east to west and then, reversing course, duplicated the line by steaming from west to east, further doubling the number of airgun shots available for stacking of the seismic signal. Although the receiver array averaged 6 km in length, the Lee continued profiling for approximately half a spread length on either side of the array (approximately 12 km total for each pass) so as to maximize offsets and fold. Two passes through the array typically required 10-12 hours; profiling was timed so as to acquire as much data as possible while operating against the tides. Slack tides proved to be the best time for profiling due to the reduced currents and therefore reduced noise conditions.

In addition to the 13 reflection lines acquired in San Francisco Bay and the Golden Gate region, the airgun array on the S.P. Lee was used primarily as a seismic source for wide-angle profiles recorded along three other lines: the N-S trending line OBS1 in San Francisco Bay; a short transit leg, line TR1 southwest of the Golden Gate; and the E-W trending line OBS2 on the continental margin (Fig 2). Lines OBS1 and OBS2 were obtained along two separate deployments of 6 USGS ocean bottom seismometers (OBS) deployed by investigators from the USGS Branch of Atlantic Marine Geology and Woods Hole Oceanographic Institution (Holbrook and ten Brink, 1991; Brocher and others, 1991). The shot interval along line 112/OBS1 was 50 m. The shot interval for lines TR1 and OBS2 was approximately 75 m. Line OBS2 was located coincident to seismic reflection line 13 of Lewis (1990) in order to provide deep structural control on the continental margin west of the San Andreas fault (Fig. 2). The reader is referred to Holbrook and others (1993) and Brocher and Moses (1993) for a discussion of these wide-angle data.

DATA REDUCTION

After acquisition, the data were demultiplexed, and the 16-s records with 4 ms sample intervals were truncated and resampled to 12-s records with 8 ms sample intervals. The data were then sorted to common receiver gathers for editing. Once the navigation data were reduced, the shot-receiver offsets and midpoints were calculated and inserted into the trace headers. Trace editing, which was the most effort-intensive part of the data reduction process, followed.

There were several sources of noise that required different editing techniques. Because of the use of a telemetry recording system, data drop-outs resulted when radio communication was hampered by land barriers and large ship-to-receiver distances. Signal was rarely present beyond 6 km offsets. The crooked nature of the reflection profile also resulted in a smearing of midpoints out of the plane of the common midpoint (CMP) profile. Even though the data were binned using a crooked line geometry, offsets perpendicular to the plane of the profile needed to be restricted to minimize smearing. (Note that because the receivers were on the edge of the shipping channel and the *Lee* profiled down the center of the channel, zero offset data were rarely acquired. A certain amount of smearing was thus required by the geometry of the study.) In-plane offsets were typically restricted to 6 km, and out-of-the-plane mid-point scattering was restricted to \pm 0.5 km (except lines 201-202, which had up to a \pm 1 km midpoint scatter).

A second source of noise was the jostling of the hydrophones during times of strong currents. As tides increased during the night, noise on the receiver gathers also increased, often overwhelming any seismic signal (Fig. 13). This noise appeared as hundreds of individual spikes of approximately the same amplitude as the first-arrival refractions. A noise reduction algorithm was developed to edit these noise spikes. The algorithm was designed to compare the amplitude of a reference trace with the average amplitude of adjacent traces (the number of which is specified by the user). When the ratio of the amplitude of the reference trace exceeded that of the surrounding traces by a userdesignated amount, that sample was zeroed (compare Fig. 13 with Fig. 14 to see the effect of this noise editing routine).

Although these two editing techniques reduced the noise content of the data set significantly, areas still remained where either there was no signal or where coherent noise (i.e.: refractions) obscured weak reflections. Manual editing for each receiver gather was thus required to further reduce the noise content of the data. This manual editing was done twice, once for receiver gathers and once following sorting to CMP gathers.

Once the editing was complete, the data volume was approximately one tenth of the original volume (reduced from 790,000 traces to 95,535 traces); this equates to a reduction in the nominal fold from 500 to approximately 65. Although this reduced the effectiveness of the stacking process, the final image was improved simply due to the absence of the dominant noise.

