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# GEOTHERMAL DATA OF THE UNITED STATES 

## INCLUDING MANY ORIGINAL DETERMINATIONS OF UNDERGROUND TEMPERATURE

BY<br>N. H. DARTON



## WASHINGTON

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# GEOTHERMAL DATA OF THE UNITED STATES 

## INCLUDING MANY ORIGINAL DETERMINATIONS OF UNDERGROUND TEMPERATURE.

By N. H. Darton.

## INTRODUCTION.

The purpose of this report is to present all available published data bearing on the rate of increase of underground temperature with increasing depth in the United States, together with several hundred original observations by myself and my associates. A canvass of the governmental, State, and serial publications has yielded many records of temperature of flowing wells and also a few observations made with thermometers in borings and deep mines. Additional data on temperature of flows have been given by correspondents. Some of the data, especially those relating to flows from wells, may not be reliable, and as a rule these could not be discriminated. Many records are omitted because no facts are available as to the source of flow.

During extended investigations of the geology of underground waters in many parts of the United States I and my assistants have had opportunity to observe the temperature of flows and to sink thermometers in deep borings. One very important contribution has come from an associate, J. E. Todd, who recorded the temperature of flows from a large number of wells in central-eastern South Dakota.

Mr. C. E. Van Orstrand, of the U. S. Geological Survey, has taken observations in various exceptionally deep wells in Pennsylvania, West Virginia, Oklahoma, and Texas, with very accurate apparatus, but only a few of the data are published.

The principal feature brought out by the data here presented is the fact that the rate of increase in underground temperature with increase in depth varies widely from place to place, though probably subject to certain regional relations. There is also a local and possibly a general variation in rate at different depths, but few of the observations have afforded data on this matter. The fact that there are regional variations has been recognized for half a century and was brought out in considerable detail in Prestwich's compilation of data
up to $1886^{1}$ and his later publication in 1895. ${ }^{2}$ This observer and others later have grouped earth temperatures and attempted to show their relation to rocks and minerals of various kinds, ${ }^{3}$ petroleum, ${ }^{4}$ flowing water, etc. There are undoubtedly in the earth many factors that may influence the rate of temperature increase, such as variation in conductivity of rocks, underground tension, mineralization, volcanic influences, and movement of underground waters. Variation in radioactivity has been suggested as a factor, and the influence of bodies of cold water, such as Lake Superior, and the former presence of glacial ice are thought to be causes of local diminution of earth temperature. Positive evidence is lacking as to all these matters, however, and we must await the results of extended special investigations before the weight of the several factors can be evaluated.

In nearly all my calculations of the rate of increase in temperature with increase in depth mean annual air temperatures were used as a standard, These were obtained from publications of the Weather Bureau, United States Department of Agriculture, and nearly all of them represented averages up to and including 1916. ${ }^{5}$ In places for which no observations are on record temperatures were deduced from means of near-by stations, with due consideration of difference in latitude and altitude. It is realized that these means of air temperature may differ from the ground temperature a few yards below the surface to the amount of $2^{\circ}$ or more, but they were the only data available for comparison. ${ }^{6}$ In a few places the temperatures at shallow depths were recorded, and the rate of increase could also be calculated from them.

The thermometers used by me and my assistants and supplied to certain correspondents were maximum self-registering instruments made to order by Henry Green, of Brooklyn, N. Y. The style is similar to that used by William Hallock, but certain modifications were found desirable. The modified form has been termed the Darton thermometer, but the difference hardly merits the title. The main features are shown in figure 1. A 10-inch outer tube of heavy glass, sealed, carries an 8 -inch thermometer held by plugs of cork, C; S, stricture; r, rounded end of mercury column. The calibra-

[^0]tion is in degrees Fahrenheit and is readable to half a degree with certainty.

Most of the instruments were verified at the Bureau of Standards. As a rule they were used singly, but in certain places two instruments were used either simultaneously or in succession. They are intended to be sunk bulb end up, for the mercury in the bulb end does not easily jar past the stricture. In warm weather care is taken to chill the instrument before lowering. . It is wrapped in flannel, placed in a 14 -inch piece of 1 -inch lead pipe, and usually lowered by means of No. 18 wire, with care to avoid jars.

Generally the instrument was left down the hole overnight or at least four hours. In observing temperatures of flowing water I and my assistants used ordinary 6 -inch thermometers readable to half degrees and tested as to accuracy. Nothing is known as to the kind of instruments used for most of the observations given in published reports or furnished by correspondents. They probably varied in character.
Throughout this report temperatures are given in degrees Fahrenheit, depths in feet, and yield of wells in gallons a minute. Although the geothermal gradient should strictly be stated in units of temperature per unit of depth, it is customary, in order to get larger and more readily comparable quantities, to use the reciprocal statement, units of depth per unit of temperature. The figures are here given in feet p $\dagger$ r degree Fahrenheit. A high gradient is of course represented by a small number of feet per degree, and vice versa.

## ALABAMA.

## TEMPERATURE OBSERVATIONS.

Many observations of the teqmperature of waters from flowing and other wells in Alabama have been recorded, most of them by the State geologist. ${ }^{1}$ Unfortunately, in the greater part of these observations the source of the water is not stated, and there is uncertainty as to the accuracy of the thermometers used. The following selected data are believed to be of sufficient value to be considered.


FIGURE 1.-Thermometer used in taking deep-weld cemperatures.

Temperaturcs in wells in Alabama.

\begin{tabular}{|c|c|c|c|c|c|c|}
\hline Well. \& Total depth (feet) \& Depth of flow (feet) \& $$
\begin{gathered}
\text { Flow } \\
\text { (gallons per } \\
\text { minute). }
\end{gathered}
$$ \& $$
\begin{gathered}
\text { Tempera- } \\
\text { ture } \\
\left({ }^{\circ} \mathrm{F} .\right) .
\end{gathered}
$$ \& Mean annual air temperature ( ${ }^{\circ} \mathrm{F}$.). \& Depth per degree of increase in tempera-
ture (feet). <br>
\hline Barbour County: Eufaula, 1 mile south of \& 350 \& \& ${ }^{6}$ \& 68 \& a 64.6 \& 103 <br>
\hline Bullock County: Union Springs........ \& 848 \& ${ }^{6} 843$ \& Pumps 140 \& 76 \& c64.6 \& 74 <br>
\hline Dallas County: \& \& \& \& \& \& <br>
\hline Great well d. \& 728 \& \& 1,200 \& 773 \& c 64.8 \& 57 <br>
\hline Courthouse d \& 555 \& \& Many. \& 75 \& e 64.8 \& 55 <br>
\hline Bell's Hotel d. \& 400 \& \& Many. \& $733^{4}$ \& e 64.8 \& 45 <br>
\hline Huntersfield. \& 550 \& \& 70 \& 69 \& e64.8 \& 131 <br>
\hline Selma: $\quad$ Ferrill well $d$ d \& \& \& \& \& \& <br>
\hline Ferrill well $d$ d.......... \& 487 \& \& 110 \& 68 \& e 64.8 \& 152 <br>
\hline Main and Water streets $d$
Foundry $d$. \& 470 \& \& Many. \& 72 \& c 64.8 \& 65 <br>
\hline Fscambia Coundy ${ }^{\text {Foundry }}$ Browton........ \& - 409 \& \& 230
5 \& \& e 64.8
$f 66.1$ \& 136 <br>
\hline Greene County: \& \& \& \& \& \& <br>
\hline Boligee.. \& 500 \& 300 (first). \& 40 \& 70 \& $g 63.7$ \& 80 <br>
\hline Do. \& 415 \& \& \& 68 \& 063.7 \& 96 <br>
\hline Do. \& 250 \& \& 2 \& 68 \& $g 63.7$ \& 58 <br>
\hline Boligee, 4 milos north \& 142 \& \& \& 66 \& g 63.7 \& 62 <br>
\hline Boligee, Canfield's ${ }^{\text {d }}$. \& 522 \& \& \& 70 \& g63.7 \& 83 <br>
\hline Boligee, Finch's Ferry d \& 550 \& \& \& 71 \& g63.7 \& 75 <br>
\hline ${ }_{\text {Do }}{ }^{\text {Do. }}$ \& 320 \& 290 \& 10 \& 67 \& ${ }^{h} 63.8$ \& ${ }_{8}^{97}$ <br>
\hline Erie.................. \& 544
330 \& 300 \& \& 68 \& h
$h 63.8$

6 \& 80 <br>
\hline Eutaw, $4 \frac{1}{2}$ miles south of. \& 200 \& \& 2 \& 71 \& $i 63.5$ \& 27 <br>
\hline Eutaw, 6 miles south of. \& 340 \& 300 \& 10 \& 70 \& ${ }^{i} 63.5$ \& 46 <br>
\hline Eutaw, 9 mo. miles south of. \& 495 \& \& 2 \& 71 \& ${ }^{6} 63.5$ \& 68 <br>
\hline Eutaw, io miles south of. \& 450 \& \& 1 \& 71 \& i63.5 \& <br>
\hline Creswell No. 1... \& 456 \& \& \& $72 \frac{1}{2}$ \& i63.5 \& 50 <br>
\hline Creswell No. 2. \& 440 \& \& \& 71 \& i63.5 \& 59 <br>
\hline Creswell No. ${ }^{\text {William }}$, \& 550 \& \& \& $72 \frac{1}{2}$ \& ¿63.5 \& 61 <br>
\hline William Glover's ${ }^{\text {d }}$. \& 495 \& \& \& $72 \frac{1}{3}$ \& ${ }^{2} 63.5$ \& 57 <br>
\hline Morrison, 3 miles southwest of \& 330 \& 270 \& $4 \frac{12}{2}$ \& 68 \& j63.6 \& 66 <br>
\hline Steeles Bluff...... \& 400 \& 360 \& 22 \& 69 \& ${ }^{6} 63.6$ \& 74 <br>
\hline Locality not given $k$
Do............ \& 415 \& \& \& 68 \& j 63.6 \& ${ }_{8}^{60}$ <br>
\hline Do..... \& 525 \& \& \& 70 \& ${ }^{j} 63.6$ \& 82 <br>
\hline  \& 544 \& \& \& 72 \& j63.6 \& 63 <br>
\hline Hale County: Cypress Switch. \& 320 \& \& \& \& \& 86 <br>
\hline Evans, if miles west \& 210 \& 160 \& $24^{1 / 2}$ \& 67 \& ${ }_{m 63.8} 6$ \& 53 <br>
\hline Evansville station. \& 200 \& 180 \& 30 \& 68 \& m63.8 \& 43 <br>
\hline Evansville, several near. \& 160-200 \& 160 \& 31-18 \& 68 \& m63.8 \& 38 <br>
\hline Evansville, $\frac{1}{3}$ mile west of. \& 160 \& 140 \& 18 \& 67 \& m63.8 \& 42 <br>
\hline Greensboro. \& 200 \& 140 \& 18 \& 68 \& $n 64$ \& 35 <br>
\hline Greenwood, 3 miles east of. . \& 200 \& 160 \& 65 \& 67 \& 61.8 \& 32 <br>
\hline Greenwood, 3 miles southwest of... \& 140 \& 140 \& 40 \& 67 \& 61.8 \& 27 <br>
\hline Greenwood, 3 to 5 miles southwest of \& 160 \& 140 \& 17 \& 68 \& 61.8 \& 23 <br>
\hline Do. \& 185 \& 175 \& 12 \& 68 \& 61.8 \& 23 <br>
\hline Do.. \& 272 \& 175, 240 \& 18 \& 68 \& \& <br>
\hline Greenwood, 1 mile east of. \& 210 \& 170 \& 18 \& 67 \& 61.8 \& 77 <br>
\hline Laneville............... \& 710 \& \& Many. \& 75 \& 065 \& 71 <br>
\hline Do.. \& 719 \& \& 22 \& 75 \& 065 \& 72 <br>
\hline Mays station, $1 \frac{3}{3}$ miles south of. \& 285 \& \& Many. \& 69 \& $p 61.8$ \& 40 <br>
\hline Millwood.... \& 236 \& \& 850 \& 72 \& $q 64$ \& 29 <br>
\hline Do.. \& 360 \& 300 \& 6 \& 70 \& $q 64$ \& 50 <br>
\hline Sawyerville, 3 miles west of \& 440 \& 440 \& 17 \& 70 \& $q 64$ \& 73 <br>
\hline Stewart \& \& \& \& 69 \& \& 72 <br>
\hline Stewart.... \& 363 \& 300 \& 2 \& 67 \& $r 63.8$ \& $\stackrel{94}{9}$ <br>

\hline | Lock No. 4... |
| :--- |
| Marengo County: | \& 280 \& 200. \& 35 \& 66 \& 61.8 \& $5{ }^{\prime}$ <br>

\hline Marengo County: \& 1,200 \& 1,115 \& 20 \& 873 \& $t 65.2$ \& 1. <br>
\hline Mobile County: \& \& \& ¢ 20 \& \& \& <br>
\hline Mobile. \& 700 \& \& 400 \& 76 \& u 66.1 \& 71 <br>
\hline Do. \& 800 \& ........ \& 1,000 \& 78 \& ${ }^{4} 66.1$ \& 67 <br>
\hline
\end{tabular}

a Average for 33 years.
6 Cased to 843 feet.
c Average for 30 years.
${ }^{\text {a }}$ Am. Assoc. Adv. Sci. Proc., vol. 10, p. 95, 1856.
$e$ Selma average for 27 years.
$f$ Flomatom average for 25 years.
$y$ Livingston average minus $0.1^{\circ}$ :
4. Livingston average for 33 years.
© Livingston average minus $0.3^{\circ}$.
$j$ Livingston average minus $0.2^{\circ}$.
4. Tuomey, Michael, Geology of South Carolina, p. 247, 1848.
$l$ Mean of Greensboro and Tuscaloosa.
$m$ Greensboro average minus $0.2^{\circ}$.
$n$ Average for 38 years.

- Uniontown average for 31 years.
p Average at Lock No. 4 for 20 years.
$q$ Greensboro average.
${ }^{5}$ Greensboro average minus $0.2^{\circ}$.
- Probably cooled by mingling of upper flows.

Uniontown average plus $0.2^{\circ}$.
$u$ Averagefor 45 years.

Temperatures in wells in Alabama-Continued.

| Well. | Total dopth (feet). | Depth of flow (feet). | $\left\lvert\, \begin{gathered} \text { Tlow } \\ \text { (gallous per } \\ \text { minute). } \end{gathered}\right.$ | $\begin{gathered} \text { Tempera- } \\ \text { ture } \\ \left({ }^{\circ} \mathrm{F} .\right) . \end{gathered}$ | Mean annual air tem( ${ }^{\circ} \mathrm{F}$.). | Depth per degreo of increase in tempera(feet). |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Pickens County: |  |  |  |  |  |  |
|  |  |  |  |  |  |  |
| Sherman, 1 mile south of. | 600 |  | $6 \frac{1}{2}$ | 72 | a63. 3 | 69 |
| Sherman, 3 miles west of. | 725 |  |  | 71 | ${ }^{\text {a63. }} 3$ | 94 |
| Vienna, near. | 380 |  | 3 | 67 | ${ }^{\text {b63.3 }} 3$ | 103 |
|  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |
|  | 175 |  | 60 | $\stackrel{67}{72}$ | ${ }^{6} 64.6$ | 73 |
| Pittsboro. | 219 |  | ${ }^{1}$ | 72 | ${ }_{6} 64.6$ | 30 |
| Do...... | 185 |  | 30 | 68 | ${ }_{c}^{\text {c64. }} 6$ | 54 |
| SumterCounty: ${ }_{\text {Rut. }}$ |  |  |  |  |  |  |
| Epes... | 750 |  | 1 | 70 | d63.7 | 120 |
| Epes, 3 miles southeast of. ......... | 930 | 800-900 | 28 | 79 | ${ }^{163.7}$ | 52 |
| Epes, $3 \frac{1}{2}$ miles north of west of..... | 700 | 700 | 1 | 72 | ${ }^{\text {d }} 63.7$ | 84 |
| Gainesville, near ........ | 630 | 630 | 5. | 73 | e63. 5 | 66 |
| Gainesville, 13 miles west of........ | 700 |  | 2 | 74 | ${ }^{\text {e63. }} 5$ | 67 |
| Gainesville, 2.2 miless 0 atheast of. . . | 600 |  | 3 | 72 | e63. 5 | $70 \frac{1}{2}$ |
| Gainesville, 4 miles southeast of.... | 550 |  | 4 | 71 | ${ }^{6} 63.5$ |  |
| Gainesville, $3 \frac{1}{2}$ milcssouth of....... | 660 |  | 12 | 71 | e63.5 |  |
| Gainesville, 6 milesnor th of........ | 500 | 450-500 | 1 | 71 | 763.4 | 60 |
| Grinesville, 4 miles southwest of. . . | 700 | 700 | 1 | 71 | c63.5 | 93 |
| Warsaw ............................ | 560 | 560 | 53 | 71. | 963.8 | 78 |
| Do.. | 400 |  | 1 | 68 | 963.8 | 95 |
| Warsaw, $\frac{1}{2}$ mile west of... | 460 | 460 | 1 | 70 | 963.8 | 74 |
|  | 702 | 702 | 2 | 71 | ${ }^{063.5}$ | ${ }_{58}^{97}$ |
| Warsaw, 1 mile west of.............. | 300 |  | $1^{\frac{4}{4}}$ | ${ }_{69}^{69}$ | ${ }^{963} 8$ | 58 |
| Warsaw, near......... | 300 |  | 13 | ${ }_{69}^{68}$ | 963.8 | ${ }_{78}$ |
|  |  |  |  |  |  |  |
| Tuscaloosa County: | 210 | 210 | 38 |  | b63.3 | 57 |
| Ifuls station, i mile southwest of... | 200 | 200 | 30 | 66 | 663.3 | 74 |
| Tuscaloosa, 9 miles southwest of..... | 160 | 160 | 60 | 67 | $\square 63.3$ | 43 |
|  | 234 |  | $2{ }_{2}^{1}$ | 66 | ${ }^{\text {h63.1 }}$ | 81 |
| a Livingston average minus $0.5^{\circ}$. $b$ Tuscaloosa average plus $0.2^{\circ}$. |  | c Livingston average minus $0.3^{\circ}$. <br> $f$ Livingston average minus $0.4^{\circ}$. |  |  |  |  |
|  |  |  |  |  |  |  |
| c Union Springs average. ${ }^{\text {a }}$ Livingston average minus $0.1{ }^{\circ}$. |  | ${ }^{\circ} \mathrm{l}$ Tuscaloosa average for 36 | $o$ Livingston average. |  |  |  |

Temperatures of flows from wells that derive part or all of their water from unknown depths are reported as follows: Boguechitto, Dallas County, 460 feet deep, $68^{\circ}$; Boligee, Green County, 468 fect, $66^{\frac{1}{2}}$, and 420 feet, $66^{\circ}$; Montgomery, 550 feet, $68^{\circ}$, with some water from 480 feet.

## sUmmary.

Probably several of the wells represented in the table have admixtures of water from different depths. The well at Union Springs, in Bulloch County, is pumped with an output of 150 gallons a minute, and as it is cased to a depth of 843 feet its water probably indicates the underground temperature with fair accuracy; the rate of increase indicated is $1^{\circ}$ in 74 feet. The two deeper wells at Cahaba if cased to their bottoms indicate temperature increases of $1^{\circ}$ in 57 and 55 feet, but the record of the shallower well at Bell's Hotel, with a rate of $1^{\circ}$ in 45 feet, in some measure invalidates the data from the others. The 409 and 487 foot wells at Selma may have mixed flows. The exceptional temperature of $72^{\circ}$ reported for
the 470 -foot well may be a mistake. The Eutaw wells show rates of increase of $1^{\circ}$ in 27 to 66 feet. The Steeles Bluff well is cased to a depth of 360 feet. Its recorded rate of increase of $1^{\circ}$ in 74 feet is probably reliable if the thermometer reading is correct. This rate is subject to a small plus or minus correction due to uncertainty as to the mean annual temperature. The Hale County wells of which the depths to flowing water are recorded show considerable range in rate of increase, mostly $1^{\circ}$ in 34 to 54 feet.: The wells at Mobile indicate rates of $1^{\circ}$ in 67 and 71 feet. The temperature of the water from the Pittsboro and Millwood wells, $72^{\circ}$, is remarkably high for wells only 219 and 236 feet deep and indicates rates of $1^{\circ}$ in 30 and 79 feet, respectively. These figures might be modified slightly, however, if the exact mean annual temperature were known. The same is true of the ' 200 -foot well $4 \frac{1}{2}$ miles south of Eutaw, in Greene County. Hale County also shows some high rates of increase. The low rate of the 200 -foot well at Linden, in Marengo County, is a notable exception in the region southwest of Tuscaloosa. The Epes 750 -foot well, the Huntersfield well, and the Selma wells (except the 470 -foot well) are others having lower rates than the many wells that show increases of $1^{\circ}$ in 50 to 90 feet.

## GEOLOGIC RELATIONS.

All the wells given in the list are in the southern or Coastal Plain part of the State and penetrate sands, marls, and clays of Cretaceous and Tertiary age. These strata form a succession of sheets which dip southward at a low angle and mostly thicken in that direction. They lie on a basement of older granite, schist, and other rocks, which is nearly reached by the deeper wells in the northern part of the Coastal Plain district, among them the deeper wells in Tuscaloosa County, the wells 500 to 550 feet deep in Greene County, the deeper wells in Pickens County, the 930 -foot well 3 miles southeast of Epes, in Sumter County, and the 1,200 -foot well at Linden, Marengo County. These wells appear to show a wide range in rate of increase of temperature, the rate in the well at Linden, the deepest well reported, being very low- $1^{\circ}$ in 143 feet, if the temperature of the flow is given correctly. The group of wells in Bullock, Russell, and Barbour counties, most of which draw from younger formations, show no special variations. The temperature recorded for the wells in Selma, which draw from the upper part of the Tuscaloosa formation, is about $68^{\circ}$, indicating a low rate of increase. The wells at Mobile draw from higher beds and show a rate of $1^{\circ}$ in about 70 feet.

## ARIZONA.

## SOURCE OF INFORMATION.

Underground temperatures in Arizona have been determined by W. T. Lee, C. A. Fisher, and A. T. Schwennesen, of the United States Geological Survey. In 1904 Mr. Lee made for me a number of observations in wells and shafts at the Congress mine, near Phoenix, and in the Plateau region, and from 1913 to 1915 Mr . Schwennesen ${ }^{1}$ obtained temperatures in the San Simon Valley.