Additional processing steps for the BASIX data included first-break mutes, sorting to CMP gathers (using the crooked-line geometry), automatic gain control (1-s gate), velocity analysis, surface-consistent statics, residual velocity analysis, normal moveout, stacking (variable fold, average is approximately 40), and bandpass filtering (14-44 Hz). Both surface-consistent statics and detailed trace editing were important steps in generation of the final stack.

The data are written to two 8mm Exabyte tapes. The first tape contains the edited receiver gathers (12 s, 8 ms sample rate), with geometry applied so that the data can be resorted into CMP order. Receiver locations were unique, and the dataset is written as a single SEG-Y volume. The second tape contains stacked data (12 s, 8 ms sample rate) of the complete transect. This tape was also written as a single SEG-Y volume, but consists of four separately processed line segments. A line identifier is included in the SEG-Y header, allowing each segment to be extracted by unique CMP or line segment. These two data tapes are available from the National Geophysical Data Center at the following address:



Figure 13. A single common receiver gather from San Pablo Bay recorded during two passes of the S.P. Lee. No scaling has been applied in order to illustrate noise problems. The 744 traces shown on this gather are from six hours of shooting. Noise from tidal currents and surface-wave action increased during the night.



Figure 14. Common receiver gather (Figure 13) after application of automatic noise burst editing software. The amplitude of a sliding window on each trace is compared to the same time windows on adjacent traces, and anomalously high values are zeroed. Coherent energy such as reflection and refraction events are not affected by this editing. The bar chart above the record shows the percentage of each trace that has been zeroed. National Geophysical Data Center NOAA/EGCI 325 Broadway Boulder, CO 80303 Telephone: (303) 497-6123.

The EBCDIC trace headers for each of the three data tapes are listed in Appendices B and C.

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Appendix A

Line #	Location	Date	<u>#Shots</u>	<u>Group Int.</u>	<u>#Telseis</u>	<u>#Passes</u>
101	Rio Vista	9/4/91	723	100	118	2
102	Antioch	9/5/91	483	100	118	1
103	Suisun Bay	9/6/91	599	100	118	1
104	Martinez	9/7/91	165	100	60	1
105	Carquinez	9/8/91	946	200	60	2
107	San Pablo Bay	9/10/91	878	100	60	2
108	San Pablo Bay	9/11/91	869	100	60	2
109	San Pablo Bay	9/12/91	856	100	60	2
110	San Rafael	9/13/91	855	200	60	2
111	South Bay	9/14/91	791	200	105	1.5
OBS1	South Bay	9/15/91	871	0	0	0
113	Berkeley	9/16/91	9 07	200	60	2
201	Alcatraz	9/17/91	609	100	80	2
202	Golden Gate	9/18/91	956	100-200	71	2
TR1	Pacifica	9/19/91	102	0	0	1
0BS2	Cont. Margin	9/20/91	1024	0	0	1

Total:

11,634

APPENDIX B TAPE #1 -- SEG-Y EBCDIC REEL HEADER BASIX EDITED RECEIVER GATHERS:

C1	BRANCH OF PACIFIC MARINE GEOLOGY	U. S. GEOLOGICAL SURVEY
C2	BASIX: BAY AREA SEISMIC EXPERIMENT	SURVEY_ID: L1-91-NC
C3	AREA: SAN FRANCISCO BAY	YEAR COLLECTED: 1991
C4	LINES: 101 TO 202	TRACES IN DATASET: 95535
C5	REEL ID: BSXRECED	8MM LOW-DENSITY (2,5 GB)
C6	SAMPLE INTERVAL: 8 MSEC	SAMPLES/TRACE: 1500
C7	TRACE INTERVAL: 50 METERS	TIME RANGE: 0 TO 12 SEC
C 8		
C9	THIS SEG-Y TAPE CONTAINS COMMON-RECEIV	ER SORTED AND NOISE EDITED
C10	DATA FROM ALL BASIX LINES. HEADER NAM	E IS REC-STAT: BYTES 187-188
C 11	LINE 101 INCLUDES RECEIVER GATHERS 1-1	10
C12	LINE 102 INCLUDES RECEIVER GATHERS 111	-207
C13	LINE 103 INCLUDES RECEIVER GATHERS 208	-307
C14	LINE 104 INCLUDES RECEIVER GATHERS 312	-430 (EVEN NUMBERS ONLY)
C15	LINE 105 INCLUDES RECEIVER GATHERS 384	-502 (EVEN NUMBERS ONLY)
C 16	LINE 107 INCLUDES RECEIVER GATHERS 504	-563
C17	LINE 108 INCLUDES RECEIVER GATHERS 564	-623
C18	LINE 109 INCLUDES RECEIVER GATHERS 624	-683
C19	LINE 110 INCLUDES RECEIVER GATHERS 684	-804 (EVEN NUMBERS ONLY)
C20	LINE 113 INCLUDES RECEIVER GATHERS 806	-924 (EVEN NUMBERS ONLY)
C21	LINE 111 INCLUDES RECEIVER GATHERS 926	-1144 (EVEN NUMBERS ONLY)
C22	LINE 201 INCLUDES RECEIVER GATHERS 210)1-2180
C23	LINE 202 INCLUDES RECEIVER GATHERS 218	31-2260
C24		
C25	SHOT NUMBERS ARE LOCATED IN HEADER BY	TES 9-13
C26		
C27		
C28	FUNDING FOR THIS WORK PROVIDED BY THE	
C29	NATIONAL EARTHQUAKE HAZARDS REDUCT	ION PROGRAM (NEHRP)
C30	-	
C 31	*******	*******
C32		
C33	FOR FURTHER INFORMATION, CONTACT:	
C34		
C35	PATRICK HART OR JILL MCCARTHY	
C36	U. S. GEOLOGICAL SURVEY, M/S 999	
C37	345 MIDDLEFIELD RD., MENLO PARI	K, CALIFORNIA 94025
C38	(415) 853-8300	
C39		
C40		

APPENDIX C TAPE #2 -- SEG-Y EBCDIC REEL HEADER BASIX STACKED PROFILES:

C 1	BRANCH OF PAC	CIFIC MARINE GE	OLOGY		U.S. GEOI	LOGICAL SURVEY	
C2	BASIX: BAY AREA SEISMIC IMAGING EXPERIMENT				SURVEY_ID: L1-91-NC		
C3	AREA: SAN FRANCISCO BAY			YEAR COLLECTED: 1991			
C4	LINES 101 TO 20	2			TRACES	IN DATASET: 2678	
C5	REEL ID: BSXSTI	K			8MM LOV	W-DENSITY (2.5 GB)	
C6	SAMPLE INTERV	VAL 8 MSEC			SAMPLES	S/TRACE: 1500	
C7	TRACE INTERV	AL: 50 METERS			TIME RAI	NGE: 0 TO 12 SEC	
C 8							
C9	THIS SEG-Y TAP	E CONTAINS STA	CKED DATA H	ROM A	LL BASIX	LINES	
C10	LINES 101-110	CDP RANGE	111-1705	LINE_	ID: 101	NTRACES: 1587	
C11	LINES 113	CDP RANGE	1747-2046	LINE_	ID: 113	NTRACES: 282	
C12	LINES 111	CDP RANGE	2049-2493	LINE_	ID: 111	NTRACES: 456	
C13	LINES 201-202	CDP RANGE	3104-3465	LINE_	ID: 201	NTRACES: 353	
C14				_			
C15	CDP AND LINE_I	D ARE DEFINED	IN THE FOLLO	WING I	HEADER B	YTES:	
C16							
C17	HEADER	STARTING	LENGTH		VARIABL	Е	
C18		BYTE	(BYTES)		TYPE		
C19							
C20	CDP	21	4		INTEGER		
C21	LINE_ID	191	2		INTEGER		
C22							
C23							
C24							
C25							
C26							
C27							
C28	FUNDING FOR T	HIS WORK PROV	IDED BY THE:				
C29	NATIONAL EAR	THQUAKE HAZA	RDS REDUCTI	ON PRC	GRAM (N	EHRP)	
C30							
C 31							
C32							
C33	****	*****	****	* * * * * *	*******	****	
C34							
C35	FOR FURTHER IN	NFORMATION, CO	DNTACT:				
C 36							
C37	PATRICK HART OR JILL MCCARTHY						
C38	U. S. GEOLOGICAL SURVEY, M/S 999						
C39	345 MIDDLEFIELD RD., MENLO PARK, CALIFORNIA 94025						
C40	(415) 85	3-8300					