## PHOENIX.

Several moderately deep borings were made in Salt River Valley near Phoenix, most of them sunk to develop water supplies from the deep desert-valley fillings. In January, 1904, Mr. Lee ${ }^{2}$ carefully determined temperatures in these holes by sinking Darton thermometers with the results set forth below. The mean annual air temperature at Phoenix for 21 years, $69.4^{\circ}$, is used in the calculations of gradient.

Temperature in wells near Phoenix, Ariz.

| Well. | Depth (fect). | $\begin{aligned} & \text { Temper- } \\ & \text { ature } \\ & \left({ }^{\circ} \mathrm{F} .\right) . . \end{aligned}$ | ```Depth per degree of increase in temperature (feet).``` | Remarks. |
| :---: | :---: | :---: | :---: | :---: |
|  | 345 | 78.1 | 40 | Well 1,305 feet deep but clogged at |
| Murphy \& McQueen ranch............ <br> Consolidated Canal Co $\qquad$ | 150 | 77.6 | 18 | 3 349 feet. |
|  | 375 | 84.6 | 25 | Well 705 feet deep but clogged at |
|  | 150 | 81.1 | 13 | ) 375 fect. |
|  | 835 | 83 | 61 | All water cased off; thermometer down 20 hours. |
|  | 800 | 82.1 | 63 | Thermometer down 2 hours. |
|  | 700 | 81.9 | 56 | Do. |
| Sec. 30, T. 2 N., R. 4 E................ | 600 | 81.1 | 51 | Do. |
|  | 500 | 80.6 | 45 | Thermometer down 17 hours. |
|  | 400 | 80 | 38 | Thermometer down 7 hours. |
|  | 300 | 79.8 | 29 | Thermometer down 16 hours. |
|  | 200 | 79.1 | 21 | Jhermometer down 3 hours. |
| E. F. Kellner, sec. 6, T. 1 N., R. 1 E.. | 324 | 82.3 | 25 |  |
| Valley Grape Co., sec. 1, 'T. 1 S., R. 4 E. | 319 | 74.4 | 64 |  |

The increase is $3 \frac{1}{2}^{\circ}$ from 150 to 375 feet in the well of the Consolidated Canal Co., though it is only half a degree from 150 to 349 feet in the well at the Murphy \& McQueen ranch. Lee suggests that this difference may be due to freer circulation of the underflow at the latter place.

[^1]
## CONGRESS MINE.

The large Congress gold mine, south of Prescott, Ariz., which has been worked to a depth of 1,300 feet, was selected for a series of careful determinations of temperature. The tests were made in March, 1904, by W. T. L'ee, in shafts No. 6 and No. 3, which had not been worked for several months. A week in advance of the tests holes from 5 to 9 feet deep were drilled at various depths into the walls of side drifts 30 to 142 feet from these shafts but out of the main currents of ventilation. The thermometers were lightly covered and allowed to remain in the holes from 10 to 24 hours. The shafts are steep inclines, No. 6 being at an angle of about $45^{\circ}$ and No. 3 at about $21^{\circ}$, and as they extend downward the land above rises considerably. The depths given below are taken from a surface profile in which the elevations were determined barometrically by Mr. Lee.

Temperatures in shafts at Congress mine, near Prescott, Ariz.

|  | Vértical depth (feet). | Thermometer remained in hole (hours). | Temperature ( ${ }^{\circ} \mathrm{F}$.). | Depth per degree of increase in temperature (feet). |
| :---: | :---: | :---: | :---: | :---: |
| Shaft No. 3. | 205 | 23 | 68.1 |  |
| Do. | 687 | 101 | 71.4 | 156 |
| Do. | 1,125 | 102 | 74.6 | 167 |
| Do. | 1,285 | $22 \frac{1}{2}$ | 82.1 | 89 |
| Shaft No. 6. | 560 | 24 | 76.6 | 63 |
| Do. | 1,158 | 24 | 84.2 | 70 |

$a$ Based on mean annual temperature of $67.7^{\circ}$, the average for 1897 to 1904 at the mine. This is probably too high, as the mean at Wickenburg, about 1,000 feet lower in altitude, was $64.2^{\circ}$ for 1916.

The rate of increase in shaft No. 6 is $1^{\circ}$ in 78.7 feet between 560 and 1,158 feet. The rates in shaft No. 3 are $1^{\circ}$ in 146 feet between 205 and 687 feet, $1^{\circ}$ in 137 feet between 687 and 1,125 feet, $1^{\circ}$ in 141 feet between 205 and 1,125 feet, and $1^{\circ}$ in 21.3 feet between 1,125 and 1,285 feet. The rock is granite, mostly massive, but at the 1,285 -foot level in shaft No. 3 the observation was made 8 feet below a 2 -foot zone of crushed schistose material with slickensided surfaces indicating much movement. Probably this condition caused the much greater increase of temperature between 1,125 and 1,285 feet than in the intervals above and in shaft No. 6. Shaft No. 6 is all in massive granite down to the vein, which is cut at 1,158 feet. The temperature at 1,285 feet in shaft No. 3 was verified by a duplicate observation.

## PLATEAU REGION.

Several observations were made in borings along the Atchison, Topeka \& Santa Fe Railway in Paleozoic rocks in the Platoau region of north-central Arizona, resulting as follows:

Temperatures in wells alony Atchison, Topeka \& Santa Fe Railway in north-central Arizona.

| Well. | Depth (feet). | Temperature ( ${ }^{\circ}$ F.). | Mean annual air temperature ( ${ }^{\circ} \mathrm{F}$.) | Depth per degrec of increhse in temperature (fect). | Remarls. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Yucca. | \{ 1,004 | a90. 5 | b 67 | 45 | Standing water below 346 fect. Test by W. T. Lee, September, 1903. |
|  | 455 | 80.6 | 67 | 34 | No work for 3 months. Test by W. T. Lee, September, 1903. |
| Nelson.. | $\{1,043$ | 78.4 | c 52.3 | 50 | 553 feet of water in well. Test by W. T. Lec, Scptember, 1503. |
|  | ( 200 | 01.2 | 523 | 23 | No work for 4 months. Test by W. T. Lee, Scptember, 1903. |
| Seligman. | 700 | 84.2 | 52.3 | 22 | Thermometer down 2 hours. Well was being drilled all day and had been "shet" 8 hours before. Test hy C. A. Fisher, June, 1902. |
|  | ( 1,216 | 77.9 | 52.3 | 47.5 | Dry hole; tools in bottom ( 1,479 feet). No work for 24 hours. 'J'est by W. T. Lee, September, 1903. |
| Do... | 227 | 69.5 | 52. 3 | 13 | Old hole. Test by C. A. Fisher, June, 1802. |

a Duplicate tests, closely accordant.
b Mean of Needles, Catif., and Kingman, Ariz.; may be somewhat lower or higher.
c Seligman average for 10 years.
In the Yucca well the rate of increase was $1^{\circ}$ in 55.5 fcet between 455 and 1,004 feet, and in the Nelson well it was $1^{\circ}$ in 69.1 feet between 200 and 1,043 feet. Both holes are in nearly horizontal limestone of Carboniferous age and perhaps had reached the underlying Cambrian shale. The Seligman well was tested at two levels, first at 700 feet and 14 months later at 1,216 feet, where the temperature was $6.3^{\circ}$ cooler. At Seligman there is a mass of later Quaternary basalt several hundred feet thick which may retain some of its original heat and so cause the higher rate of increase near the surface. In another hole at this place 227 feet deep and entirely in the basalt the rate is still higher. Below the basalt the boring was in Carboniferous limestone in a shallow syncline.

## SAN SIMON VALLEY.

San Simon Valley, in the southeast corner of Arizona, contains many flowing wells 200 to 1,230 feet deep, drawing their supply from the thick valley fill of sand and clay. Mr. Schwennesen made the following tests.

Temperatures in wells in San Simon Valley, Ariz.

| Location. | Depth (feet). | $\begin{gathered} \text { Flow } \\ \text { (gallons } \\ \text { per } \\ \text { minute). } \end{gathered}$ | $\begin{gathered} \text { Tem- } \\ \text { perature } \\ \left({ }^{\circ} \mathrm{F} .\right) . \end{gathered}$ | Depth per degree of increase in tem(feet).a | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| T. 13 S. R. 29 E.: |  |  |  |  |  |
| NW. $\frac{1}{4}$ sec. 18. | $80_{0}$ |  | $\varepsilon 3$ | 48.6 |  |
| NE. $\frac{1}{2}$ sec. 24. | 960 | 26 | 105 | 24.2 |  |
| T. 13 S S, R. $30 \mathrm{E}$. : |  |  |  |  |  |
| NE. ${ }^{\frac{1}{4} \frac{1}{4} \text { sec. }}$ Se. 9. | 860 900 | 10 | 93, 109 | 27.6-32.5 | Two tests. <br> Do. |
| NW. $\frac{1}{1} \mathrm{sec} .11$. | 950 | 4 | 96 | 30.9 |  |
| NE .1 | 760 | 11 | 88 | 33.4 |  |
| SW. $\frac{1}{15}$ sec. 18. | 900 |  | 67 |  | Very small flow. |
| NW. 1 sec. 23. | 900 | 9 | 92 | 33.7 |  |
| NE. $\frac{1}{4}$ sec. 24. | 860 | 3 | 80 | 58.5 |  |
| NE. $\frac{1}{2 s e c .25 .}$ | 880 | 1 | 75 | 90.7 59 |  |
| NE. ${ }^{4}$ sec. 25. | 900 | 9 | 84 | 48.1 |  |
| NE. $\frac{1}{4} \mathrm{sec} .25$ | 935 | 1 | 84 | 50.0 |  |
| NW. $\frac{1}{4}$ sec. 26. | 860 | 1 | 78 | 67.7 |  |
| NW. $\frac{1}{4}$ sec. 30. | 930 | 10 | 105 | 23.4 | Two tests. |
| SW. ${ }^{\frac{1}{4}} \mathrm{sec} .17$ | 840 | 2 | 79 | 61.3 |  |
| NE. ${ }^{\text {a }}$ sec. 18. | 610 | 3 | 83 | 34.4 |  |
| SW. $\frac{1}{4}$ sec. 19. | 840 | 12 | 82 | 50.3 |  |
| NW. $\frac{1}{2}$ sec. 20. | 560 | 24 | 84 | 30.0 |  |
| NE. $\frac{1}{1}$ sec. 20. | 590 | 64 | 83 | 33.3 |  |
| SW. $\frac{1}{4}$ sec. 20 | 660 | 1 | 73 | 85.7 |  |
| $\mathrm{SE}_{\mathrm{SE}}+\frac{1}{2} \mathrm{sec} .20$. | 640 | 24 | 84 | 3.4 .2 |  |
| SE. 1 sec. 20. | 615 | 111 | $\stackrel{8}{8}$ | 34.8 |  |
| SE. ${ }_{\text {S }}$ sec. sec 21. | 550 600 | 58 12 | 80 72 | 37.4 90.0 |  |
| NW. 1 sec. 28. | 673 | 164 | 82 | 40.3 |  |
| NE. $\frac{1}{}$ sec. 28. | 549 | 51 | 82, 80 | 37.3-32. 3 | Do. |
| NE. $\frac{1}{4} \mathrm{sec} .28$. | 625 | 107 |  | 40.0 |  |
| NE. $\frac{1}{1}$ sec. 29. | 618 | 54 | 79, 78 | 48.7-45. 1 | Do. |
| T. 13 SW, R. 31 E.: |  |  |  |  |  |
| SW. ${ }^{\text {S }}$, sec. 30. | 1,008 | $\stackrel{5}{5}$ | 79 | 73.6 |  |
| SW. $\frac{1}{4}$ sec. 31. | , 884 | 36 | 84 | 47.3 |  |
| SE. ${ }^{4}$ sec. 31. | 850 | 36 | 82 | 51.0 |  |
| SE. ${ }_{\text {, }}^{2}$ sec. 33. | 590 | 200 | 80 | 40.0 |  |
| $\mathrm{SE}_{\mathrm{SE}}$. ${ }^{\text {sec. }}$ Sec. 33. | 690 | 21 | 81 | 44.0 |  |
|  | 663 | 170 | 81 | 42.0 |  |
| SW. $\frac{1}{4}$ sec. 1. | 920 |  | 81 | 58.6 |  |
| NE. $\frac{1}{3}$ sec. 12 | 1,040 | 45 | 82 | 62.3 |  |
| NF. $\frac{1}{2} \mathrm{sec} .12$ | 880 | 1 | 78 | 69.3 |  |
| SE. 12 sec .13. | 743 |  | 72 | 111.0 |  |
| $\text { T. } 14 \mathrm{~S} . \mathrm{R} .31 \mathrm{E} .:$ | 626 | 95 |  |  |  |
| SW. ${ }^{\frac{1}{4} \text { sec. } 3 \text { S. }}$ | 704 | 166 | 75 | 42.6 72.6 |  |
| SW. ${ }^{\frac{1}{4} \text { sec. } 4 .}$ | 570 | 48 | 76 | 53.2 |  |
| SW, $\frac{1}{}$ see. 5 | 770 | 19 | 78 | 60.6 |  |
| NW. $\frac{1}{4}$ sec. 6 . | 900 | 2 | 78 | 70.8 |  |
| SW. ${ }^{\frac{1}{3}}$ sec. 7. | 760. | $\cdots 45$ | 77 | 65.0 |  |
| NE. ${ }_{\text {NW }} \frac{1}{2}$ sec. 8. | 796. 800 | 8 46 | 78 80 | 62.7 <br> 54.4 |  |
| SE. $\frac{1}{4}$ sec. 9. | 775 | 107 | 76 | 73.1 |  |
| NF. $\frac{1}{4} \mathrm{sec} 10$ | 603 | 222 | 80 | 41.0 |  |
| NW. $\frac{1}{1}$ sec. 10. | 624 | 24 | 74 | 71.7 |  |
| SE. $\frac{1}{4} \mathrm{sec} 10$. | 650 | 260 | 78 | 51.2 |  |
| SW. $\frac{1}{1}$ sec. 11. | 433 | 28 | 76 | 40.4 |  |
| SW. 1 sec. 11. | 730 | 1 | 72 | 109.0 |  |
| SW. ${ }_{\text {SW }}$ sec. 13. | 490 |  | 80 | 33.3 |  |
|  | 530 440 | 3 | 76 | 49.5 |  |
| NW. $\frac{1}{4}$ sece. 15. | 820 | $\begin{array}{r}60 \\ 206 \\ \hline\end{array}$ | 75 79 | 44.4 60.0 |  |
| NW. $\frac{1}{4}$ sec. 15 | 790 | 274 | 80 | 53.7 |  |
| SW. $\frac{1}{4}$ sec. 15. | 726 | 125 | 83 | 41.0 |  |
| SE. $\frac{1}{4}$ sec. 16. | 715 | 31 | 79 | 52.2 |  |
| NE. $\frac{1}{1}$ sec. 17. | 830 1 140 | 18 | 84 | 45.4 |  |
| NW. $\frac{1}{4}$ sec. 19. | 1,140 | 1 | 70 77 | 242.5 74.9 |  |
| NE. ${ }_{4}$ | 730 |  | 84 | 74.9 39.0 |  |
| SE. ${ }_{4}^{4}$ sec. 21. | 800 | 34 | 85 | 40.7 |  |
| NE. 4 sec. 22. | 620 | 28 | 78 | 48. 8 |  |
| SW. ${ }^{1}$ Sec. 22. | 750 | 40 | 88 | 38.1 |  |
|  | 770 | 120 250 | 83 84 | ${ }_{32.3}^{43.5}$ |  |
| SW. $\frac{1}{4}$ sec. 23. | 750 |  | 83 | 42.4 |  |

Temperatures in wells in San Simon Valley, Ariz.-Continued.

| Location. | Depth (feet). |  | $\begin{aligned} & \text { Tem- } \\ & \text { perature } \\ & \left.0^{\circ} \mathrm{F} .\right) . \end{aligned}$ | Depth per degree of increase in temperature (feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| T. 14 S., R. 31 E.:-Continued. |  |  |  |  |  |
| SE. ${ }^{1}$ sce. 23. | 620 | 180 | 80 | 42.2 |  |
| SE. $\frac{1}{4}$ sec. 23. | 720 |  | 80 | 49.0 |  |
| SE. ${ }^{\frac{1}{4}}$ sec. 23. | 580 | 90 | 79 | 42.3 |  |
| NW. $\frac{1}{4}$ sec. 24 | 590 | 87 | 77 | 50.5 |  |
| NE. $\frac{1}{4}$ sec. 25. | 640 | 134 | 80 | 43. 5 |  |
| - SE. ${ }^{\text {a }}$ sec. 25. | 660 | 137 | 83 | 37.3 |  |
| - NW. $\frac{1}{2}$ sce. 26. | 735 | 22 | 81 | 46.7 |  |
| NW. $\frac{1}{2}$ sec. 20. | 920 | 25 | 84 | 49.2 |  |
| NW. $\frac{1}{3}$ sec. 27. | 740 | 3 | 83 | 41.8 |  |
| $\text { T. } 14 \text { SW. R. } 32 \text {. } 32 .$ |  |  |  |  |  |
| NW. ${ }^{4}$ sscc. 19. | 441 | 300 | 75 80 | 40.1 30.0 |  |
| SW. ${ }^{\text {d }}$ sec. 19. | 207 | 254 | 73 | 26.9 |  |
| SWW | 350 | 34 | 70 | 74.5 |  |
| NW. ${ }^{\text {a sec. 31 }}$ | 430 | 20 | 78 | 33.8 |  |
| $\text { T. } 15 \text { SW. R. } 32 \text { E. }$ | 603 |  | 76 | 56.3 |  |

## ARKANSAS.

## TEMPERATURES.

Very few temperature determinations from Arkansas are available The following are taken mostly from a report by Veatch: ${ }^{1}$

Temperatures in wells in Arkansas.

a Centerpoint average for 16 years.
© Greenville, Miss., average minus $0.1^{\circ}$.
c Camden average for 30 years minus $0.3^{\circ}$.
$d$ Amity average for 24 years.
e Average for 24 years.
$f$ Warren average for 34 years.
$g$ Average for 32 years.
$h$ Camden average plus $0.7^{\circ}$.
$i$ A verage for 21 years.

[^2]
## SUMMARY.

The data for Allbrook, Hot Springs, Pine Bluff, Stamps, and Texarkana apparently are not valid, probably because the true source of the flow is not given. The temperatures in other wells appear to be consistent. These wells are in the southern quarter of the State and reach water-bearing strata in the great Tertiary sedimentary series of the Mississippi embayment.

## GEOLOGIC RELATIONS.

The Gurdon well draws from the Nacatoch sand in the Upper Cretaceous; the Arkansas City well, the shallow well at Wilmar, and the Fordyce well, all with high rates of increase in temperature, draw from Eocene beds at the horizon of the Yegua ("Cockfield") formation, about 1,500 feet higher; the Monticelle well and the 455 -foot well at Wilmar, both with high rates of increase, draw from coarser deposits of Eocene age not far below the Yegua horizon.

## CALIFORNIA.

## SOUTHERN CALIFORNIA.

W. C. Mendenhall, ${ }^{1}$ in studying the hydrology of southern California, determined the temperature of water flowing from many artesian and other wells and from several hot springs. The following list sets forth the principal data obtained in San Bernardino Valley:

Temperatures in wells in San Bernardino Valley, Calif.

| Well. | Depth (feet). | $\begin{gathered} \text { Tempera- } \\ \text { ture } \\ \left({ }^{\circ} \mathrm{F} .\right) . \end{gathered}$ | Mean annual air temperature ( ${ }^{\circ} \mathrm{F}$.). | Depth per degree of increase in temperaturo (feet). |
| :---: | :---: | :---: | :---: | :---: |
| Riverside, Waterman Avenue. | 984 | 75. | - 63. | 82 |
| Mound City: |  |  |  |  |
| $\frac{1}{2}$ mile northwest of. | 784 | 74 | a 63 . | 71 |
| $\frac{1}{2}$ mile west of. | 656 | 75 | $a 63$ | 55 |
| 1 mile west of. | 125 | 68 | a 63 | 25 |
| $\frac{1}{2}$ mile northwest of | 582 | 74 | a 63 | 53 |
| $\frac{2}{2}$ mile west of...... | 534 | 74 | a 63 | 49 |
| 1 mile northeast of | 517 | 73 | a 63 | 52 |
| Do... | 506 | 70 | a 63 | 72 |
| 2 miles north by east of | 472 | 70 | a 63 | 67 |
| $\frac{1}{2}$ mile northwest of.... | 482 | 74 | a 63 | 44 |
| $\frac{1}{2}$ mile northwest of. | 320 | 73 | a 63 | 32 |
| 2 miles northwest of | 83 | 68 | b 62.7 | 16 |
| $1 \frac{1}{2}$ miles north of... | 196 | 70 | \% 62.7 | 27 |
| Do...... | 190 | 70 | $\checkmark 62.7$ | 26 |
| Do. | 142 | 70 | b 62.7 | 19 |
| 1 ? miles north of | 148 | 68.5 | b 62.7 | 25 |
| 1. Do.......... | 141 | 67 | $b 62.7$ | 33 |
| Do. | 181 | 68 | b 62.7 | 34 |
| 2 miles north of | 472 | 70 | ${ }^{6} 62.7$ | 65 |
| Urbita, new woll. | 740 | 78 | c 62.5 | 48 |

[^3][^4]Temperatures in wells in San Bernardino Valley, Calif.-Continued.

| Woll. | Depth (feet). | Tempera- ture ( ${ }^{\circ} \mathrm{F}$.). | Mean annual air tem$\underset{\left({ }^{\circ} \mathrm{F} .\right)}{ }$ perature | Depth perdegree ofincrease in temperature (fect). |
| :---: | :---: | :---: | :---: | :---: |
| San Bernardino: |  |  |  |  |
| Payne well, 2 miles southeast of.. | 642 | 112 | $a 62.5$ | 13 |
| City well on Antell tract, 2 miles east of | 614 | 69 | $a 62.5$ | 94 |
| 1 mile east of....... | 544 | 68 | $a 62.5$ | 99 |
| 2 milos east of. | 682 | 68 | $a 62.5$ | 129 |
| 3 milos northeast of | 460 | 72 | a 62.5 | 48 |
| 2 miles south of. | 225 | 71 | a 62.5 | 27 |
| Do... | 121 | 98 | $a 62.5$ | 3 |
| Do.. | 158 | 85 | a 62.5 | 7 |
| 2 miles north by east of | 169 | 76 | $a^{6} 2.5$ | 13 |
| Earlem Springs: |  |  |  |  |
| 1 mile scuth of. | 451 | 90 | a 62.5 | 17 |
| 1 mile west cf. | 231 | 72 | a 62.5 | 24 |
| 14 miles west of. | 185 | 80 | $a 62.5$ | 11 |
| 1 mile west by south of | 150 | 68 | a 62.5 | 27 |
| $\frac{1}{3}$ mile north of....... | 300 | 115 | a 62.5 | 6 |
| \% mile northwest of. | 68 | 86 | $a 62.5$ | 3 |
| 2 miles east of | 92 | 73 |  |  |
| Do........ | 81 | 73 | 63 | 8 |
| Do. | 68 | 71 | 63 | 8 |

© San Bornardino average for 25 years.
The distribution, depths, and geothermal gradients of most of these wells are shown in figure 2.

The data in the table present variations that are difficult to understand. Some of them are not reliable, because flows from higher horizons may have ingress to the well, and in a few wells the outflow is small, but these two conditions cause diminished temperatures. The materials penetrated by the borings are a thick series of Quaternary sand, gravel, and clay lying horizontal and constituting a wide desert plain. The conditions are favorable for rapid circulation of water, so that the mean annual air temperature $62.5^{\circ}$ should be expected to extend some distance below the surface. On the edge of the basin and at Harlem Springs there are thermal springs, and in a local crumple of the beds passing through Bunker Hill warm waters come very near the surface. The areas of high temperature are probably due to seepage of warm spring waters along or near recent faults. Admixture with the shallow underflow causes the great diversity of temperatures shown.

The 642 -foot boring 2 miles southeast of San Bernardino, whose water has a temperature of $112^{\circ}$, appears to have tapped a particularly warm spot, although shallower borings a short distance southwest show nearly the same rate of temperature increase. The water of the Urbita Springs has a temperature of $102^{\circ}$ to $105^{\circ}$. Several wells near the Bunker Hill anticlinal ridge show temperatures of $85^{\circ}$ to $105^{\circ}$, and the shallow wells of the Riverside Water Co. along the north bank of Santa Ana River show temperatures of $71^{\circ}$ to $73^{\circ}$. Deeper wells, 320 to 656 feet deep, near Loma Linda show temperatures
of $70^{\circ}$ to $75^{\circ}$, but the rate of increase is not so high in these wells as in the other wells just mentioned.


Figure 2.-Map of part of the San Bernardino Valley, Calif., showing location of wells and indicating geothermal gradient. Depths of wells are represented by gothic figures, and italic figures show the depth in feet per degree of increase in temperature. Data from Mendenhall's report.

Mendenhall ${ }^{1}$ has recorded the temperatures in many wells in the Anaheim and Santa Ana districts of the eastern coastal-plain region of southern California, and, although most of them are too shallow to indicate the geothermal gradient, they present some interesting features. Most of the wells are from 80 to 200 feet deep, and the waters show unaccountable variations in temperature, mostly from $61^{\circ}$ to $65^{\circ}$ and a few to $70^{\circ}$. Of the few deeper wells covered by the

[^5]records one flowing well 397 feet deep in the San Joaquin district showed a temperature of $79^{\circ}$, which indicates a rate of increase of about $1^{\circ}$ in 23 feet. Another well near by, 165 feet deep, with a temperature of $81^{\circ}$, indicates a still higher rate. The temperature of water from a 386 -foot well was $67^{\circ}$, a 390 -foot well $70^{\circ}$, a 264 -foot well $71^{\circ}$, a 360 -foot well $71^{\circ}$, a 336 -foot well $71^{\circ}$, a 400 -foot well $68^{\circ}$, two 450 -foot wells $66^{\circ}$, and a 118 -foot well $71^{\circ}$. Some of the wells showing the higher temperatures are in groups or districts, but others appear to be irregularly distributed.
Temperatures in the Cucamonga and Pasadena region have been recorded by Mendenhall, as follows:

Temperatures in wells in Cucamonga and Pasadcna quadrangles, Calif.

| Location. | $\begin{aligned} & \text { Depth } \\ & \text { (feet). } \end{aligned}$ | $\begin{gathered} \text { Flow } \\ \text { (gallons } \\ \text { per } \\ \text { minute). } \end{gathered}$ | $\begin{gathered} \text { Tem- } \\ \text { peraturo } \\ (\mathrm{F}) . \end{gathered}$ |  | Depth per degree of increase in tem- perature (fect). |
| :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |  |
|  |  |  |  |  |  |
| Do... | ${ }_{456}^{479}$ | $\begin{array}{r}280 \\ 56 \\ \hline\end{array}$ | 72 72 | 62.5 62.5 | ${ }_{49}^{50}$ |
| Do. | ${ }_{433}^{486}$ | ${ }_{33}^{56}$ | ${ }_{72}$ | 62.5 | ${ }_{48}$ |
| Do. | ${ }_{666}^{428}$ | 45 | ${ }_{64}^{73}$ | 62.5 | ${ }^{41}$ |
| ${ }_{\text {D }} \mathrm{D}$ | 638 |  | 62 | 62.5 |  |
| Do.. | ${ }_{624}$ | 850 | 62 | 62.5 |  |
| Do. | 640 | 850 | 62 | 62.5 |  |
| Do. | 540 649 | 850 1.675 | ${ }_{70}^{62}$ | ${ }_{62}^{62.5}$ |  |
| Santa Ana del Chino. | ${ }_{375}$ |  | 64 | 62.5 | 250 |
| Do. | 126 |  | ${ }^{67}$ | ${ }^{62.5}$ | 30 |
| Do. | 331 | 900 | ${ }_{66}^{69}$ | 62.5 | - 16 |
| Do. | 347 | ${ }_{225}$ | ${ }_{65}^{60}$ | 62.5 | 140 |
| San Jose... | ${ }_{400}$ | 500 | ${ }^{62}$ | 62.5 |  |
|  | 385 367 | 340 280 | 64 70 | 62.5 62.5 | 49 |
|  |  |  |  |  |  |
|  |  |  |  |  |  |
| ${ }_{\text {DO }}$ |  |  |  |  | ${ }_{36}^{50}$ |
| Sec. 9, T. $2 \mathrm{~S} ., \mathrm{R} .13$ w | ${ }_{210}$ | ....... | 65 | 60.3 |  |
|  |  |  |  |  |  |
| Do.. | 175 |  | ${ }_{68}^{62}$ | 60.3 | 23 |
| Do. | 202 |  |  | 60.3 |  |
|  |  |  |  |  |  |
| Do. | 492 | 280 |  | 62.5 | 197 |
| Do. | 140 | 560 | ${ }_{71}^{67}$ | 62.5 | $\begin{array}{r}33 \\ 18 \\ \hline\end{array}$ |
|  | 130 <br> 156 <br> 158 | ${ }_{45}^{22}$ | ${ }_{71}^{71}$ | 62.5 <br> 62.5 | 18 15 |
|  |  |  |  |  |  |
| San Francisquito....... | 130 | 1,750 | 64 | 62.5 | 87 |

[^6]In this region other wells from 113 to 300 feet deep show temperatures of $62^{\circ}$; wells 68 and 172 feet deep, $63^{\circ} ; 280$ feet, $64^{\circ}$; 250 feet, $65^{\circ} ; 60$ and 160 feet, $66^{\circ} ; 147,160$, and 300 feet, $67^{\circ} ; 139$, $160,180,250$, and 255 feet, $68^{\circ}$; and 140 and 220 feet, $69^{\circ}$. These figures indicate great irregularity in the rate of increase at slight depths.

## SAN LUIS OBISPO.

A flow of sulphur water from a 928 -foot well at San Luis Obispo has a temperature of $103^{\circ}$. As the mean annual air temperature at this place is $57.2^{\circ}$, this indicates a rate of increase of $1^{\circ}$ in 20 feet.

SAN JOAQUIN VALLEY.

Mendenhall and Dole ${ }^{1}$ have given records of many observations of temperatures of wells in San Joaquin Valley. Most of the wells are shallow," but others are deep and flowing, so that their temperatures throw much light on the rate of increase. These wells are all in the filling of loam, clay, sand, and gravel, occupying a broad valley between mountain ranges of older rocks. These deposits are known to be more than 2,000 feet thick in places. It is believed that none of the wells here recorded have reached their base. The data which appear to be most useful are given in the following list:

Temperatures in wells in San Joaquin Valley, Calif.

| Location. | Depth (feet). | Flow (gallons per minute). | Temperature ( $\left.{ }^{\circ} \mathrm{F}.\right)$. | Mean annual air temperature ( ${ }^{\circ} \mathrm{F}$.). | Depth per degree of increase in temperaturo (feet). |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Fresno County: |  |  |  | . |  |
| Sec. 22, T. 13 S., R. 14 E. | 700 | 70 | 78 | $a 63.2$ | 47 |
| Sec. 31, T. 13 S., R. 15 E. | 640 | 80 | 78 | 63.2 | 43 |
| Sec. 2, T. 14 S., R. 15 E.. | 750 | 225 | 78 | 63.2 | $50 \frac{1}{2}$ |
| Sec. 10, '1. 14 S., R. 16 E. | 800 | 180 | 78 | 63.2 | 54 |
| Sec. 12, T. 15 S., R. 16 E. | 800 | 112 | 77 | 63.2 | 58 |
| Sec. 25, T. 15 S., R. 16 E . | 825 |  | 76 | 63.2 | $64 \frac{1}{2}$ |
| Sec. 19, T. 15 S., R. 16 E. | 700 | 22 | 79 | 63.2 | $44^{2}$ |
| Sec. 11, T. 15 S., R. 15 E. | 733 | 125 | 76 | 63.2 | 57 |
| Sec. 25, T. 15 S., R. 17 E. | 750 | 345 | 80 | 63.2 | 44 |
| Sec. 14, T. 16 S., R. 17 E. | 690 | 90 | 75 | 63.2 | $58 \frac{1}{2}$ |
| Sec. 2 , T. $17 \mathrm{~S} ., \mathrm{R} .18 \mathrm{E}$. | 1,200 |  | -78 | 63.2 | 81 |
| Sec.17, T. 17 S., R. 19 E. | 700 | 190 | 77 | 63.2 | 502 |
| Sec. 21, T. 17 S., R. 19 E. | 1,500 | 260 | 82 | 63.2 | 79. |
| Sec. 28, T. $17 \mathrm{~S} ., \mathrm{R} .19 \mathrm{E}$ | 1,200 | 168 | 80 | 63.2 | 71 |
| Sec. 1, T. 17 S., R. 20 E . | 700 | 11 | 80 | 63.2 | $41{ }^{1}$ |
| Sec. 36, 'T. 18 S., R. 16 E. | 1,178 | 80 | 80 | 63.2 | 70 |
| Sec. 5, T. $19 \mathrm{~S} ., \mathrm{R} .20 \mathrm{E}$. | 500 | 35 | 76 | 63.2 | 39 |
| See. 28, T. 18 S., R. 20 E. | 700 | 225 | 76 | 63.2 | 54. |
| Kern County: |  |  |  |  |  |
| Sec. 20, T. 25 S., R. 24 E. | 800 | 2,900 | 78 | b 62.2 | $50 \frac{1}{2}$ |
| Sec. 26, T. 25 S., R. 24 E. | 995 | 45 | 82 | 62.2 | 50 |
| Sec. 34, T. 25 S., R. 24 E. | 647 | 230 | 78 | 62.2 | 41 |
| Sec. 3, T. 26 S., R. 24 E.. | 500 |  | 80 | 62.2 | 28 |
| Sec. 11, T. 26 S., R. 24 E. | 1,000 |  | 82 | 62.2 | $50 \frac{1}{2}$ |
| Sec. 31, T. 25 S., R. 25 E. | 700 | 155 | 79 | 62.2 | 41 |
| Sec. 17, T. 26 S., R. 24 E. | 512 | 345 | 79 | 62.2 | $30_{2}$ |
| Sec. 19, T. 26 S., R. 24 E. | 480 | 580 | 78 | 62.2 | 30 |
| Sec. 34, T. 26 S., R. 23 E. | 369 | 620 | 78 | 62.2 | 23 |
| Sec. 32, T. 26 S., R. 23 E. | 284 | 1,330 | 78 | 62.2 | 18 |
| Sec. 34, T. 26 S., R. 22 E . | 320 | -975 | 76 | 62.2 | 23 |
| Sec. 6, T. 27 S., R. 23 E. | 800 | 650 | 82 | 62.2 | 40 |
| Sec. $10, \mathrm{~T} .27$ S., R. 23 E | 423 | 70 | 76 | 62.2 | 55 |
| Sec. 3,'T. 27 S., R. 25 E.. | 700 |  | 78 | 62.2 | 44 |
| Sec. 5, T. 28 S., K. 23 E. | +600 | 1,160 | 76 | 62.2 | 40 |
| Sec. 24, T. 28 S., R. 23 E. | 390 | - 515 | 76. | 62.2 | 28 |
| Sec. 25, T. 28 S.; R. 23 E. | 450 | 335 | 73 | 62.2 | 41 |

a Fresno average for 30 years
$b$ Wasco mean annual average for 17 years.

[^7]Tcmperatures in wells in San Joaquin. Valley, Calif.-Continued.

|  |  |
| :--- | :--- | :--- | :--- | :--- | :--- | :--- |
|  |  |

[^8]Temperatures in wells in San Joaquin Valley, Calif.-Continued.

| Location. | Depth (feet). | $\underset{\substack{\text { Flow } \\ \text { (gallons } \\ \text { per }}}{ }$ minute). | Temperature ( ${ }^{\circ} \mathrm{F}$.). | Mean annual air temperature ( ${ }^{\circ} \mathrm{F}$.). | Depth degree of increase in temperature (feet). |
| :---: | :---: | :---: | :---: | :---: | :---: |
| San Joaquin County-Continued. |  |  |  |  |  |
| Campo de los Franceses.... | 1,045 | Few. | 79 | 60.1 | 55 |
|  | 1,128 | 22 |  | 60.1 | 47 |
| Sec. 22, T. 1 S., R. 6 E | 11,400 | 100 | 84 84 | 60.1 | 46 59 |
| Sec. 8, T. IS.,R. 1 E. | 1,250 |  | 77 | 60.1 |  |
| Sec. 25, T. 1 S., R. 6 E | 1,200 | 90 | 74 | 60.1 | 86 |
| Tulare County: |  |  |  |  |  |
| Sec. 30, T. 18 S., R. 24 E | 800 | 11 | 78 | $a 61.6$ | 48 |
| Sec. 29, T. 19 S., R. 24 E. | 420 | 90 | 75 | 61.6 | 31 |
| Sec. 31, T. 19 S., R. 24 E | 430 | 60 | 76 | 61.6 | 30 |
| Sec. 7, T. 20 S., R. 24 E. | 470 |  | 74 | ${ }^{6} 63.2$ | 4312 |
| Sec. 16, T. 20 S., R. 24 E | 462 |  | 74 | 63.2 | 43 |
| Sec. 10, T. 20 S., R. 23 E | ${ }^{505}$ | 70 | 74 | 63.2 | $46 \frac{1}{2}$ |
| Sec. 11, T. 20 S., R. 23 E | 1,150 | 230 | 75 | 63.2 | 972 |
| Sec. 5, T. 20 S., R. 24 E. | - 430 | 270 | 73 | 63.2 | 44 |
| $\mathrm{Sec}_{\text {Sec }} 32 \mathrm{~T}$ T. $20 \mathrm{~S} ., \mathrm{R} .24 \mathrm{E}$ | 1,000 | 11 45 | 72 | 61.3 | 114 |
| Sec. 25, T. 20 S., R. 23 E. | 640 | $\stackrel{45}{45}$ | 74 | 63.2 63.2 | $4{ }^{42}$ |
| Sec. 25, T. 20 S., R. 23 E | 787 | 270 | 76 | 63.2 | 61 |
| Sec. 35, T. 20 S., R. 23 E | 600 | 540 | 74 | 63.2 | $55 \frac{1}{2}$ |
| Sec. 22, T. 20 S., R. 23 E | 418 | 56 | 73 | 63.2 | 43 |
| Sec. 28, T. 20 S., R. 23 E. | 521 | 80 | 74 | 63.2 | 48 |
| Sec. 32, T. 20 S., R. 24 E | 530 | 56 | 74 | 63.2 | 49 |
| Sec. 32, T. 20 S., R. 23 E | 800 | 325 | 75 | 63.2 | 68 |
| Sec. 32, T. 20 S., R. 23 E | 868 | 270 | 76 | 63.2 | 68 |
| Sec. 8,'T. 21 S.,'R. 23 E. | 954 | 450 | 78 | 64.8 | 72 |
| Sec. 3, T. 21 S., R. 23 E | 900 | 475 | 75 | 64.8 | 88 |
| Sec. 3, T. 21 S., R. 23 E. | 750 | 270 | 74 | 64.8 | 81 |
| Sec. 10, T. 21 S., R. 23 E. | 800 | 225 | 74 | 64.8 | 87 |
| Sec. 10, T. 21 S., R. 23 E | 778 | 425 | 75 | 64.8 | 72 |
| Sec. 26, T. 21 S., R. 24 E | 600 | 22 | 72 | 64.8 | 83 |
| Sec. 4, T. 22 S., R. 24 E. | 600 | 134 | 72 | 64.8 | 83 |
| Sec. 23, T. 21 S., R. 24 E | 500 |  | 70 | 64.8 | 96 |
| Sec. 28, T. 21 S., R. 24 E. | 400 | 80 | 72 | 64.8 | 551 |
| Sec. 29, T. 21 S., R. 24 E. | 500 | 100 |  | 64.8 | 70 |
| Sec. 31, T. 21 S., R. 24 E. | 630 | 56 | 71 | 64.8 | 1013 |
| Sec. 17, T. 22 S., R. 23 E. | 815 | 340 |  | 64.8 | 73 |
| Sec. 17, T. 22 S., R. 23 E. | 918 | 875 | 78 | 64.8 | 69 |
| Sec. 21, T. 22 S., R. 23 E. | 800 | 560 | 78 | 64.8 | 60. |
| Sec. 30, T. 22 S., R. 23 E | 1,000 | 850 | 81 | 64.8 | 61 |
| Sec. 30, T. 22 S., R. 23 E | 1,200 | 340 | 77 | 64.8 | 982 |
| Sec. 29, T. 22 S., R. 23 E | 700 |  | 68 | 64.8 | 218 |
| Sec. 29, T. 22 S., R. 23 E. | 830 | Few. |  | 64.8 | 74 |
| Sec. 8, T. 22 S., R. 24 E. | 606 | 180 | 70 | 64.8 | 116 |
| Sec. 20 , T. 22 S., R. 24 E. | 606 | 270 | 70 | 64.8 | 116 |
| Sec. 10, T. 22 S., R. 24 E | 435 |  | 69 | 64.8 | 104 |
| Sec. 10, T. 22 S., R. 24 E | 480 | 150 | 69 | 64.8 | 114 |
| Sec. 22, T. 22 S., R. 24 E | 600 | 180 | 70 | 64.8 | 115 |
| Sec. 22, T. 22 S., R. 244 E | 1,400 | 56 | 71 | 64.8 | 226 |
| Sec. 22, T. 22 S., R. 24 E | 450 | 35 | 69 | 64.8 | 107 |
| Sec. 24, T. 22 S., R. 24 E | 500 | 200 | 70 | 64.8 | 96 |
| Sec. 23, T. 22 S., R. 24 E. | 1,100 | 400 | 70 | 64.8 | 211 |
| Sec. 13 , T. 22 S., R. 24 E | 1,043 | 35 | 72 | 64.8 | 145 |
| Sec. 3, 'T. 23 S., R. 25 E. | 680 |  |  |  |  |
| Sec. 7, T. 23 S., R. 25 E. | 500 | 90 | 78 | 64.8 | 38 |
| Sec. 7, T. 23 S., R. 25 E. | 800 | 35 | 77 |  | 65 |
| Sec. 8, T. 23 S., R. 25 E. | 1,387 | 450 | 79 | 64.8 | 98 |
| Sec. 18, T. 23 S., R. 25 E | 500 | 35 | 76 | 64.8 | 45 |
| Sec. 11, T. 23 S., R. 24 E. | 1,142 |  | 80 | 64.8 | 75 |
| Sec. 11, T. 23 S., R. 24 E | , 815 |  | 70 | 64.8 | 157 |
| Sec. 11, T. 23 S., R. 24 E | 1,000 |  | 78 | 64.8 | 76 |
| Sec. 35, T. 22 S., R. 24 E | 900 | 125 | 70 | 64.8 | 173 |
| Sec. 34, T. 22 S., R. 24 E | 582 | 205 | 75 | 64.8 | 57 |
| Sec. 3, T. 23 S., R. 24 E. | 600 | 360 | 74 | 64.8 | 65 |
| Sec. $9, \mathrm{~T}, 23 \mathrm{~S}$., R. 24 E | 1,000 | 850 | 18 | 64.8 | 76 |
| Sec. 16, T. 23 S., R. 24 E | 1,100 | 2,300 | 80 | 64.8 | 72 |
| Sec. 3, T. 23 S., R. 23 E. | 1,103 | 465 | 78 | 64.8 | 83 |
|  | 1,025 | 215 | 78 | 64.8 | 77 |
| Sec. 33, T. 23 S., R. 25 E | 1,300 | 245 | 75 | 64.8 | 127 |

[^9]The water in many shallow wells in San Joaquin Valley has temperatures ranging from $65^{\circ}$ to $72^{\circ}$, or considerably higher than the mean air temperature.

## INDIO BASIN.

Mendenhall ${ }^{1}$ has given the temperatures of flow from many of the artesian wells in the Indio Basin. They all draw from gravel and sand in the deep Quaternary valley fill, the bottom of which has not been reached by any of the borings.

Temperatures in wells in Indio Basin, Calif.

|  | Location. | Depth (feet). | Flow (gallons per minute). | Temperature ( $\left.{ }^{\circ} \mathrm{F}.\right)$. |
| :---: | :---: | :---: | :---: | :---: |
| Scc. 23, T. 6 S., R. 8 E |  | 499 | 135 | 74 |
| Sec. 23, T. 6 S.,', R. 8 E |  | 509 | 260 | 74 |
| Sec. 23, T. 6 S., R. 8 E |  | 497 | 225 | 74 |
| Sec. 23, T. 6 S., R. 8 E |  | 512 | $225+$ | 76 |
| Sec. 23, T. 6 S., R. 8 E |  | 526 | 58 | 74 |
| Sec. 25, T. 6 S., R. 8 E |  | 538 |  | 74 |
| Sec, 26, T. 6 S., R. 8 E |  | 487 | 170 | 74 |
| Sec. 26, T. 6 S., R. 8 E |  | 514 | 170 | 74 |
| Sec. 26, T. 6 S., R. 8 E |  | 601 | 90 | 74 |
| Sec. 23, T. 6 S., R. 8 E |  | 519 | - 170 | 74 |
| Sec. 22, T. 6 S., R. 8 E |  | 506 |  | 74 |
| Sec. 22, T. 6 S., R. 8 E. |  | 498 | 280 | 74 |
| Sec. 16, T. 7 S., R. 9 E |  | 484 | 225 | 73 |
| Sec. 18, T. 7 S., R. 9 E |  | 525 |  | 76 |
| Sec. 8, T. 7 S., R. 9 E. |  | 547 | 125 | 77 |
| Sec. 8, T. 7 S., R. 9 E |  | 531 | 150 | 77 |
| Sec. 12, T. 5 S., R. 7 E |  | 165 |  | 74 |
| Sec. 23, T. 5 S., R. 7 E |  | 46.5 | 22 | 73 |
| Sec. 13, T. 5 S., R. 7 E |  |  |  | 78 |
| Sec. 13, T. 5 S S., R. 7 R 7 E |  | 480 |  | 78 |
|  |  | 652 | 45 | 74 |
| Sec. 34, T. 6 St, 6 S., R. ${ }^{\text {R }} 8.8$ |  | 510 |  |  |
| Sec. 16, T. 7 S., R. 8 E |  | 480 | 383 | 81 |
| Sec. 18, T. 7 S., R. 9 E |  | 380 | 70 | 74. |
| Sec. 12, T. 7 S., R. 8 E |  | 409 | 157 | 76 |
| Sec. 12, T. 7 S., R. 8 E |  | 480 | 280 | 74 |
| Sec. 17, T. 6 S., R. 8 E |  | 558 | 80 | 74 |
| Soc. 17, T. 6 S., R. 8 E |  | 500 | 336 | 72 |
| Sec. 20, T. 6 S., R. 8 E |  | 518 | 112 | 73 |
| Sec. 20, T. 6 S., R. 8 E |  | 500 |  | 73 |
| Sec. 20, T. 6 S., R. 8 E |  | 520 | 180 | 74 |
| Sec. 30, T. 6 S., R. 8 E |  | 407 | 80 | 74 |
| Sec. 30, T. 6 S., R. 8 E |  | 437 | 90 | 74 |
| Sec. 30, T. 6 S., R. 8 E |  | 474 | 257 | 74 |
| Sec. 30, ${ }_{\text {Sec }}$ T. 6 S., R. 8 R 8 E |  | 430 |  |  |
| Sec. 30, ${ }_{\text {Sec }} 30$ T. 6 S., 6 R. 8.8 E |  | 446 | 22 | 74 |
|  |  | 410 | 125 45 | 74 |
| Sec. 14, T. 6 S.,'R. 7 E |  | 577 |  | 72 |
| Sec. 20, T. 7 S ., R. 9 E |  | 480 | 415 | 76 |
| Soc. 20, T. 7 S., R. 9 E |  | 640 | 370 | 76 |
| Sec. 20, T. 7 S., R. 9 E |  | 565 | 335 | 75 |
| Sec. 8, T. 7 S., R. 9 E . |  | 550 | 347 | 76 |
| Sec. 8. T. 7 S., R. 9 E . |  | 563 | 34 | 77 |
| Sec. 12, T. 7 S S., R. 8 EE |  | 464 | 390 | 74 |
| Sec. 2 , T. T. 7 S. ${ }^{\text {S., R. R. }} 8.8 \mathrm{E}$. |  | 496 |  | 74 |
| Sec. 2, T. 7 S., R. 8 E. |  | 499 |  | 74 |
| Sec. 2, T. 7 S., R. 8 E. |  | 477 | 291 | 74 |
| $\mathrm{Sec} .2, \mathrm{~T} .7 \mathrm{~S}$, , R. 8 E. |  | 550 |  | 74 |
| Sec. 35, T. 6 S., R. 8 E |  | 584 | 145 | 74 |
| Sec. 35, T. 6 St, T. 6 S., R. 8 R E |  | 497 | ${ }_{291}^{335}$ | 74 |
| Sec. 35, T. 6 S.,'R. 8 E |  | 504 | 620 | 74 |
| Sec. 35, T. 6 S., R. 8 E |  | 515 | 370 | 75 |
| Sec. 34, T. 7 S., R. 8 E |  | 780 |  | 90 |
| Scc. 5, T. 6 S., R. 8 E |  | 549 | 168 | 73 |

[^10]Tomperatures in wells.in. Indio Basin, Calif.-Continued.

| Location. | Depth (feet): | Flow (gallons per minute). | Temperature ( ${ }^{\circ} \mathrm{F}$.) |
| :---: | :---: | :---: | :---: |
| Sec. 5, T. 6 S., R. 8 E. | 548 | 45 | 73 |
| Sec. 6, T. 6 S., R. 8 E. | 557 | 22 | 73 |
| Sec. 6, T. $6 \mathrm{~S} ., \mathrm{R} .8 \mathrm{E}$ | 586 | 200 | 70 |
| Sec. 5, T. 6 S., R. 8 E | 550 |  | 73 |
| Sec.5,.T.6.S., R. 8 E | 553 |  | : |
| Sec. 5, T. 6 S., R. 8 E | 540 | 180 | 73. |
| Sec. 5, T. 6 S ., R. 8 EL |  |  | 73 |
| Sec. $5, \mathrm{~T} .6$ S., R. 8 E. | 557 | 340 | 73 |
| Sec. 8, T. 6 S., R. 8 E. | 545 |  | 73 |
| Sec. 8 , T. 6 S., R. 8 E. | 563 | 450 | 73 |
| Sec. 5, T. 6 S., R. 8 E. | 541 | 45 | 73 |
| Sec. 8 , T. 6 S., R. 8 E | 510 | 225 | 73 |
| Sec.12, T. 6 S.,R. 7 E. | 500. |  | 72 |
| Sec.12, T. 6 S., R. 7 E | 500 | 130 | 72 |
| Sec. 26,T. 6 S., R. 7 E | 329 | 145 | 74 |
| Sec. 26, T.6S., R. 7 E. | 320 | 112 | 74 |
| Sec. 25, T. 6 S., R.. 7 E | 385 | 100 | 74 |
| Sec. 36, T. 6 S. R. 7 E | 340 | 100 | 74 |
| Sec.1, T. 7 S., R. 7 E | 385 | 170 | 74 |
| Sec. 32, T., 6 S.,R. 8 E | 397 | 170 | 73 |
| Sec. 32, T. 6 S.,R. 8 E. | 400 | 235 | 73. |
| Sec. 32, T. 6 S., R. 8 E | 540 | 190. | 7.4 |
| Sec. 32, T.6SS.,R. 8 E | 484 | 358. | 74 |
| Sec. 32, T. 6 S.,'R. 8 E | 448 | 257 | 74 |
| Sec. 32, T. 6 S., R. 8 E | 438 | 70 | 74: |
| Sec.4, T. 7 S., R. 8 E. | 474 | 80 | 7.4 |
| Sec. 4, T. 7.S., R. 8 E. | 440 | 258 | 74 |
| Sec. 22, T. 6 S., R. 8 E | 498 | 90 | 7.4 |
| Sec. 22,T. 6 S.,.,R. 8 E. | 375 | 157 | 74 |
| Sec. 22,T. 6 S., R: 8 E | 580 | 280 | 7. |
| Sec. 26, T. 6 S., R. 8 E | 572 | 180 |  |
| Sec. 26,T. 6 S., R. 8 E | 510 | 210 | \% |
| Sec. 26, T. 6 S., R. 8 E. | 500 | 125 | 74 |
| Sec. 26,T. 6 S.,.R. 8 E. | 511 | 100. | 74 |
| Sec. 35, T. 6 S., R. 8 E. | 509 | 292 | 74 |
| Sec. 35, T. 6 S., R. 8 E | 500 | 215 | 74 |
| Sec. $22, \mathrm{~T} .6$ S. ${ }^{\text {R }}$. 8 E | 577 | 225 | 74 |
| Sec. 4, T. 6 S., R. 8 E. | 542 | 130 | \% |
| Sec. 5, T. 6 S., R. 8 E. | 574 | 130. | 74 |
| Sec. 10, T. 6 S., R. 8 E. | 526 | 90 | 75 |
| Sec. 10, T. 6 S., R. 8 E | 532 | 112 | 75 |
| Sec. 10, T. 6 S., R. 8 E | 517 | 90. | 75 |
| Sec. 34, T. 7 S., R. 8 E | 220 | 390 | 91 |
| Sec. 34, T. 7 S., R. 8 E | 250 | 390 | 91 |
| Sec. 14, T. 8 S.,.R. 8 E | 300 | 33 | 92 |
| Sec. 24, T. 8 S., R. 8 E | 315 | 390 | 94 |
| Sec. 24, T. 8 S., R. 8 E | 315 | 640 | 90 |
| Sec. 12, T. 7 S., R. 8 E.. | 464 |  | 70 |

The flows from most of these wells have remarkably low temperatures. Some are only a degree or two warmer than the mean annual air temperature of $73.9^{\circ}$ (the average for 39 years at Indio), and several are colder. One well 400 feet deep in sec. 16, T. 7 S., R. 8 E., has water $8^{\circ}$ warmer than that found in closely adjoining wells of similar depth. The 780 -foot well in sec. 34 has a flow with a temperature of $90^{\circ}$, indicating a rate of increase of $1^{\circ}$ in 48 feet, and two wells near by 220 and 250 feet deep have flows with a temperature of $91^{\circ}$, indicating rates of increase of $1^{\circ}$ in 13 and 15 feet, respectively. About 3 and 4 miles to the south, along a prolongation of the same line, are wells yielding large flows with temperatures of $92^{\circ}$, $94^{\circ}$, and $90^{\circ}$. It is probable that these wells tap a supply from a fissure, possibly a fault, along which heated waters are rising from a deep-seated source. It is. stated that a 500 -foot hole was sunk in
the bottom of the Salton Basin before the inundation and obtained a flow of water with a temperature of $92^{\circ}$.

OIL REGION.
In investigating the chemistry of the waters of the oil fields in San Joaquin Valley Rogers ${ }^{1}$ determined temperatures of flow from several deep wells in Fresno and Kern counties.

Temperaiures in several deep wells in Fresno and Kern counties, Calif.

| Location. | Depth of source of flow (feet). | $\begin{gathered} \text { Temper- } \\ \text { ature } \\ \left({ }^{\circ} \mathrm{F} .\right) . \end{gathered}$ | $\begin{array}{\|c\|} \text { Mean } \\ \text { annual air } \\ \text { tempera- } \\ \text { ture } \\ \left({ }^{\circ} \mathrm{F} .\right) . \end{array}$ | Depth per degree of increase in temperature (feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Fresno County: <br> Sec. 12 「 21 S R 14 E |  | 118 | $a 61.5$ |  |  |
|  | 2,077 | 118 | $a 01.5$ | 37 | Sulphur water 600 feet below oil sand. |
| $\underset{\text { Sec. 2, T. } 20 \text { S., R. } 14 \text { E.. }}{\text { Kern County }}$ | 1,104 | 86 | 01.5 | 45 |  |
| Sec. 9, T. 31 S., R. 22 E.. | 3,860 | 131 | 665 | 58.1 | Flow probably from 3,000+ |
| Sec. 32, T. 12 N., R. 23 E. | 2,505 | 1.15 | 65 | 50 | feet below main oil sand. Flow from below oil sand about 100 feet above second oil sand. |
| Sec. 35, T. 31 S., R. 22 E.. | 1, 495-1,727 | 120 | 65 | $2 \overline{7}+$ | About 100 feet below lowest oil sand. |
| Sec. 15, T. 32 S., R. 23 E.. | 1,765-1,820 | 102 | 65 | 48 | Flow from about 150 feet below oil sand. |
| Sec. 25, T. 32S., R. 23 E.. | 1,460 | 97 | 65 | 46 | Flow from about 50 feet beloy oil sand. |
| Sec. 31, T. 32 S., R. 24 E.- | 1, 334-1, 609 | 109 | 65 | $30+$ | Do. |
| Sec. 26, T. 12 N., R. 24 W. | 2,540-2,560 | 93 | 65 | 91 | Flow from about 125 feet below oil sand. |
| Sec. 12, T. 11 N., R. 24 W. | 3,550 | 104 | 65 | 01 | Flow from about 2,550 feet below oil sand. |
| Sec. 35, T. $32 \mathrm{~S} ., \mathrm{R} .23 \mathrm{E} .$. | 1,090 | 84 | 65 | 58 | Shale of Monterey group. |
| Sec. 23, T. 31 S., R. 23 E.. | 3,000 | 1.25 | 65 | 50 | Flow from a fow feet below oil sand. |

${ }^{a}$ Hanford average minus $7^{\circ}$ for difference in altitudo.
$b$ Bakersfield average minus $1^{\circ}$.
It is reported that in a 5,390 -foot hole at Rosemary station the oil at 3,000 feet had a temperature of $140^{\circ}$ to $160^{\circ}$.

## GRASS VALLEY.

Lindgren ${ }^{2}$ made observations of temperatures in the mines of gold-bearing quartz in the Grass Valley and Nevada City districts in 1894. Thermometers were left in holes with the following results: 30 feet below surface, $53 \frac{1}{2}^{\circ} ; 1,513$ feet below surface, $66^{\circ}$; 1,513 feet below surface (dry hole in quartz), $67 \frac{1}{4}^{\circ} ; 1,553$ feet below surface, $66^{\circ}$. These observations indicate a rate of increase of $1^{\circ}$ in 122 feet at 1,553 feet and of $1^{\circ}$ in 105 feet at 1,513 feet.

[^11]
## COLORADO.

## ARKANSAS VALLEY.

There are flowing wells at short intervals along the valley of Arkansas River from Canon City to Granada, Colo. The temperatures in some of them are as follows:

Temperatures in wells in Arkansas Valley, Colo.

| Location. | Depth (feet). | Flow (gallons per min- ute). | Tem-perature ( ${ }^{\circ} \mathrm{F}$.). | Mcan annual air tem-perature ( ${ }^{\circ} \mathrm{F}$.). | Depth per degrec of increase in tem-perature (feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Canon City: |  |  |  |  |  |  |
| - Oil company, sec. 23a | 1,600 | 700 | 90 | b 52.9 | 31 | Flcw probably from 1,150 fect. |
| Sanitarium 3 miles northeast, sec. 26. | 1,670 | Many. | $98 \frac{1}{2}$ | 52.9 | 36 | Water from 1,650 feet, test by owner. |
| Florence: |  |  |  |  |  |  |
| Sec. 14. | 800 | 1,000 | 86 | 52.9 | 35 | Test by N.H. Darton. |
| 10 miles southeast of. | 1,085 | 750 | 92 | 52.9 | 28 | Test January, 1905, by Dick Rule with Darton thermometer. |
| SW. $\frac{1}{4}$ sec. 26, T. 18 S., R. 69 W.c | 1,230 | 360 | 87 | 52.9 | 36 |  |
| Pueblo: |  |  |  |  |  |  |
| Grand Hotel. | 1,219 | 20 | 76 | b 51.5 | 4912 | Pressure 50 pounds. Test by N. H.Darton Nov. 7, 1898. |
| Sec. 12, T. 21 S., R. 65 W. | 1,260 | 25 | $80 \frac{1}{2}$ | 51.5 | 43 | Pressure 15 pounds. |
| Ferris Hotel. | 1,400 | $\frac{1}{2}$ | 77.2 | 51.5 | $54 \frac{1}{2}$ | Pressure 61 pounds. Test by N. H. Darton Nov. 7, 1898. |
| Clarks. | 1,402 | 87 | 79.5 | 51.5 | 50 | Flow through 8 feet of rubber hose. Test by N. H. Darton Nov. 7, 1898. |
| South of......--....... | 1,404 | 100 | 82 | 51.5 | 39 | Flow at 1,180 feet. |
| Baxter, $4 \frac{1}{2}$ miles north of..... | 660 |  | 73.2 | 51.5 | $30 \frac{1}{2}$ | Test by C. A. Fisher February, 1904. Down 3 hours, water stands at 550 feet. Some gas. |
| Boone, $2 \frac{1}{2}$ miles north of. | 1, 000 | 49 | 82.5 | 51.5 | 32 | Test by C. A. Fisher, February, 1904; down 3 hours. Dry hole. |
| Manzanola | 1,113 | 42 | 80 | d 51.5 | 39 | Large flow. |
| Rocky Ford. | 707 | 130 | 78 | 51.5 | 29 | Do. |
| Rocky Ford, sec. 18 | 790 | 68 | 74 | 51.5 | 35 | Do. |
| Rocky Ford | 793 | 68 | 78 | 51.5 | 30 | Do. |
| Do.................... | 820 | 38 | 75 | 51.5 | 35 | Do. |
| Rocky Ford, W ychoff Park.. | 845 | 115 | 72 | 51.5 | 41 | Do. |
| Do.. | 1, 003 | 17 | 75 | 51.5 | 43 | Do. |
| La Junta. | 412 | 8 | 68 | - 52 | 26 | Do. |
| La Junta mill. | 420 | 45 | 70 | 52 | 23 | Do. |
| La Junta, railroad | 420 |  | 68 | 52 | 26 | Air pumped. |
| Do.... | 439 |  | 68 | 52 | $.27 \frac{1}{2}$ | Do. |
| Do... | 700 |  | 68 | 52 | 44 | Water probably from higher level, at least in part. |
| La Junta. | 640 | 35 | 68 | 52 | 40 |  |
| La Junta Home Co. | 766 | 25 | 68 | 52 | 48 | Do. |
| La Junta, near. | 740 | 12 | 75 | 52 | 32 | - Do. |

[^12]
## SAN LUIS VALLEY.

Siebenthal ${ }^{1}$ has supplied many observations of well temperatures in the San Luis Valley. The following data are compiled from his report:

Temperatures in wells in or near Alamosa, Colo.

| Location. | Depth (feet) | $\begin{array}{\|c} \text { Depth } \\ \text { of } \\ \text { flow } \\ \text { (feet). } \end{array}$ | $\begin{array}{\|c\|} \text { Flow } \\ \text { (gallons } \\ \text { per } \\ \text { min- } \\ \text { ute). } \end{array}$ | Tem-perature <br> $\left({ }^{\circ} \mathrm{F}.\right)$.$\qquad$ | Depth per degree of increase in temper- ature. (feet).. | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Northeast corner sec. 10. | 1,000 | 932 | 600 | 74.7 | 20 | Smaller flow at 500 fect may mingle. |
| Northwest corner sec. 10... | 865 |  | 400 | 72 | 30 | Cased to 852 feet. |
| Goodall well................ | 883 | 876 | Many. | 70 | 32 | Small flow at 261, 507, and 617 feet; larger ones at 776 and 883 feet. |
| North of Alamosa.. | 725 |  | Few. | 66 |  | Probably a higher flow mingles. |
|  | 680 | 680 | Many. | 69 | 26 | Cased to bottom. |
| Blectric-light plant. | $800+$ | $800+$ | Many. | 73 | $40 \pm$ | Do. |
| $1 \frac{1}{2}$ miles south of Alamosa.. | 810 | 810 |  | 72 | 28 | Do. |
| 1 mile southeast of Alamosa.. | 800 | 800 | Many. | 72 | 28 | Do. |
| 4 miles north of Alamosa. | 840 | 840 |  | 72 | 29 | Do. |
| Near Tountain. | 58.5 | ...... | Many. | 64 | 28 |  |
| La Jara Do. | 325 |  | Many. | ${ }_{49}{ }_{4}$ | 62 50 |  |
| Los Sauces, near | 200-300 |  | Many | 53-55 ${ }^{\text {2 }}$ | 20-25 | Casca to 300 feet. |
| Lockett, near. | 805 | 800 | 150 | 63 | 40 |  |
| Veteran, near | 265 | 265 | 250 | 54 | 24 | Cased to bottom. |
| Do...... | 383 | 335 | Many. | 55 | 28 |  |
| Swede Corners. | 156 | 147 | Many. | 61 | 8 |  |
| Swede Corners, 1.12 miles | 189 | 189 |  | ${ }_{58}$ | 12 |  |
| northeast of. Swede Corners, $1 \frac{1}{2}$ miles northeast of. | 126 |  |  | 45 | 83 |  |
| Moffat, near.. | 640 | ${ }_{6}^{616}$ |  | 61 | 34 |  |
| Do. | 300 | 300 | 8 | 51 | 37 |  |
| San Isabel, near | 380 |  | 55 | $55 \frac{1}{2}$ | 30 |  |
| Do. | 375 |  | 58 | 58 | 25 |  |
| Baca grant. | 481 | $\ldots$ | 10 | 62 | 25 |  |
| Mosca. | 500 |  | 50 | 66 | 22 |  |
| Do. | 600 |  |  | 69 | 23 |  |
|  | ( $\begin{array}{r}385 \\ 500\end{array}$ |  |  | 60 63 | $\stackrel{23}{25}$ | Second flow. Scparate casing. |
| Do. |  | $\begin{aligned} & 500 \\ & 340 \end{aligned}$ |  | 63 | 25 18 | Scparate casing. Do. |
| Do. | 800 |  | 90. | 71 | 29 |  |
| Mosca, 5 miles west of | 712 |  | 120 | ${ }_{69}^{63}$ | 36 |  |
| Hooper. | 740 300 | $\begin{aligned} & 740 \\ & 300 \end{aligned}$ | 70 1 | 69 53 | ${ }_{30}^{28}$ | Cased to bottom. Do. |
| Hooper, ne | 425 |  | Many. | 54 | 39 |  |
| Do. | 550 |  |  | 59 | 34 | Flow from above 380 feet. |
| Do.. | ${ }_{7}^{630}$ |  | 70 | 60 | 36 | Cased to 360 feet. |
| Kinney ranch Jacobs ranch. | 766 |  |  | 68 | ${ }_{33}^{51}$ | Considerable gas. |
| Jacobs ranch Do....... | 400 |  | 25 | 58 57 | 33 29 | Cased to 125 feet. Cased to 224 feet. |
| San Luis. | 22-32 |  |  | 45-47 |  |  |
| Del Norte. | 450 | 450 | 2 | 54 | 41 |  |

a Temperature about Alamosa observed by Prof. Carpenter.
$b$ Based on an approximate mean annual of $43^{\circ}$ averaged from observations in adjoining regions.
It is suggested that the high rate about Swedes Corners is probably due to the presence of an old lava flow, a part of which crops out 1 mile west and 3 miles north from the Corners.

[^13]The flow from a 405 -foot well on Cotton Creek, in Saguache County, is reported to have a temperature of $58^{\circ}$, but as the mean annual air temperature at that locality is not known the rate of increase can not be calculated.

## LEADVILLE.

It is reported ${ }^{1}$ that a 400 -foot well in sec. 25 near Leadville yields a flow with a temperature of $58^{\circ}$. As the mean annual air temperature at Leadville is $35^{\circ}$ (average for 18 years) this would indicate a rate of increase of about $1^{\circ}$ in 17 feet.

TRINIDAD.
In June, 1905, W. T. Lee made temperature determinations at various depths in a 2,500 -foot boring at Trinidad. The hole was found to be plugged at 770 feet, and a Darton thermometer left at that depth for several hours recorded $69.6^{\circ}$. A test at 370 feet gave $67.8^{\circ}$. As the mean annual temperature of Trinidad is $51.3^{\circ}$ (average for 15 years) these readings indicate rates of increase of $1^{\circ}$ in 42 feet and 22 feet, respectively. The difference between the two readings, however, indicates a rate of increase of $1^{\circ}$ in 222 feet between depths of 370 and 770 feet.

## DENVER BASIN.

Few data are available as to the temperatures in many wells in the vicinity of Denver that formerly flowed. The following figures ${ }^{2}$ may possibly be significant:

Temperatures in wells in and near Denver, Colo.

| Depth (feet). | Flow (gallons per minute). | Tem-perature ( ${ }^{\circ} \mathrm{F}$.). | Depth <br> degree of increase of temperature (feet). $a$ | Remarks. |
| :---: | :---: | :---: | :---: | :---: |
| 350 |  | 56 | 56 |  |
| 340 | 3 | 58 | 41 | Some flow from 240 feet perhaps. |
| . 407 | 30 | 62 | 33 |  |
| 239 | Many. | 55 | 46 |  |
| 445 | 50 | 62 |  |  |
| 410 |  | 58 | 50 |  |
| 563 | 20 | 66 | 35 |  |
| 635 352 | 10 | $\bigcirc 62$ | 52 | Littleton. |
| 352 | 140 | ${ }^{\text {b }} 52$ | 160 | Cased to 248 feet. |
| 620 | 325 | ${ }^{6} 65$ | 38 | Principal flows from 250 and 575 feet. Both same temperature. |
| 480 | 300 | ${ }^{6} 57$ | 67 | Cased to 350 fect. |

$a$ Calculated from mean annual temperature of $49.8^{\circ}$, average for 45 years.
b Emmons, S. F., Cross, Whitman, and Eldridge, 'G. H., Geology ol the Denver Basin in Colorado: U. S. Geol. Survey Mon. 27, pp. 456-459, 1896.

MISCELLANEOUS LOCALITIES.
Temperatures of wells in Montrose, Akron, and Loveland, Colo.

a 30 -gallon flow.
$b$ A verage for 21 years.
c Approximated from Fort Morgan, $48.7^{\circ}$ (20-ycar average).

- Fort Collins average for 33 years.
$e$ Calculated to $1,365^{\text {f fect. }}$


## SUMMARY.

The rate of increase of temperature in wells in the Arkansas Valley varies considerably, but in most of them it is high. From Baxter to La Junta the average is near $1^{\circ}$ in 34 feet and variations from the average are mostly less than 10 feet. The Pueblo wells show an average near $1^{\circ}$ in 50 feet, but farther west the rate is higher. In San Luis Valley the rate is $1^{\circ}$ in 33 to 42 feet so far as the few observations show. The Trinidad hole showed too much variation to be significant. The rate in the Denver wells was $1^{\circ}$ in 33 to 56 feet, at Montrose $1^{\circ}$ in $33 \frac{1}{2}$ feet, and at Akron $1^{\circ}$ in 45 feet. The observation at the Loveland well is open to question as to the source of the flow.

## GEOLOGIC RELATIONS.

Wells in the Arkansas Valley below the Pueblo draw from the Dakota and associated sandstones, which dip gently eastward. The well in T. 18 S., R. 69 W., also reaches the Dakota sandstone. The Florence wells penetrate much higher sandstones in a basin holding coal measures, and the Canon City holes are on the west rise of this basin. The San Luis Valley wells are in Quaternary valley filling. The Trinidad hole was in the Pierre shale at 770 feet. Wells in the Denver Basin penetrate the nearly horizontal younger Tertiary deposits. At Montrose the Dakota sandstone was probably reached, and the Akron hole was in the Pierre shale. The precise relations at the Leadville well are not known, but the region contains limestone cut by igneous rocks and greatly mineralized.

## FLORIDA.

Although many determinations of temperature of flows from artesian wells in Florida have been published, there are few reliable data to indicate the geothermal gradient. This is due largely to lack of information as to the depth from which the water is derived and as to whether or not higher flows mingle with the main flow.

ST. AUGUSTINE.
In the large well at the Ponce de Leon Hotel, St. Augustine, the main flow, from 1,340 to 1,390 feet, has a temperature of $86^{\circ} .{ }^{1}$ As the mean annual temperature of St . Augustine is $69.5^{\circ}$ (65-year average), this would indicate a rate of increase of $1^{\circ}$ in 81.2 feet, providing no water enters the casing above 1,340 feet. When the well was being bored the temperatures of water from the principal large flows brought rapidly to the surface in the sand pump from various depths were recorded as follows: ${ }^{2}$

Temperature at different depths in well at Ponce de Leon Hote l, St. Augustine, Fla.

| Depth (feet). | Tem- <br> pera- <br> ture <br> (oF.). | Depth per <br> degree of <br> increase in <br> tempera- <br> ture (feet). |
| ---: | ---: | ---: |
| 170 | 74 | 38 |
| 410 | 76 | 63 |
| 520 | 79 | 55 |
| 1,110 | 80 | 106 |
| 1,225 | 85 | 79 |
| $1,340-1,390$ | 86 | 81 |

The temperature at a depth of 35 feet was reported as $62^{\circ}$, or $7 \frac{1 \frac{1}{2}^{\circ}}{}$ less than the mean annual air temperature. The rate of increase from 170 to 1,340 feet is $1^{\circ}$ in $87 \frac{1}{2}$ feet.

[^14]
## MISCELLANEOUS. DATA.

Data for Florida, some of which may be reliable, are as follows:
Temperatures in wells in Florida.

| Location. | Depth (feet) | $\begin{gathered} \text { Flow } \\ \text { (gals.). } \end{gathered}$ | Tem- pera- ture (are $)$ | $\begin{gathered} \text { Mean } \\ \text { annual } \\ \text { air } \\ \text { tem- } \\ \text { pera- } \\ \text { ture } \\ \left({ }^{\circ} \mathrm{F} .\right) . \end{gathered}$ | Depth per degree of increase perature (fect). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |
| Malabar. | 350 |  | 78 | 72.6 | 65 | Large flow; first flow at 280 |
| Melbourne, 1 mile west of. | 400 ? |  | 77 | 72.6 | 90 | Large flow; possibly some |
| Micco. | 500 ? |  | 78 | 72.6 | 93 | Large flow; may be from 250 |
| Rockledge. | 350 |  | 80 | b 72 | 44 | Large flow; cased to bottom. |
|  | 218 |  | 75 | c 71.2 | 60 | Large flow. |
|  |  |  |  |  |  |  |
| Magnolia Springs.. | 362 |  | 74 | 68.3 | 64 | Large flow. ${ }^{\text {a }}$ |
|  |  |  |  |  |  | Do. |
|  |  |  |  |  |  | "Flow from 550 feet." |
| Jacksouville Electric Co.. | 1,020 |  | 81 | 68.2 |  | Main flow probably from 555 |
| Jacksonville city well. | 984 |  | 772 | 68.2 |  | Water probably also from 524; 750 , and 860 feet. |
| Do. | 498 |  | 71 | 68.2 |  |  |
| Do. | 635 |  | 74 | 68.2 |  | Mixed flows. |
| Do. |  |  |  | 68.2 |  |  |
| Do. | 900 |  | 74 | 68.2 |  | Do. |
| Do.. | 1,031 |  | 74 | 58.2 |  | Do. |
| Jacksonville, 3 mileseast of <br> Lee County: 800 $\ldots \cdots$ 75 68.2 118 Flow from bottom. |  |  |  |  |  |  |
|  |  |  |  |  |  |  |
| Fort Myers 1 mile east of. | 403 |  | 80 | 73.1 | 56 | 150 -gallon flow. |
| Manatec County: 17 . |  | Many. | 72 | 071.7 |  | Flow from 347 feet. |
| Ellentown. | 400 | Many. | 80 | $\bigcirc 71.7$ | 48 | Flow mostly from 200 feet. |
| Manatee. | 368 | 150 | 74 | g 71.7 | 150 | Flow from 348 feet. |
| Terra Ccia. | 378 | Many. | 75 | ${ }^{\text {h }} 71.9$ | 116 | Flow from 360 feet. |
| Do. | 350 | 250 | 76 | h 71.9 | 73 | Flow from 300 feet. |
| Marion County: Ocala. | 1,210 | Many. | 742 | i 70.2 | 284 | Pumped 74 feet. |
| Nassan County: Fernandina. | 700 |  | 74 | ${ }^{\text {e } 68.2}$ | 121 |  |
|  |  |  |  |  |  |  |
| Putnam County: | 300 |  | 72 |  | 125 | Water from 300 feet. |
| Federal Point. | 160-250 | 200-600 | 72-75 | $i 69.6$ |  | Water from 155 to 200 fect. |
| 10.. | 250 | 500 | 72 | '69.6 | 83 | Water from 200 feet. |
| Do. | 225 | 600 | 75 | $i 69.6$ | 31 | Water from 168 feet. |
| Palatka | 206 | Many. | 76 | ${ }^{*} 69.9$ | 31 | Water from 190 teet. |
| Do.. | 250 | Many. | 72 | ${ }^{k} 69.9$ | 119 | Water from 250 feet. |
|  |  |  |  |  |  |  |
| Orchid...... | 368 | Many. | 78 | ${ }_{m} 72.6$ | 66 |  |
| Orchid, $\frac{1}{2}$ mile east of | 400 | Many. | 78 | m 72.6 | 74 |  |
| Sebastian. | 460 | Many. | 75 | $n 72.7$ | 200 |  |
| Do.............. <br> St. John County: |  |  |  |  |  |  |
|  |  |  |  |  |  |  |
| Hastings... <br> Switzerland | 155 337 | 120 400 | 75 | O 69.6 $p 68.9$ | 29 107 | Water from 155 feet. |
|  | 300 | 400 | 72 | ${ }_{p} \mathbf{p} 8.9$ | 90 | Water from 280 feet. |
|  |  |  |  |  |  |  |
| Many wells. | 100-163 | Many. | 70-73 | 970 |  |  |
| Ormond.... Lake Helen | 227 | $\underset{115}{ }$ | 70 | 70 70 | 28 |  |
|  |  |  |  |  |  |  |

[^15]Average for 18 years.
$k$ Federal Point average plus $0.3^{\circ}$.
${ }^{2}$ Federal Point average plus $0.4^{\circ}$.
$m$ Fort Pierce average minus $0.3^{\circ}$.
$n$ Fort Pierce average minus $0.2^{\circ}$.

- St. Augustine average plus $0.1^{\circ}$.
${ }_{p}^{\circ}$ St. Augustine average plus $0.1^{\circ}{ }^{\circ}$.
$q$ Deland average for 20 years.

Apparently the most important data on temperature here presented are those from St. Augustine, Fort Myer, Mandarin, and Rockledge. Some of the others may be reliable, but there is a possibility that in many of the wells higher flows mingle with the main flow. The source of the flows at Melbourne and Micco is not known, and most of the other wells in Brevard County are shallow. The flows at Jacksonville are known to be mixed, for although the main flow is derived from a depth of 555 feet other flows come in, notably in the deeper wells, which tap a flow at a depth of 975 feet, below a body of clay. Temperatures of flows in the 984 -foot city well, which is cased only to 494 feet, are stated ${ }^{1}$ to be as follows: 524 to 727 feet, $76^{\circ}$ to $77 \frac{1}{2}^{\circ}$; 865 to 970 feet, $78^{\circ}$ to $79^{\circ}$; flow at mouth, $77 \frac{1}{2}^{\circ}$. The large flow from the 496 -foot well at Fort Myer probably affords a true indication of underground temperature, although it shows a smaller rate of increase than is indicated by the well a mile east of Fort Myer.

## GEOLOGIC RELATIONS.

The wells in Florida penetrate sands, clays, and limestones of Tertiary age, which lie nearly horizontal and are uniform in character over wide areas. Those at St. Augustine, Jacksonville, and Mandarin penetrate deeply into Oligocene limestones of the Vicksburg group, which are overlain by clays and sands of later Tertiary age. These limestones were also reached by the holes in Brevard and Clay counties and at Fort Myer. The variations in geothermal gradient have no obvious connection with geologic features.

## GEORGIA.

## TEMPERATURES.

Local observers have furnished records of the temperature of flows from several wells in Georgia which afford a few data as to the geothermal gradient. Other facts regarding the wells are given by McCallie ${ }^{2}$ and by Stephenson and Veatch. ${ }^{3}$

[^16]Temperatures in wells in Georgia.

| Locality. | Depth (feet). | Flow (gal- lons per min- ute). | Tem- pera- ture $\left({ }^{\circ} \mathrm{F}.\right)$ | Mean air tem-perature ( ${ }^{\circ} \mathrm{F}$.). | Depth per degree of increase in tem-perature (feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Chatham County: Savannah, 12 miles west of. | 659 | Many. | 72 | a 66.6 | 122 |  |
| Dougherty County: b Albany city well. | 750 | 200 | 732 | c 66.9 | 103 | Cased to 600 feet, principal flow from 680 fect. |
| Albany, well No. $2 d$ | 1,320 | 125 | 78 | c 66.9 | 119 | Cased to bottom. |
| Evans County: Claxtone. | 5461 | 100 | 75 | $f 65.8$ | 59 | Cased to 459 feet, water begins at 460 fect. Water pumped to 80 feet. |
| Glynn County: Brunswick... | 465 | 600 | 70 | $g 67$ | 155 | Possibly some flow from 300 feet also. |
| McIntosh County: Valona.... | 455 | 120 | 70 | $h 66.1$ | 116 | Possibly some flow from 320 to 365 feet also. |
| Pierce County: Offerman e... | 6751 | 250 | 76 | 167.2 | 73 | Cased to 635 feet. Water 640 to $675 \frac{1}{2}$ feet. Rises to within 33 feet of surface. |
| PulaskiCounty: Fiavkinsville | 490 | Many. | 71 | 362.4 | 57 | Cased to bottom. Flows. |
| Ware County: Way Cross, 2 wells. | 691 | Many. | 722 | $k 67.2$ | 127 | Air-pumped, main supply 670 to 691 feet. |
| Wilcox County: Bowens mill. | 673 | Many. | 74 | $l 66.2$ | 86 | Source of water not, given. |
| a Average for 67 years. <br> b Georgia Geol. Survey Bull. 7, pp. 178-181, 1902. <br> c Average for 30 years. <br> d Georgia Geol. Survey Bull. 15, pp. 97-99, 1908. <br> e U. S. Geol. Survey Bull. 298, pp. 201-202, 1900. <br> f. Statesboro average for 17 years. |  |  | $g$ Average of various distant stations. <br> $h_{\text {Savannah average }}+0.5^{\circ}$. <br> i Waycross average for 28 years. <br> j A verage for 11 years. <br> * Average for 28 years. <br> $l$ Eastman average for 25 years. |  |  |  |
|  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |

## SUMMARY.

The data given indicate considerable variation in the geothermal gradient in Georgia, but it is difficult to know which are reliable. The figures for the Claxton, Hawkinsville, and Offerman wells appear to be valid, and the source of flow in the wells at Albany and a well 12 miles west of Savannah, which show a low rate of increase, may be correctly reported. The Waycross water is probably cooled by air lift. The Georgia wells can not be grouped geographically by their rates of increase or by their depths.

## GEOLOGIC RELATIONS. ${ }^{1}$

The wells in the list are all in the Coastal Plain province, in which sheets of sand, gravel, and clay of Cretaceous and Tertiary age dip gently to the east and southeast. The older beds lie on a basement of pre-Cambrian crystalline rocks, which slopes to the southeast and is far beneath the surface along the coast and at the Florida State line. The Albany deeper wells draw from the Ripley formation, high in the Cretaceous. The water-bearing stratum at Hawkinsville and Bowens Mill is sand, probably at the top of the Cretaceous or the base of the Eocene. The Offerman, Claxton, and Waycross wells draw from beds in the Eocene at a considerably higher horizon, and
the Savannah, Brunswick, and Valona wells may reach beds still higher. From these statements it will be seen that there is no apparent relation between the rates of increase in temperature and the water horizons.

> IDAHO.
> BOİSE.

The water of the artesian wells at and near Boise is reported to have temperatures of $80^{\circ}$ to $170^{\circ}$. As the depths of these wells are less than 500 feet, the higher temperatures, at least, indicate a very high geothermal gradient, which is due, doubtless, to local volcanic conditions.

## SNAKE RIVER VALLEY.

Russell ${ }^{1}$ reports temperatures in a number of wells along the Oregon Short Line Railroad on the Snake River Plains as observed by Scott Turner. They all indicate high temperatures. The most notable are as follows:

Temperature in wells along Oregon Short Line Railroad in southwestern Idaho.

| Location. | Depth (leet). | Depth of flow (feet). | Temperature ( $\left.{ }^{\circ} \mathrm{F}.\right)$. | Mean annual air temperature ( $\left.{ }^{\circ} \mathrm{F}.\right)$. | Depth per degree of increase in temperature (feet). |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Bliss . | 483 | 430 | 70 | a50.3 | 22 |
| Cleft. | 450 |  | 73 | $b 48.6$ | 19 |
| Owyhee | 600 | 530 | 70 | 52 | 30 |
| Nampa. | 114 | 40 | $61 \frac{1}{2}$ | 50 |  |

$a$ Approximated from averages at Boise and Garnet. b Mountain Home average.
Russell ${ }^{2}$ has also given data of wells on the south side of Snake River valley in Owyhee County. The temperatures are high, and as there are numerous hot springs in the region it is believed that the wells draw in part from deep-seated thermal waters rising in fissures. The wells all flow from 20 to 120 gallons a minute and are sunk to sands in the Payette formation, of lower Tertiary age.

[^17]Temperature in wells in Owyhee County, Idaho.

| Locality. | Depth (feet). | Temperature ( ${ }^{\circ} \mathrm{F}$.). | Depth per degree of increase in temperaturo (feet). |
| :---: | :---: | :---: | :---: |
| Guffey. | 538 | $76 \frac{1}{2}$ | 22 |
| Central | 1,033 | $10{ }^{2}$ | $21 \frac{1}{2}$ |
| Do. | , 940 | 98 | 20.2 |
| Do. | 1,035 | 106 | 19 |
| Do. | 720 | 100 | 15 |
| Enterprisc. | 340 | 87 | 91 |
| Do:.... | 385 | 90 | 10 |
| Hot Spring, 5 miles north of | 700 | 90 | $18 \frac{1}{2}$ |
| De Lamar ${ }^{\text {a }}$. | 975 | 120 | 14 |

a Lindgren, Waldemar, and Drake, N. F., U. S. Geol. Survey Geol. Atlas, Silver City folio (No. 104), p. 8, 1904.

These rates are based on an assumption that the mean annual temperature is $52^{\circ} \mathrm{F}$., a figure approximated from observations at distant stations.
Russell also states that the water in a 240 -foot well near Hot Springs, in Bruneau Valley, Owyhee County, had a temperature of $109^{\circ} \mathrm{F}$. The well flows 7 gallons a minute and is near a hot spring of the same temperature. There are many shallow wells on these lava plains of Snake River which yield warm water, but as the region is one of relatively recent volcanic activity the lava flows are doubtless the source of the heat.

## NEZ PERCE COUNTY.

In his report on Nez Perce County, Russell ${ }^{1}$ gives a number of temperatures of deep wells and warm springs. A 220 -foot well on the Dowd farm, 8 miles southeast of Lewiston, has a flow of 15 to 20 gallons a minute, which in December, 1900, John Adams found to have a temperature of $583^{3}$. A copious spring near by had a temperature of $54^{\circ}$. Several wells about 100 feet deep, not far away, had flows with temperatures of $66^{\circ}$ to $68^{\circ}$, according to the same observer. The mean annual temperature of the locality is estimated by Russell as $48^{\circ}$ to $49^{\circ}$. Russell believed that the water must come from a depth of 540 to 1,000 feet in these wells in order to account for the high rate of increase indicated, but as the region is underlain by recent basalt it appears to me more likely that this rock at a moderate depth is the source of the heat.

## ILLINOIS.

## TEMPERATURES.

About fifty records of temperatures in deep wells in various parts of Illinois were obtained, and although in many of the wells the flows are mixed, in others the depth of flow is known and the rate of increase is indicated. At Streator and St. John Darton thermometers were sunk to the bottom of the wells.

[^18]Temperatures in wells in Illinois.

| Location. | Depth (feet). | Tem-perature ( ${ }^{\circ} \mathrm{F}$.). | $\begin{gathered} \text { Mean } \\ \text { annual } \\ \text { tempera- } \\ \text { ture } \\ \left({ }^{\circ} \mathrm{F} .\right) . \end{gathered}$ | Depth per degree of increase in temperature (feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Alexander County: Cairo: a Halliday Hotel | 824 | 62 | 57.5 | 155 | Water possibly all or mainly from |
|  | 824 | 62 | 57.5 | 155 | ater possibly all or mainity from 498 to 518 fect. • |
| Halliday residence. | 811 | 62 | 57.5 | 155 | 60 -gallon flow from 753 feet; no increase below 700 feet. |
| Near mouth of Cache River | 806 | 62 | 57.5 | 163 | Water from 735 feet or higher. |
| Cook County: |  |  |  |  |  |
| Evanston................... | 1,602 | 65 | $b 48.5$ | 97 | Source of flow not given. Probably higher waters mingle. |
| Oak Park | 2,180 | 64 | 48.5 |  |  |
| Chicago c. | 751 | 55 | 48.5 | 115 |  |
| Hancock County: Warsaw... | 844 | 60 | d 51.4 | 98 | Source of flow not given. |
| Henry County: Kewanee, 3 wells. | 1,500 | 65 | e 49.4 | 90 | Cased to 1,330 feet. |
| Knox County: <br> Galesburg $f$. | 1,226 | 60 | $g 50$ | h 106 | Cased to St. Peter sandstone at 1,060 feet. Pumped 160 feet; main supply from 1,100 to 1,215 feet. |
| Knoxville. | 1,360 | 68 | 50 | $i 66$ | Cased to St. Peter at 1,180 feet. |
| La Salle County: Streator...... | 2,496 | j 76 | ${ }^{k} 50.4$ | 97 | Thermometer to bottom. |
| Mercer County: Aledo......... | 3,115 | 68 | $l 50$ |  | Cased 1,705 to 1,805 feet. "Most of flow from bottom." |
| McDonough County: Macomb. | 1,630 | 68 | $m 51$ | $n 67$ | Cased to St. Peter sandstone at 1,135 feet. |
| Perry County: St. John Coal \& Salt Co. | 3,735 | - 101 | 56.3 | 83.5 | Thermometer down 24 hours. |
| Putnam County: Hennepin... | 800 | 58 | $p 49.4$ | 93 | Cased nearly to bottom. |
| Peoria County: $q$ |  |  |  |  |  |
| Puoria Asylum | 1,600 900 | 78 | l 49.9 | 56.9 59.6 | source of flow not known. <br> Do. |
| Stockyards................. | 850 | 65 | . | 56.3 | Do. |
| Sulphur water house....... | 800 | 62 . |  | 66.1 | Do. |
| Rock Island County: $n$ |  |  |  |  |  |
| Carbon Clift | 950 | 60 | 849.4 | 90 | Do. |
| East Moline | 1,050 | 61.5 | 49.4 | 87 | No casing. |
| Milan $t$. | 1,027 | 61.5 | 49.4 | 85 |  |
| Moline: <br> Prospect Park | 1,121 | 61 | 49.4 | 99 | Flows may mingle |
| Paper mills. | 1,375 | 63 | 49.4 | 101 | Do. |
| Rock Island: |  |  |  |  |  |
| Hubers Brewery....... | 1,187 | 60.5 | 49.4 | u 82 | Cased to 912 feet. |
| Mitchell \& Lynie...... | 2,049 | 68 | 49.4 | 83 | Cased to 1,200 feet. "Probably main flow from 1,550 feet." |
| St. Clair County: Marissa...... | 685 | 65 | $v 55.2$ | 70 | Pumps 30 gallons 200 feet. |
| Stark County: Bradford....... | 2,054 | 68 | w 49.4 | 110 | Cased to 1,600 feet. Water pumped 175 feet. |
| Tazewell County: Pekin....... | 950 | 70 | $x 49.9$ | 42 | Sulphur water. . |
| Whiteside County: Sterling.... | 1,450 | 62 | $y 48.2$ | 105 |  |

${ }^{a}$ Glenn, L. C., Underground water of Tennessee and Kentucky west of Tennessee River and of an adjacent area in Illinois: U. S. Geol. Survey Water-Supply Paper 164, pp. 150-152, 1906.
$b$ Chicago average for 42 years.
c Schott, C. A., On underground temperatures: Smithsonian Rept. for 1874, p. 250.
${ }^{d}$ Keokuk, Iowa, average for 46 years.
e Gaiva average for 24 years.
$f$ Udden, J. A., Some deep borings in Illinois: Illinois Geol. Survey Bull. 24, p. 65, 1914.
$\theta$ Approximated.
${ }^{h}$ Calculated to 1,060 feet.
$i$ Calculated to 1,180 feet.
$j$ Taken by Prof. R. Williams with thermometer from U. S. Weather Bureau.
$k$ Average for 23 years.
$l$ Average for 61 years.
$m$ Approximate mean annual.
$n$ Calculated to 1,135 feet.

- Ohservation by superintendent in 1899 with Darton thermometer down 24 hours. Dry hole. Limestone below 2,500 feet. Another 24 -hour observation gave $95^{\circ}$ and a short one gave $93^{\circ}$, all in the midst of working double shift.
$p$ La Salle average to 1908.
$q$ Udden, J. A., Artesian wells in Peoria and vicinity: Illinois Geol. Survey Bull. 8, p. 334, 1908.
${ }^{r}$ All taken by Prof. J. A. Udden, 1898, with a Sargent thermometer reading $1^{\circ}$ higher than instruments in U. S. Weather Bureau, Davenport, Iowa.
- Daveriport average for 45 years.
$t$ Upper tow cased off, but casing may be corroded through as suggested by diminished pressure.
$u$ Calculated to 912 feet.
$v$ Maccoutah average for 26 years.
${ }^{w}$ Henry average for 23 years.
$x$ Peoria average for 26 years.
y Dixon average for 26 years.


## SUMMARY.

The most satisfactory observations of underground temperature in Illinois are those made at St. John and Streator, where thermometers were sunk to the bottoms of the holes. At St. John the instrument was left down 24 hours. When withdrawn it read $101^{\circ}$, indicating a rate of increase of $1^{\circ}$ in 83.5 feet. The temperatures in wells in Rock Island County were accurately taken by Prof. Udden. The variation which some of them show is possibly due largely to mingling of flows. However, the observations were made just after vigorous drilling, and some heat of impact may be manifest. The wells at Peoria give results so closely accordant that presumably the flows are all from the bottom or near it. The Hennepin well, "cased nearly to the bottom," and the Knoxville, Chicago, Marissa, and Galesburg wells give close approximations of rate if the water is not cooled by pumping. The temperatures in the Cairo wells may not be accurately determined.

Flows from the deep well at Aledo may be mixed, and if so the rate of $1^{\circ}$ in 168 feet is not valid. The Oak Park flow is undoubtedly mixed. The high temperature of water from the 950 -foot well at Pekin probably indicates a source at the bottom of the well, and the rate of $1^{\circ}$ in 42 feet is considerably higher than that of any other well in the State.

The rates of increase given in the list appear not to have any regional relations but vary from place to place. Galesburg and Knoxville, not far apart, apparently have rates of $1^{\circ}$ in 106 feet and 66 feet, respectively. Pekin is not far south of Peoria, where the average rate is $1^{\circ}$ in 60 feet.

## GEOLOGIC RELATIONS.

The Illinois wells penetrate limestone, shale, and sandstone of Carboniferous to Ordovician age, all lying nearly horizontal, without local disturbances of large amount and with no volcanic rocks. The Streator well is stated to be on a low anticline trending northwestward. The deep wells in the southern part of the State reach the oldest strata, most of them drawing from the St. Peter sandstone. At Peoria the shallower wells in the Niagara limestone show no material difference in rate of increase in temperature from the deeper ones, which draw from the St. Peter sandstone. A similar condition is presented by various wells of Rock Island County. The Cairo wells penetrate Cretaceous deposits to a depth of about 500 feet and then enter flint rock and sandstone of Mississippian age. If the flows are from about 500 feet, the base of the Cretaceous, the rate is $1^{\circ}$ in about 110 feet.

## INDIANA.

TEMPERATURES.
The few well temperatures on record for Indiana are not very significant, because no statement is furnished as to the depth from which the water comes. The following data may be of interest:

Temperatures in wells in Indiana.

| Location. . | Depth (feet). | Temperature (feet). | Mean annual air temperature ( ${ }^{\circ}$ F.). | Depth per degree of increase in temperature (feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Brown County: Nashville. | 530 | 56 | a 53 | 176 | Flows 300 gallons a minute. |
| Carroll County: |  |  |  |  |  |
| Delphi, several wiells..... | 898-928. | 62 | a 50.7 | 80 | Oil wells. |
| Flora.................. | 1,040 | 62 | 50.7 | 92 | Do. |
| Cass County: Logansport $b$. | 447 | 54 | c 50.9 |  |  |
| Crawford County: English.... | 887 | 60 | a 56 | 222 | 300-gallon flow. |
| Decatur County: Sandusky... | 830 | 60 | a 53 | 118 |  |
| Lake County: Hammond, 2 wells. | 1,780 | 60 | d 49 | 162 | Water próbably from shallow source or chilled by pumping. |
| Fort Wayne e.. .......... | 2,635 | 51.5 | $f 50.1$ |  |  |
| Wabash e... . . . . . . . . . . | 2,270 | 50.5 |  |  |  |
|  | 1,923 | 81 | g 54.6 | 73 |  |
| Laporte County: Michigan City. | 833 | 57 | h 49 | 104 | Flows 300 gallons a minute. |

$a$ From U. S. Weather Bureau map.
${ }^{b}$ Capps, S. R., Underground waters of north-central Indiana: U. S. Geol. Survey Water-Supply Paper 254, p. 93, 1910.
c Average for 36 years.
$d$ Average for 29 years.
e U.S. Geol. and Geog. Surveys W. 100th Mer. Rept., vol. 1, p. 209, 1889.
$f$ Average for 20 years.
$g$ Average for 26 years.
$h$ Average at Laporte.

## SUMMARY.

The rate in Indiana probably is not indicated accurately by these data except possibly by those for the Carroll County, Sandusky, Terre Haute, and Michigan City wells, which give figures that appear reasonably accordant.

## IOWA.

## TEMPERATURES.

A report of the State Geological Survey ${ }^{1}$ and a report by Norton and others ${ }^{2}$ give temperatures of flows from many deep wells in Iowa. Doubtless most of the figures are correct, but for most of the wells either the flows are known to be from various depths or no facts are given as to their source. Some of the wells are pumped, a condition which is unfavorable for accurate results, especially if the air lift is used. The 3,000 -foot well at Des Moines is one in which many flows mingle and the water is pumped to the surface. Temperatures of flows from several of the deep wells at Davenport were taken for me by J. A. Udden in 1898, but unfortunately in most of the wells there are mingled flows. The well at Ames was tested by S. W. Beyer, ${ }^{3}$ who lowered a Miller-Casella thermometer and took readings every 100 feet.

[^19]Temperatures in wells in Towa.

| Location. | Total depth (feet). | Flow (gallons per minute). | Tem-perature ( $\left.{ }^{\circ} \mathrm{F}.\right)$. | Mean air tem-perature ( ${ }^{\circ} \mathrm{F}$.). | Depth juer degree of increase in tem-perature (feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Allamakee County: |  |  |  |  |  |  |
| New Albin..... | 500 |  | 51 | 45.7 |  |  |
| Lansing. | 675,748 |  | 50 | 45.7 |  | Source of water not given. |
| Waukon, 3 miles north of. $a$ | 396 | Pumps many. | 52 | ${ }^{6} 47$ | 80 | Sandstone 347 to 396 feet, with water within 137 feet of surface. |
| Blackhawk County: |  |  |  |  |  |  |
| Waterloo. | 1,360 |  | 56 | 46.8 |  | Source of water not given. |
| Do. | 1.365 |  | 54 | 46.8 |  |  |
| Cedar County: Tipton c $d$, e | 2, 6902 | Pumps | 57 | $f 48.3$ | 118 | Water pumped 65 feet. "NQ water below 1,290 feet"; probably all from St. Peter sandstone, 1,030 to 1,070 feet. |
| Clayton County: McGregor.c | 520 | Many. | 52 | 45 | 974 | Flows. |
| Clinton County: Clinton.... | 1,605 | 500 | 72 | 4 S .6 | 69 | Probably cased. Many other wells with mixied flows. |
| Henry County: Mount Pleasant Asylum. $h$ | 1,125 | 165 | 62 | $i 50.2$ | 84 | Water pumped 30 fect. Water from 990 feet. |
| Des Moines County: $j$ Burlington. | 509 | 15 | 56 | k 50.5 | 93 |  |
| Do.. | 450 | Many. | 60 | 50.5 | 47 | Flow presumably from bottom. |
| D0. | 852 | 500 | 64.5 | 50.5 | 61 | Some water also at 430, 700, and 800 feet. |
| Do | 460 | 40 | 60 | 50.5 | 48 | Flow probably from bottom. |
| Iowa County: Amana Society.l | 1,640 | 100 | 68 | m 47.9 | 81 | Cased to 400 fect. Main flow begins at 1,200 feet; gradual increase to 1,640 feet. |
| Lee County: <br> Fort Miadison: |  |  |  |  |  |  |
| Atlee well $n$ | 740 | Many. | 64 | 051.9 | 57 | Water pumped 85 feet. |
| Paper compan | 689 | 600 | 62 | 51.9 | 62 | Water from about 680 feet. |
| $\cdots$ Do. | 681 | $200+$ | 65 | 51.9 | 43 | Water from 607 feet. |
| Atchison, Topeka \& Santa Fe Ry. ${ }^{n}$ | 764 | Many. | 60 | 51.9 | 77 | Water probably from sandstone, 692 to 756 feet. |
| Keokuk: $\boldsymbol{P}$ |  |  |  |  |  |  |
| Brewery. | 700 | Many. | 65 | Q 51.4 | 51.5 |  |
| Jottery company | 701 | 250 | 60 | 51.4 | 73 | First sandstone 628 to 701 feet. |
| Y. M. C. A.... | 769 | 350 | 64 | 51.4 | 55.5 | Water mainly from 700 feet. |
| Pickle company | 710 | 250 | 64 | 51.4 | 42 | Water from 530 fect. |
| Carter Co. | 661 | 5 | 61 | 51.4 | 8 | Water mainly from 648 to 661 feet. |
| Mooar ${ }^{\text {r }}$. | 800 | 165 | 67 | 51.4 | 51 | Cased to 600 fect; water from 800 feet. |
| Polk County: <br> Des Moines city pa |  |  | 70.0 | 849.3 | 109 |  |
| Des Momes city park | 2,250 |  | 70.0 | 849.3 | 109 | Beyer, $t$ |
| Do. | 2,000 |  | 69.4 | 49.3 | ${ }_{97} 9{ }^{2}$ |  |
| Do | 1,750 |  | 69.5 | 49.3 | 87 |  |
| Do | 1,500 |  | 69.1 | 49.3 | ; ${ }^{0}$ |  |
| Do | 1,250 |  | 68.3 | 49.3 | 66 |  |
| Do. | 1,000 |  | 65.4 | 49.3 | 62 | Rate 250 to 2,250 feet is 138 feet to $1^{\circ}$. |
| Do. | 750 |  | 60.5 | 49.3 | 67 | Rate 250 to 1,000 feet is 75 feet |
| Do. | 500 |  | 58.2 | 49.3 | 56 | Rate 1,000 to 2,250 feet is 272 feet |
| Do. | 250 |  | 55.5 |  |  |  |
| $a$ U. S. Geol. Survey Water-Supply Paper 293, p. 252, 1912. |  |  |  |  |  |  |
| 6 Dubuque average minus $0.9^{\circ}$ : |  |  |  |  |  |  |
| c Iowa Geol. Survey Rept., vol. 6, pp. 185, 189, 200, 266-267, 1896. |  |  |  |  |  |  |
| d Mead, D. W., Hydrogeology of the Upper Mississippi Valley: Assoc. Eng. Soc. Jour., vol. 13, p. 368, 189.4. e U. S. Geol. Survey Water Supply Paper 293, p. 374, 1912. |  |  |  |  |  |  |
| $f$ Cedar Rapids average plus $0.6^{\circ}$. |  |  |  |  |  |  |
| o Postville average. |  |  |  |  |  |  |
| ${ }^{h}$ lowa Geol. Survey Rept., vol. 6, pp. 320-321, 1896; U. S. Geol. Survey-Water Supply Paper 293, p. 535, 1912. |  |  |  |  |  |  |
| $i$ Average for 36 years. |  |  |  |  |  |  |
| j U. S. Geol. Survey Water-Supply Paper 293, pp. 528-529, 1912. |  |  |  |  |  |  |
|  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |
| $n$ U. S. Geol. Survey Water-Supply Paper 293, p. 563, 1912. - Keokuk average minus $0.4^{\circ}$. |  |  |  |  |  |  |
| p U. S. Geol. Survey Water-Supply Paper 293, pp. 564-566, 1912. |  |  |  |  |  |  |
| $q$ Average for 46 years. |  |  |  |  |  |  |
| t Average for 38 years. ${ }^{\text {t }}$ ( Geol. Survey Water-Supply Paper 293, pp. 734-738, from Iowa College Water-Supply; Ames, 1897. |  |  |  |  |  |  |

Temperatures in wells in Jowa-Continued.

| Location. | Total depth (feet). | Flow (gallons per minute). | Tem-perature ( ${ }^{\circ} \mathrm{F}$.). | Mean annual air tem-perature ( ${ }^{\circ} \mathrm{F}$.). | Depth per degree of increase in tera-perature (feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Pottawattamie County: Council Bluffs. ${ }^{a}$ | 1,280 | 105 | 62 | b 50 | 83 | Flow from 1,000 to 1,100 feet. |
| Scott County: <br> Davenport: $c$ |  |  |  |  |  |  |
| Glucose works, three wells. | 1,969 | 'Many. | 66.5 | d 49.4 | 116 | Flows probably mingle. |
| Malting company*. . | 800 | Many. | 60 | d 49.4 | 75 | Cased to bottom. ${ }^{\text {c }}$ |
| Do............ | 1,653 | Many. | 64 | d 49.4 | e 95 | First flow from 1,385 to 1,585 feet. |
| Witts Bottling Co.. | 780 | Many. | 59.5 | d 49.4 | 77 |  |
| Woolen mills... | 1,050 | Many. | 60.5 | d 49.4 |  | Flows probably mingle. |
| Ice factory... | 1,050 | Many. | 60 | d 49.4 |  | Do. |
| Kimball House | 1,050 | Many. | 59.5 | d 49.4 |  | Do. |
| Story County: <br> Ames. | 2, 100 | …...... | 63.4 | $f 46.8$ | 127 | Thermometer sunk by S. W. Beyer. $g$ |
| Do. | 2,000 |  | 64.1 | $f 46.8$ | 116 | Do. |
| Do. | 1,500 |  | 60.9 | $f 46.8$ | 107 | Do. |
| Do. | 1,000 |  | 57.0 | $f 46.8$ | 98 | Do. |
| Do. | 500 |  | 53.0 | $f 46.8$ | 81 | Do. |
| Do. | 200 |  | 51.2 | $f 46.8$ |  | Rate from 200 to 2,100 feet is 156 feet to $1^{\circ}$. |
| Nevada. . | 940 | 200 | 59 | $f 46.8$ | 77 | Cased 810 feet. Water at 940 feet pumped. |
| Wapello County: Ottumwa: $h$ |  |  |  |  |  |  |
| Water company.... | 2,047 | 700 | 70 | $i 49.3$ |  | Water mostly from 1,015 feet. |
| Do............. | 2,047 | 300 | 70 | $i 49.3$ |  | Cased to 1,200 feet and from 1,705 to 2,047 feet. |
| Morrell Co. | 1,554 | 1,000 | 64 | $i 49.3$ | 74 | Water from 1,085 feet. |
| Iron works | 1,150 | Many. | 62 | $i 49.3$ | 82 | Water from 1,040 feet. |
| Y. M. C. A | 800 | 33 | 65 | 249.3 |  | Source of flow not given. |
| No.3............. | 1,702 | 1,500 | 67 | 149.3 | 77 | Cased to 1,360 feet. |
| Washington County: Washington. $j$ | 1,611 | 95 | 72 | $k 49.4$ | 71 | Air lift pumps water 54 feet. Water from various depths; much at 1,115 feet, main supply at 1,600 feet. |

$a$ U. S. Geol. Survey Water-Supply Paper 293, p. 951, 1912.
$b$ Omaha, Nebr., average for 46 years.
c Taken by J. A. Udden, 1898, with a Sargent thermometer reading $1^{\circ}$ higher than the one in the Weather Bureau at Davenport.
$d$ Average for 45 years.
$e$ Calculated to 1,385 feet.
$f$ Boone average of 14 years.
$g$ U.S. Geol. Survey Water-Supply Paper 293, pp. 734-738, from Towa College Water Supply, Ames, 1897.
$h$ Iowa Geol. Survey Rept., vol. 6, pp. 317-320, 1897; U. S. Geol. Survey Water-Supply Paper 293, pp. 606 609, 1912.
i Mean of Keosauqua and Oskaloosa.
$j$ Average for 35 years.
${ }_{k}$ U.S. Ceol. Survey Water-Supply Paper 293, p. 613, 1912.

## SUMMARY.

The data here given from the Iowa wells present considerable diversity in value. The Beyer determinations of temperature at 2,250 feet at Des Moines and at 2,100 feet at Ames are valuable records. They were taken with a Miller-Casella self-registering instrument at intervals of 250 feet in the Des Moines well and 100 feet at Ames. The wells were full of water and had not been disturbed for a month or more. The rates of increase in temperature, $1^{\circ}$ in 109 feet and 127 feet, respectively, appear abnormally low and are considerably less than the rate in the 940 -foot Nevada well, 9 miles east of Ames. As shown in the table there were some remarkable local variations at different depths. The rates indicated by many
records in the southeast corner of the State vary greatly. The deeper Ottumwa wells give $1^{\circ}$ in $62 \frac{1}{2}, 77 \frac{1}{2}, 80$, and 87 feet, and the 800 -foot well at the Young Men's Christian Association Building indicates $1^{\circ}$ in 53 feet or less. The Amana data appear to be not worth considering. Some of the Davenport data appear to be consistent and indicate a rate of $1^{\circ}$ in 75 to 77 feet. The Keokuk, Fort Madison, and Burlington rates of $1^{\circ}$ in 42 to 73 feet appear to be reliable, but the precise rate for each place is not deducible from the data given. The Mount Pleasant and Washington rates are reasonably consistent with the figures for the adjoining counties. The McGregor and Waukon wells, in the northeast corner of the State, although of moderate depth, appear to afford reliable data. The same is true of Council Bluffs, in the extreme western part of the State, where the rate of increase is the same as in the wells at Omaha, across Missouri River- $1^{\circ}$ in 83 feet.

## GEOLOGIC RELATIONS.

The Iowa wells penetrate limestones, shales, and sandstones of Carboniferous to Cambrian age. They lie nearly horizontal, but most of the State is occupied by a broad, open syncline whose axis extends from east and northeast to southwest and passees near Muscatine, Oscaloosa, Indianola, Osceola, and Bedford. The Algonkian rocks are deeply buried, for they were not reached by the 2,696 $\frac{1}{2}$-foot hole at Tipton, and no igneous rocks or marked local disturbances are known to be near the wells of which data were recorded. The deeper borings and those in the northeast corner of the State draw from the sandstones of Cambrian or Ordovician age, especially the St. Peter sandstone, which is the source of supply at Tipton and Washington and is about 1,300 feet below the surface at Ottumwa, 1,100 feet at Davenport, 950 feet at Burlington, and 900 feet at Keokuk. The Ames, Amana, Davenport, and Des Moines holes penetrate the Jordan sandstone; the Nevada well draws from much higher limestones of Silurian age. The shallow Burlington and Fort Madison wells draw from strata considerably above the Jordan sandstone. When these facts are considered in connection with the rates of increase given in the table there appears to be no connection between the geologic horizon and the rate, so far as is indicated by the data available.

The abrupt change of rate from $1^{\circ}$ in $\dot{75}$ feet to $1^{\circ}$ in 272 feet at a depth of about 1,000 feet in the Des Moines well occurs just below the base of the Devonian.

## KANSAS.

## TEMPERATURES.

Sixteen temperature determinations were obtained in Kansas, part of them by me and my assistants, who lowered thermometers to the bottom of the wells. Others are given in publications, notably in the report on mineral water. ${ }^{1}$ The following are the data:

Temperatures in wells in Kansas.

| Location. | Total depth (feet). | $\underset{\substack{\text { Flow } \\ \text { (gallons } \\ \text { per min- }}}{\text { and }}$ $\underset{\text { ate). }}{\text { per min- }}$ | Tem-pera$\left({ }^{\circ} \mathrm{F}\right.$.). | Mean annual air tem pera${ }_{(0}{ }^{\circ} \mathrm{F}$.) ( $\left.{ }^{\circ} \mathrm{F}.\right)$. | Depth per deincreas in tem-pera(feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Allen County: Iola a Bourbon County: Fort Scott $b$ | 720 621 | Many. | 61. 67.5 | c 56 | 54 | Gassy water from below 626 feet. Good flow, begins at 510 feet. |
|  | 700 | Pumps. | 67 | 56 |  |  |
| Cherokee County: Columbus $b$. | 1,400 |  | 75.2 | c 56.5 | 75 | Cased to bottom. |
| Crawford County: Cherokee d. | 913? |  | 71.5 | c 56.4 | 602 | Cased 300 feet. Water pumped |
| Girard $f$. | 980 |  | 75 | 956.3 | 52 | Cased 376 feet. Water pumped from 150 feet. |
| Decatur County: Kanona.. | 1,540 |  | 89.7 | ${ }^{\text {b }} 51.6$ | $40 \frac{1}{2}$ | In progress July, 1903. Thermometer sunk by N. H. Darton. |
| Douglas rence. $h$ County: Law- | 1,400 |  | d 65.5 |  |  | Flows from depth not given. |
| Ellsworth County: Ellsworth. $j$ | 730 |  | 70 | ${ }^{*} 55$ | 49 | Thermometer sunk by C. A. Fisher in 1900, 2 t hours. |
| Finney County: Garden | 435 |  | 63.5 | 54 | 146 | Thermometer sunk by W. D. Johnson, July, 1898. |
| Meade County: Meade..... | 325 |  | 71.5 | m 56.5 | 22 | Thermometer sunk by W. D. Johnson in July, 1898. Water was within 42 feet of surface. |
| Do. | $265 \frac{1}{2}$ |  | 61.7 | 56.5 | 51 | Flows. |
|  | 610 2,775 |  | 69 73 | $n 57.7$ 57.7 | 54 181 | Oil well not in use. Thermometer sunk by E. G. Woodruff. Abandoned boring. Thermom- |
|  |  |  |  |  |  | eter sunk by E. G. Woodruff. |
| Morton County: Richfield $q$ | 600 | 6.3 | ${ }_{66}^{62}$ | $\begin{aligned} & \text { p} \\ & r \\ & 56 \end{aligned}$ | 57 | Source of flow not given. ${ }^{\text {Flows from } 570}$ feet. In red |
| Neosho County: St. Pault.. | 700 |  | 55 |  |  | beds 265 to 600 feet. Gassy water. |
| Pawnee County: Larned s.. | 750 | 250 | 65 | $t 54.3$ | 70 | Pressure 23 pounds. Brine from 743 feet. Flow also at 430 feet cased off |
| Do. | 743 | 240 | 6.5 | 54.3 | 693 | Thermometer lowered by C. A. Fisher, 1900. |

a Kansas Univ. Geol. Survey, vol. 7, p. 251, 1902.
b İem, p. 265.
$c$ Average for 10 years.
d Kansas Univ. Geol. Survey, vol. 7, p. 39, 1902.
$e$ Kansas Univ. Geol. Survey, vol. 7, p. 268, 1902.
$f$ Kansas Univ. Geol. Survey, vol. 7, p. 151, 1908.
$g$ Columbus average minus $0.2^{\circ}$.
$h$ Norton average for 18 years.
$i$ Average for 20 years.
; On ranch of J.Cochran, SW. $\frac{1}{4}$ NW. $\frac{1}{2}$ sec. i5; T. 15, R. 8.
$k$ Salina average minus $0.1^{\circ}$.
1 A verage for 21 years.
$m$ A shland average for 18 years.
$n$ Independence average minus $0.1^{\circ}$.
o Kansas Univ. Geol. Survey, vol. 7, p. 150, 1902.
$p$ A verage for 38 years.
${ }_{q}$ q Senate Committee on Use and Reclamation of Arid Lands Rept., vol. 4, p. 60, 1890.
$r$ Ulysses average minus $1^{\circ}$.
${ }^{8}$ Kansas Univ. Geol. Survey, vol. 7, p. 158, 1902.
$t$ Marksville average minus $0.2^{\circ}$.
${ }^{1}$ Bailey, E. H. S., Kansas Univ. Geol. Survey, vol. 7, 343 pp., 1902.

Temperatures in wells in Kansas-Continued.

| Location. | Total depth (feet) | $\xrightarrow{\text { Flow }}$ (gallons per minute). | $\begin{aligned} & \text { Tem- } \\ & \text { pera- } \\ & \text { ture } \\ & \left({ }^{\circ} \mathrm{F} .\right) . \end{aligned}$ | $\begin{gathered} \text { Mean } \\ \text { annual } \\ \text { air tem- } \\ \text { pera- } \\ \text { ture } \\ \left({ }^{\circ} \mathrm{F} .\right) . \end{gathered}$ | Depth per deincrease in tem-pera(feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Rice County: <br> Lyons Gas \& Oil Co... | 950 |  | 714 | ${ }^{\text {a } 54.4}$ | 56 | Salt shaft. Thermometer lowered by C. A. Fisher, 1900. Down three hours. |
| Lyons, salt mine. | 1,083 |  | 69 | 54.4 | 74 | Saltshaft. Thermomoter placed in bottom by C. $\Lambda$. Fisher. |
| Sterling, NE. $\frac{1}{8}$ SW. $\frac{1}{3}$ sec. 22. | 946 |  | $71 \pm$ | 54.4 | 56 | Thermometer lowered by C. A. Fisher, 1900. Down 4 hours. Water is within 3 feet of surface. Rock salt and shale 710 to 946 feet. |

a Ellinwood average for 21 years.

## SUMMARY.

Most of the Kansas observations are in fairly close accord as to rates of increases in temperature. The principal exception is that of the Caney boring, which shows a temperature of only $73^{\circ}$ at a depth of 2,775 feet. Possibly this is due to gas escape. This hole and those at Independence, Coffeyville, Fort Scott, Girard, Cherokee, and Columbus are all in the southeast corner of the State. The average rate exclusive of Caney is $1^{\circ}$ in slightly less than 60 feet. The group of four holes in central Kansas have rates of $1^{\circ}$ in 48.7 to 56.2 feet, but the rate decreases to $1^{\circ}$ in 74.2 feet in the deep shaft at the Lyons salt mine. Four wells in the western part of the State have rates of $1^{\circ}$ in 41,47 ; 51 ; and 57 feet; the first and third represent careful measurements. There is a great difference in the rates of the two wells at Meade as determined by Johnson.

GEOLOGIC RELATIONS.
The wells in Kansas penetrate a variety of sedimentary formations, but the strata are nearly horizontal and no igneous rocks are known to exist near any of the holes in the list. The wells in the southeast corner of the State are in limestone, shale, and sandstone of Carboniferous age, some of the deeper ones penetrating the upper part of the Mississippian. The Ellsworth, Lynn, and Sterling wells are in the salt series of the Permian. The Richfield well is in the red beds of the Permian, Garden City, and Kanona in the Dakota sandstone (Upper Cretaceous). The Meade wells reach red beds under a thick mantle of sand and gravel of Tertiary age. The high rate of increase in the 325 -foot well at Meade, $1^{\circ}$ in 21.7 feet, may be related to a fault supposed to extend along one side of the artesian basin.

## KENTUCKY.

## TEMPERATURES.

Very few records of well temperatures have been reported from Kentucky. The most important one is that of the 2,086 -foot well sunk in 1857-58 at Louisville, which indicates a rate of increase of $1^{\circ}$ in 106 feet if the water comes from the bottom.

Temperatures of wells in Kentucky.

| Location. | Depth (feet). | Flow (gallons per minute). | Tem-perature $\left({ }^{\circ} \mathrm{F}\right)$ | $\begin{aligned} & \text { Mean } \\ & \text { annual } \\ & \text { air } \\ & \text { tem- } \\ & \text { pera- } \\ & \text { ture } \\ & \left({ }^{\circ} \mathrm{F} .\right) . \end{aligned}$ | Depth per degree of increase in tem-perature (feet). | Remarks. |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Jefferson County: Louisville. <br> Jessamine County: Harrisburg (on Kentucky River). Meade County: Fountain - Farm.e | 2,086 485 428 | $264$ | a $76 \frac{1}{2}$ c $82 \frac{3}{2}$ 77 64 | b 56.8 56.8 d 56.1 $f 56.2$ | 106 81 23 57 | Flow of poor water. Gas and brine. | - |

$a$ Owen, D. D., Geological reconnaissance of Arkansas, p. 61, 1860. Another report gives a temperature of $86 \frac{1}{4}^{\circ}$, which indicates a rate of $1^{\circ}$ in 71 feet.
$b$ Average for 44 years.
c Am. Jour. Sci., 2d ser., vol. 27, pp. 174-178, 1859.
d Richmond average for 22 years.
$e$ Report on the occurrence of petroleum, natural gas, and asphalt rock in western Kentucky, based on examinations made in 1888 and 1889: Kentucky Geol. Survey, p. 179, 1891.
$f$ Irvington average for 19 years.

## SUMMARY.

The old Louisville artesian well with a large flow should give a reliable indication of the rate of increase, but two widely different temperatures are reported. The result for the Harrisburg well is dubious. The Fountain Farm well is gassy, but the rate is reasonable.

## GEOLOGIC RELATIONS.

The strata penetrated by the wells listed in the table are limestone and shale of Paleozoic age on the west slope of the Cincinnati arch, and the dips are very low. The Louisville hole penetrates Devonian, Silurian, and Ordovician strata. The Fountain Farm well penetrates Mississippian limestone to the Devonian black shale.

## LOUISTANA.

## BELLE ISLAND.

The most significant determination of underground temperature in Louisiana is one made for me by I. N. Knapp in February, 1909, during the sinking of a 3,171 -foot boring at Belle Island, near Morgan City. A Darton thermometer was lowered to various depths with the following results.:

Temperatures at different depths in hole at Belle Island, La.

| Depth (feet). | $\begin{aligned} & \text { Temper } \\ & \text { ature } \\ & \left({ }^{\circ} \mathrm{F} .\right) . \end{aligned}$ | Remarks. |
| :---: | :---: | :---: |
| 427 | 75 | In salt water. |
| 764 | 79 | 24 hours after drilling ceased. |
| 975 | 82 | Dry salt. |
| 1,625 | 82 | Dry hole. |

At 427, 764, and 975 feet the instrument was down 36 hours. It was sent down again 24 hours later, obtaining closely accordant results. At 1,625 feet there was much gas, which caused a diminished temperature, so no observations were made at greater depths. The beds below 211 feet were salt or salt bearing. The rates of increase were as follows, as calculated from a mean annual temperature of $68.7^{\circ}$ (Franklin average for 24 years plus $0.2^{\circ}$ ):

Depth per degree of increase in temperature in hole at Belle Island, La.

|  | Feet. |  | Feet. |
| :---: | :---: | :---: | :---: |
| Surface to 427 feet. |  | 427 to 764 feet. |  |
| Surface to 764 feet. | 74 | 427 to 975 feet. | . 78 |
| Surface to 975 feet. | 73 | 764 to 975 feet | 70 |

## MISCELLANEOUS DATA.

Although there are 'many flowing wells in Louisiana they afford but few reliable data as to the geothermal gradient. Harris ${ }^{1}$ gives results of his own observations for many wells, notably those in St. Tammany and Tangipahoa parishes, and Veatch ${ }^{2}$ gives records from various sources, but only in á few of the records is information given as to the source of the water. However, as the pressures are such as to indicate considerable depth, and as the materials are soft, it is probable that the wells are cased to or nearly to the bottom. If so, the temperature records are valuable as indicating approximately the rate of increase with depth.

[^20]Temperatures in wells in Louisiana.

| Location. | Depth (feet). | Flow (gallons per minute). | Tem-perature ( ${ }^{\circ} \mathrm{F}$.). | Mean annual air tem-perature ( ${ }^{\circ} \mathrm{F}$ ). | Depth per degree of increase in tem-perature (fect). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Caddo Parish: Shreveport. | 996 | Many. | 84.0 | $a 65.1$ | 53 | Brine with gas. |
| Calcasieu Parish: Welsh..... | 190 | 1,200 | 71.5 | ${ }^{\text {b } 67.4}$ | 46 |  |
| Catahoula Parish: Leland. | 1,550 | 10 | 86.5 | c 66.8 | 78 | 1)0. |
| Orleans Parish: <br> New Orleans. | 1,229 | 55 | 81.5 | d 68.2 | 92 | Do. |
| New Orleans, 1 mile west of. | e 900 | 12 | 79 | 68.2 | 83 |  |
| Ouachita Parish: Monroc. . . | 385 | Many. | 71 | $f 65.6$ | 71 |  |
| De Soto Parish: Frierson.... St. Tammany Parish: | 1,500 | 10 | 70 | g 64.8 |  | Flows at 241 to 281 and 998 to 1,275 feet; probably mingle. |
| Covington: <br> Dummet | 572 | 20 | 74 | $h 66.8$ | 80 |  |
| Dixon Academy |  | Many. | 72.6 | 66.8 | 80 |  |
| Claiborne. . . . . | 630 | 30 | 73 | 66.8 | 101 |  |
| Courthouse, 1901. |  |  | 73 | 66.8 |  |  |
| Courthouse, 1903 |  |  | 72.4 | 66.8 |  |  |
| John Dutch..... | 600 | 20 | - 74 | 66.8 | 83 |  |
| H. Hallen. . | 520 | 30 | 72 | 66.8 | 100 |  |
| Hernandez, 2 miles north of. | 610 | 38 | 73 | 66.8 | 98 |  |
| Hernandez............... | 610 | 40 | $72 \frac{1}{4}$ | 66.8 | 111 |  |
| Maison Blanche. . . . . . . . | 480 | 16 | $72 \stackrel{1}{4}$ | 66.8 | 87 |  |
| Abita Springs. | 545 | 54 | 73 | 66.8 | 88 |  |
| Do..................: | 574 | 27 | 73 | 66.8 | 93 |  |
| Do.. | 526 | 56 | 74 | 66.8 | 73 |  |
| Mandeville. | 217 | 28 | 69.5 | ${ }^{i} 67$ | 87 |  |
| Do. | 247 | Many. | 71 | 67 | 62 |  |
| Chinchuba................ | 325 | Many. | 72 | 67 | 65 |  |
| Sabine Parish: Negreet. . . . . | 630 | 30 | 70 | j65.4 | 137 | From 630 feet, but possibly |
| Tangipahoa Parish: Ponchatoula | 232 | 3 | 71 |  |  | higher flows mingle. |
| Ponchatoula, 3 miles north of. | 170 | 10 | 69.5 | 67 67 | 70 |  |
| Hammond............... | 272 | 20 | 70.5 | 67 | 80 |  |
| Do. | 225 | 21. | 70 | 67 | 75 |  |
| Do. | 330 | 24 | 71.5 | 67 | 73 |  |
| Do. | 265 | 7 | 71 | 67 | 66 |  |
| Do. | 265 | 28 | 71 | 67 | 66 |  |
| Do. | 377 | 24 | 71 | 67 | 94 |  |
| Do................... | 300 | 25 | 71 | 67 | 75 |  |
| Hammond Ice Co........ | 340 |  | 72 |  | 68 |  |
|  |  |  |  |  |  |  |
| $1 \frac{1}{1}$ miles south-southwest of. | 212 | 8 | 69 | 67 | 106 |  |
| Do......... | 140 | 3 | 69 | 67 | 70 |  |
| $1 \frac{1}{2}$ miles south of...... | 309 | 30 | 72 | 67 | 62 |  |
| $\frac{1}{2}$ mile southeast of... | 318 | 45 | 70.5 | 67 | 91 | , |
| Do........ | 302 | 24 | 70.5 | 67 | 86 |  |
| 1 mile south of. . . . . . | 1380 | 15 | 70.6 | 67 | 1.06 |  |
| 1 mile east of......... | 235 |  | 69 | 67 |  |  |
| $1 \frac{1}{2}$ miles south-southeast of. | 298 | $5 \frac{1}{2}$ | 703 | -67 | 80 |  |
| Vermillion Parish: Gueydan, 3 miles southwest of. | 190 | 8 | 73 | $m 68.3$ | 40 |  |
| Winn Parish: Drake Salt Works. | n 1, 011 | 18 | 75 | - 66. | 123 | Brine from Cretaceous. |

a Average for 45 years.
b Lake Charles average.
c Natchez, Miss., average for 29 years.
d A verage for 45 years at Station 1.
e Cased to bottom.

- $f$ Monroe average for 28 years.

G Grand Cane average for 22 years.
$h$ Average for 24 years.
${ }^{6}$ Covington average plus $0.2^{\circ}$.
f Average at Nacodoches, Tex.
Hammond average for 24 years.
6 Water from 340 to 380 feet.
$m$ A bbeville average.
n Veatch, A. E., Louisiana Geol. Survey Rept. for 1902, pp. 51-64, 1903.
o Mean at Dodson, 1916.

Water from pumped wells 197 to 315 feet deep in Bossier Parish is reported to have temperatures from $64^{\circ}$ to $66^{\circ}$; from a 225 -foot well at Belcher, in Caddo Parish, $66^{\circ}$; from wells 137 to 280 feet deep at Shreveport, $66^{\circ}$ to $68^{\circ}$; and from several wells 115 to 368 feet deep in Webster Parish, $65^{\circ}$ to $66^{\circ} .^{1}$ All these figures are so near the mean annual air temperature that they throw no light on the rate of increase.

## SUMMARY.

None of the rates of the deeper wells given in the list are close approximations. In the Frierson and Negreet wells probably the mingling of flows gives a lower temperature than the water from the lower flow would show. The wells of Tangipahoa Parish are too shallow to be reliable within close limits, but they give an average rate of $1^{\circ}$ in very nearly 80 feet.

## GEOLOGIC RELATIONS.

All the wells in Louisiana are in the great Coastal Plain sedimentary series of the Mississippi embayment, but several horizons are tapped. The strata are sands, clays, and limestones of Cretaceous and Tertiary age, dipping at low angles in various directions. In northern Louisiana there is a general dip to the east into a basin whose bottom is near Vicksburg, Miss., at a rate of generally from 7 to 10 feet to the mile, with few local variations, some of which are due to sharp local uplifts. This basin widens and deepens to the south and becomes a deep, wide trough containing a great thickness of Tertiary strata. The briny waters of the Shreveport well come from the Nacatoch sand in the Upper Cretaceous, and the brine at the Drake salt works is also from the Cretaceous.。 Brine from the Leland well is believed to be from the Eocene Yegua ("Cockfield") formation, 1,500 feet higher. The beds tapped by the shallow wells of Calcasieu, St. Tammany, and Tangipahoa parishes are high in the Tertiary.

## MARYLAND.

Only one well temperature was obtained from Maryland-that of the 140 -gallon flow of saline water from the deep well at Crisfield,
1 Somerset County, on the eastern peninsula. The flow comes from a depth of 1,025 feet in the Potomac group, and the temperature, $79^{\circ}$, was determined by Mr. John Buxton with a reliable thermometer sent to him for the purpose. As the mean annual temperature at Pocomoke City, not far away, is $57.4^{\circ}$, this determination indicates a rate of increase of $1^{\circ}$ in 47.4 feet.
${ }^{1}$ Veatch, A. C., U. S. Geol. Survey Prof. Paper 46, pp. 201-207, 217-223, 1806.

## MICHIGAN.

## SOURCES OF INFORMATION.

Temperature observations that have been made in deep borings and deep mines in Michigan throw important light on the rate of increase in temperature underground. Some of these have been made by A. C. Lane, ${ }^{1}$ in part with thermometers which I furnished. On account of the differences in the geology of the two areas it is best to present the results of observations made in southern and central Michigan separately from those of observations made in the area of older rocks in the northern peninsula. Lane has considered the probable mean annual temperature of the northern peninsula and also gives temperatures in shallow borings and in upper levels in certain copper mines.

## NORTHERN PENINSULA.

Many determinations of underground temperature are available for the northern part of Michigan, mainly in connection with the deep mines.

In 1895 Alexander Agassiz ${ }^{2}$ reported the results of a series of observations made in copper mines of the Calumet \& Hecla Co. on Keweenaw Point. This point, which extends 70 miles out into Lake Superior, consists of a series of old igneous and sedimentary rocks dipping $7^{\circ}-55^{\circ}$ and overlain by Cambrian sandstone. The mines are mostly in the older series. The temperatures were taken at intervals, but only those at 105 feet and 4,580 feet were published. Negretti and Zambra thermometers were used, placed in 10 -foot holes, slightly inclined upward, and left from one to three months, the holes plugged with wood and clay. The results were $59^{\circ}$ and $79^{\circ}$, a difference of only $20^{\circ}$ for the 4,475 feet, or an increase of $1^{\circ}$ in 223.7 feet. However, if the $59^{\circ}$ observation is disregarded and the increase to 4,580 feet calculated from a mean annual air temperature of $43^{\circ}{ }^{3}$ the result is $1^{\circ}$ in 127 feet, which is borne out by other observations. It is evident that the reading of $59^{\circ}$ at 105 feet was defective, because it is $16^{\circ}$ to $20^{\circ}$ higher than the mean annual air temperature.

In a letter to J. D. Everett in 1896, ${ }^{4}$ Agassiz stated that further observations had shown that the rate was "different from what it was believed to be when the preliminary announcement was made."

[^21]In 1886 Wheeler ${ }^{1}$ presented to the St. Louis Academy of Sciences an account of temperature observations in copper mines of Keweenaw Point. Wheeler's results are given in the following table:

Temperature observations in copper mines on Keweenaw Point, Mich.

| Mine. | Depth. | Temperature ( ${ }^{\circ} \mathrm{F}$.). |  | Temperatureat certain depths.a |
| :---: | :---: | :---: | :---: | :---: |
| Atlantic. | 907 | 51.6 | 993 | $43.6^{\circ}$ at 111 feet. |
| Central. | 1,950 | 61 | 101 | $42.6^{\circ}$ at 90 feet. |
| Conglomerate (Delaware or | 617 | 48.3 | 95 | $42.8{ }^{\circ}$ at 90 feet. |
| Quincy. | 1,931 | 58.5 | 122 | $43^{\circ}$ at 111 feet. |
| Osceola | 996 | 54.5 | 764 | $42.3^{\circ}$ at 136 feet. |
| Tamarack. | 2,240 | 62 | $110 \frac{1}{2}$ | $43^{\circ}$ at 136 feet. |

$a$ The figures in this column were used in calculating the rates of increase.
Wheeler states that the most satisfactory results were obtained in the Atlantic mine, where observations at intermediate stations gave an accordant rate. The Conglomerate mine is 36 miles farther east. The rate in the Tamarack mine may be vitiated by the proximity of active mining. All the mines were nearly dry. Wheeler suggests that the low rate of increase in this district may be due to the proximity of Lake Superior, with its great volume of cool water ( $38.8^{\circ}$ ).

Temperature observations in mines and holes in northern peninsula of Michigan. a

| Locality. | Depth (feet) | Temperature ( $\left.{ }^{\circ} \mathrm{F}.\right)$. | Mean annual air tem$\underset{\left({ }^{\circ} \mathrm{F} .\right)}{\mathrm{p} \text {. }}$ | Depth per degree of increase in tem$\underset{\text { (feet) }}{\text { perature }}$ (feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Freda well b | 950 | 55.6 |  | 84 | Temperature $45{ }^{10}$ at 100 feet. |
| Victoria mine ${ }^{\text {b }}$ | c 1,350 | 57. | $\pm 44$ | 104 | Eighteenth level. |
| Champion mine | 465 | 45.8 | 44 |  | 9 to 250 fect west of shaft. |
|  | 1,335 | 52 | 44 |  | Twenty-fourth level, 140 feet |
| Do. | 1,500 | 55 | 44 |  | End of twenty-sixth level. |
| Do | 1,650 | 56.5 | 44 | 111 | End of twenty-eighth level, 750 |
| Vulcan iron mine ${ }^{6}$. | 1,090 | 56 | 44 | 83 |  |
|  | 1,210 | 56 | 44 | 94 | Surface water $44^{\circ}$. |
| Republic iron mine | 1,153 | 55 | 44 | ${ }^{105}$ |  |
| Central copper mine. | 2,400 | ${ }_{61}^{59}$ | 44 | ${ }_{141}^{951}$ | In water. <br> Foot of No. 2 shaft, thirtioth |
| Centennial mine. |  | 62 | 44 | 180 | Foot of shaft. |
|  | (?) | 69 |  | (?) | Slope 3,100 feet long on Osceola |
|  |  |  |  |  | lode, $72^{\circ}$ during previous summer. |
| Olibway mine. | 680 | 50 | 44 | 113 | On 1,225-foot slope, $33^{\circ} \mathrm{dip}$. |
| Franklin, jr., mine | 1,600 | 61 | 44 | 94 | Twenty-first level cross cut. |
| Old Franklin...... | 3,200 | 62-63 | 44 | 169 | Third level; G. T'ope, observer. |

a Lane, A. C., The geothermal gradient in Michigan: Am. Jour. Sci., 4th ser., vol. 9, p. 435, 1900; Michigan Geol. Survey Ann. Rept. for 1901, pp. 244-251, 1902.
${ }_{b}$ Lane, A. C., The Keweenaw series of Michigan: Michigan Geol. and Biol. Survey, J'ub. 6, vol. 1, pp. 757-773, 1911.
$c$ Vertical depth not given.

[^22]Temperature observations in mines and holes in northern peninsula of Michigan.-Con.

| Locality. | Depth (feet). | $\begin{aligned} & \text { Temper- } \\ & \text { ature } \\ & \left({ }^{\circ} \mathrm{F} .\right) . \end{aligned}$ | Mean annual air temperature ( ${ }^{\circ} \mathrm{F}$.). | Depth per degree of increase in temperature (feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Tamarack | 4,400 | 84 | 44 | 110 | R. M. Edwards, obscrver. |
| Do. | 5,223 | 85 | 44 | 127 | Do. |
| Do. | 5,367 | 88-91.4 | 44 | $93 \frac{1}{2}$ | Bottom; observed by Lane in damp mud. |
| Tamarack, No. 5. | 4,662 | 82 | 44 | 123 | W. E.Parnall. |
| Do............ | 4,900 | 87 | 44 | 114 | J. Hull, mine inspector. |
| Quincy minea | 1,080 | 57 | 44 | 83 | Diamond-drill hole, 3 days. |
| Do...... | 2,941 | 64.5 | 44 | $143{ }^{\frac{1}{2}}$ | Diamond-drill hole, 5 days. |
| Do. | 2,941 | 67.5 | 44 | 121 | Breast of drift. |
| Do. | 3,090 | 68.5 | 44 | 126 | Do. |
| Do | 3,196 | 69 | 44 | 128 | Do. |
| Do | 3,403 | 71 | 44 | 126 | Do. |
| Do. | 3,620 | 75 | 44 | 117 | Driveway between shafts. |
| Do. | 3,816 | 79, 80 | 44 | 106 | Breast of drift. |
| Do...................... | 3, 875 | 76 | 44 | 121 | Diamond-drill hole, 40 hours. |
| Calumet \& Hecla stamp mill, Lake Linden, Houghton County, SW. $\frac{1}{4} \mathrm{sec} .6$, T. 55 , | 500 | 47.5 | 44 | 143 | Two wells, 1,502 and 1,508 feet. Flow 50 gallons. All in sandstones. |
| R. 32. |  |  |  |  |  |
| Do. | 1,000 | 51.5 | 44 | 133 |  |
| Do. | 1,490 | 55.5 | 44 | 130 | Rate indicated by difference between figures for 500 and 1,490 feet is 10 in about 124 feet. |
| Lake Superior mine, Ishpeming. | 900 | 50 | 44 | 150 | Thermometer 20 minutes in 5 foot hole in shaft. |

a The depths here given assume the surface to be 500 feet above Lake Superior, for in the original paper depths are given with the lake as the datum and the surface is irregular. Temperatures were all taken by S. Smillie, engineer.

The temperature of $57^{\circ}$ at 1,080 feet in the Quincy mine appears anomalously high when compared with the record at greater depths. The rate of increase calculated from the difference between the temperature at 1,080 feet and 3,875 feet is $1^{\circ}$ in 147 feet; from 2,941 to 3,875 feet it is $1^{\circ}$ in 81 feet on the basis of a temperature of $64.5^{\circ}$ in the bore hole, and $1^{\circ}$ in 110 feet on the basis of an observation of $67.5^{\circ}$ in the breast of a drift.

## CENTRAL AND SOUTHERN MICHIGAN.

Temperatures in deep wells in central and southern Michigan.

| Locality. | Depth (feet). | Flow (gallons per minute). | Tem-perature ( ${ }^{\circ} \mathrm{F}$.). | Mean annual air tem-perature ( ${ }^{\circ} \mathrm{F}$.). | Depth per degree of increase in tem-perature (feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Alpena County: Alpena a Bay County: | 698-711 | 200 | 53 | b 41.8 | 62 | Main supply from 698 to 711 feet; surely below 588 feet. |
| -Bay City c.. | 3,455 |  | 97 | d 46.5 | 681 | Thermometer 5688 down 8 hours. |
| Do...... | 2,934 |  | 90.1 | $\bullet 46.5$ | 67 | Thermometers 5688 and 5690 |
| Do. | 2, |  | 9 90.2 | 46.5 | $66_{2}^{1}$ | down half an hour. |
| Do. | 2, 282 |  | $\left\{\begin{array}{l}77 \\ 77\end{array}\right.$ | 46.5 46.5 | 75 | Thermometers 5688 and 5690 |
| Do. | 1,793 |  | 71 | 46.5 | 73 | Thermometer 4708 down 50 |
| Do | 1,304 |  | 65 | 46.5 | 70 | minutes. <br> Thermometer 4708 down 1 hour. |

a Lane, A. C., Am. Jour. Sci., 4th ser., vol. 9, p. 435, 1900.
$b$ Average for 44 years.
c Lane, A. C., op. cit., pp. 434-438. Test by C. A. Davis with Darton thermometers after drilling had ceased two months; bore full of heavy brine. The tests at 2,282 and 1,793 feet were regarded as too short in duration.
$d$ Average for 21 years.

Temperatures in deep wells in central and southern Michigan-Continued.

| Location. | Depth (feet). | Flow (gallons per minute). | Tem- pera- ture ( $\left.{ }^{\text {F }} \mathrm{F}.\right)$. | Mean annual air tem-perature ( ${ }^{\circ}$ F.). | Depth per degree of in crease in tem-perature (feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Benzie County: |  |  |  |  |  |  |
| Frankfort $a$. | 2, 200 | 280 | 58 | ${ }^{6} 44.9$ | c 61 | Water from 800 fect; surely be- |
| Dod | 2,200 |  | 56 | 44.9 | c 72 | low 590 feet. <br> Do. |
| Do. | 1,800 |  | 54 | 44.9 | c 88 | Do. |
| Rerrien County: Niles........ | 1,438 | 15 | 60 |  |  |  |
| Cheboygan County: <br> Cheboygan $e$ |  |  | 61.6 | $f 42.6$ |  |  |
| Cheboygan e............. | 1,380 |  | 73 | 42.6 42.6 | 711 89 | Near base of Salina formation. |
|  |  |  |  |  |  | is 118 feet to $1^{\circ}$. |
| Cheboygan waterworks... | 408 |  | ${ }^{\text {h }} 51.8$ | 42.6 | 44 |  |
| Burt Lake ${ }^{\boldsymbol{i}}$. . . . . . . . . . . | 464 | 4(?) | 51 | 42.6 | 55 |  |
| Crawford County: |  |  |  |  |  |  |
| Grayling 3. | 2,600 |  | 95.9 | ${ }^{\text {d }} 42.7$ | 49 |  |
| Do. | 2,376 |  | 93.8 | 42.7 | $46 \frac{1}{2}$ |  |
| Do. | 2,150 |  | 89 | 42.7 | $46{ }^{2}$ | Thermometer 7815 down 1 hour |
| Do. | 900 |  | 58.4 | 42.7 | $k 57$ | to $2 \frac{1}{2}$ hours. |
| Do. | 500 |  | 51.8 | 42.7 | 55 |  |
| Emmet County: Petoskey.... | 575 | 750 | 50 | $l 43.3$ | 86 |  |
| Genesce County: Flintm..... | 376 | 8 | 52 | $n 45.9$ | 61 | Water at 376 fect. |
| Grand Traverse: Traverse City.o | 417 |  | 51 | 45.9 | 82 |  |
| Gratiot County: |  |  |  |  |  |  |
| Almap................... | 2,863 | Flows. | 98 | 45.6 | $54 \frac{1}{2}$ | Temperature of water brought up in sand pump. |
| Alma, same well . . . . . . . | 900. |  | 69.6 | 45.6 | $37 \frac{1}{2}$ | Test by A.C.Lane, 1900. |
| Ingham County: |  |  |  |  |  |  |
| Lansing q................. | 1,400 400 | 1 | 58.5 53 | 46.7 46.7 |  | Saline water; source not known. |
| Ionia County: |  |  |  |  |  |  |
| Ionis $0 . .$. |  | Many. | 55 | ז 46.8 | 44 | Water from 362 feet. |
| Do. | $\left\{\begin{array}{c}340 \\ 540\end{array}\right.$ | \}...... | 52.3 | 46.8 |  |  |
| Do................... | 320 |  | 53.8 | 46.8 | 46 |  |
| Lenawee County: |  |  |  |  |  |  |
| Britton s | 1,617 |  | 67.6 | $t 48$ | 83 |  |
| Mackinac County: St.Ignace $u$ | 1,155 ${ }^{945}$ | Many. | 59 51 | 48 40.9 | 86 $v 57$ | Flow at 575-681 feet; some also |
| Mackinac County. St. Ig ace 4 , | 1,155 |  |  |  |  | at 1,040 fect. |
| Macomb County: Mount Clemens.w | 1,265 | ..... | 56 | $x 47.8$ | 79 | Water from 700 feet. |

a Michigan Geol. Survey, 1881-1893, p. 59.
${ }^{6}$ Average for 13 years.
c Calculated to 800 feet.
${ }^{d}$ Lane, A. C., Am. Jour. Sci., 4th ser., vol.9, p. 435, 1900.
e Taken by A.C. Lane (Geol. Soc. America Bull., vol. 13, 1903) with Darton thermometer 7812. Down 1 hour on bailer. Temperature $51.8^{\circ}$ at 408 feet.
$f$ A verage of 20 years to 1908.
$\sigma$ Tested by F. P. Rust, contractor, who determined other temperatures with thermometer 9114 as follows: 100 feet, $53^{\circ} ; 400$ feet, $51^{\circ} ; 700$ feet, $5,1.5^{\circ} ; 1,000$ fect, $55^{\circ} ; 1,300$ feet, $60^{\circ} ; 1,400$ feet, $60^{\circ} ; 1,700$ feet, $63^{\circ}$; 1,800 feet, $65^{\circ} ; 2,000$ feet, $65^{\circ} ; 2,300$ feet, $68^{\circ}$.
${ }^{n}$ Taken by A. C. Jane with Darton thermometer 7536.
iLeverett, Frank, U. S. Geol. Survey Water-Supply Paper 182, p. 363, 1907. Other wells with flows $46^{\circ}$ to $48^{\circ}$.
'Lanc, A. C., op. cit., and Michigan Geol. Survey Ann. Rept. for 1901, p. 250, 1902.
${ }^{k}$ Rate of increase from 500 to 2,600 feet is 48 feet to $1^{\circ}$.
$l$ Average for 27 years.
${ }^{m}$ Cooper, W. F., U. S. Geol. Survey Water-Supply Paper 102, p. 492, 1904.
$n$ Average for 28 years.

- Leverett, Frank, and athers, U. S. Geol. Survey Water-Supply Paper 183, pp. 16, 242, 324, 1907.
$p$ Michigan Geol. Survey Rept., 1881-1893, pp. 45-46, 1895 . This report is somewhat discredited by A. C. Lane in Am. Jour. Sci., 4th ser., vol. 9, p. 435,1900 , but accepted in Michigan Geol. Survey Rept. for i901, p. 249.
$q$ Michigan Geol. Survey Rept., 1881-1893, p. 66, 1895.
$r$ Grand Rapids average for 28 years minus $1.3^{\circ}$.
${ }^{4}$ Lane, A.C., Michigan Geol. Survey Rept. for 1901, p. 249, by probable misprint gives decener tempera. ture ( $53^{\circ}$ ) but with data indicating that it is $67.6^{\circ}$.
${ }^{i}$ Hillsdale average for 20 years plus $0.6^{\circ}$.
u Michigan Geol. Survey Rept. 1901, p. 248, 1902.
- Calculated to 575 feet.
${ }^{w}$ Cooper, W. F., op. cit., p. 498.
$x$ Average for 17 years.

Temperatures in deep wells in central and southern Michigan-Continued.

| Location. | Depth (feet) | $\begin{aligned} & \text { Flow } \\ & \text { (gal- } \\ & \text { lons } \\ & \text { per } \\ & \text { min- } \\ & \text { ute). } \end{aligned}$ | Tem-pera( ${ }^{\text {ture }}$. . | $\begin{array}{\|c} \text { Mean } \\ \text { annual } \\ \text { air } \\ \text { tem. } \\ \text { pera- } \\ \text { ture } \\ \left({ }^{\circ} \mathrm{F}\right. \text { F.). } \end{array}$ | Depth per of increase perature (feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Midiand County: Midland $a$. | 1,205 | Flow. | 62.5 | b 46.1 | 7312 |  |
| Ingersoll c. | . 625 |  | 52 | 46.1 | 105 |  |
| Muskegon, Ryerson well.. | 1,150 |  | 67.2 | ${ }^{\text {d }} 46.7$ | 56 | Plugged at 1,200 feet; full of water; flows from base of drift at 240 feet are $53^{\circ}$ to $53 \frac{3}{2}^{\circ}$. |
| Do. | 650 |  | 58.7 | 46.7 | 54 |  |
| Do. | 240 |  | 53.2 | 46.7 | 37 | Rate of increase from 240 to 1,150 feet is $1^{\circ}$ in 65 fect. |
| St. Clair County: e <br> St. Clair. | 1,635 |  | 71.2 | $f 45.3$ | 63 | Salt well; idle for months. |
| Sti. Clair, Oakland well.... | 1,200+ | P'mp'd | 69 | 45.3 | $56+$ |  |
|  | , 300 |  | 54.2 | 45.3 |  |  |
| Do. | 600 |  | 55.3 | 45.3 | 60 |  |
| Do. | 838 |  | 57.5 | 45.3 | 69 | Rate of increase from 300 to 838 feet is $1^{\circ}$ in 163 feet. |
| Saginaw County: |  |  |  |  |  |  |
| Saginaw, salt well. | 617 |  | 54 | $g 46.3$ | 80 |  |
| Do.............. | 531 |  | 51 | 46.3 | 113 | Rate of increase from 293 to 617 feet is $1^{\circ}$ in 81 feet. |
| Do. | 293 |  | 50 | 46.3 | 80 |  |
| Do................... | 102 |  | 47 | 46.3 |  | Rate of increase from 102 to 617 feet is $1^{\circ}$ in 75 feet. |

a Lane, A. D., Am. Jour. Sci., vol. 9, p. 434, 1900.
$b$ Average of Bay City minus $0.4^{\circ}$.
c Leverett, Frank, and others, U. S. Geol. Survey Water-Supply Paper 183, pp. 16, 242, 324, 1907.
$d$ A verage for 20 years.
e Michigan Geol. Survey Rept. 1901, p. 248, 1902.
$f$ Port Huron average for 42 years.
$g$ Average for 22 years.

## MINNESOTA.

Few temperatures of flows from wells in Minnesota have been reported and the source of the flows is uncertain. The large flow from a 467 -foot well at Winona is stated to have a temperature of $54^{\circ}$, which may indicate a rate of increase of $1^{\circ}$ in 47 feet if the water is from the bottom. This well penetrates sandstones of Cambrian age which lie on granite at a depth of about 500 feet. The beds are nearly horizontal. The 600 -gallon flow from a $707-$ foot well at Henderson, in Sibley County, is stated to have a temperature of $50^{\circ}$, but doubtless part of the water is from beds some distance above the bottom.

## MISSISSIPPI.

## TEMPERATURES.

Temperatures of several flowing wells in Mississippi have been recorded by G. D. Harris ${ }^{1}$ and others, and as most of the wells flow in large volume and appear to be cased to the bottom the temperatures indicate approximate rates of increase. The Biloxi well, with readings at 500 and 900 feet, is of particular interest.

Temperatures in wells in Mississippi.

$a$ Pearlington average for 29 years.
$b$ A verage for 26 years.
c Biloxi average.
d Average for 29 years.
e U. S. Geol. Survey Prof. Paper 46, pp. 226-227, 1906.
$f$ A verage for 46 years.
$g$ Yazoo City average for 23 years.

## SUMMARY.

The rates of increase deduced in Mississippi show considerable variation, but this may be due, in part at least, to lack of information as to source of water, notably in the wells at Canton, Vicksburg, and Sartartia. If, as stated, the Scranton well draws from the bottom, a low rate is indicated.

## GEOLOGIC RELATIONS.

The Bay St. Louis, Biloxi, Pass Christian, and Scranton wells are in the higher Tertiary strata of the southern part of the State. All these strata dip gently to south or west. The Vicksburg, Sartartia, and Canton wells are deep in the Tertiary. The Columbia well is in the Cretaceous.

## MISSOURI.

## TEMPERATURES.

About 50 well temperatures have been obtained for Missouri, some of them taken by Schweitzer ${ }^{1}$ and Shepard, ${ }^{2}$ who were competent observers, and others by various persons who may or may not have obtained reliable figures. The depth from which the flows may

[^23]come is not indicated for some of the wells, and other wells are pumped. The following is a list of the more reliable data:

Temperatures in wells in Missouri.

| Location. | Depth | Flow (gallons per ute). | $\begin{gathered} \text { Tem- } \\ \text { pera- } \\ \text { ture } \\ \left({ }^{\circ} \mathrm{F} .\right) . \end{gathered}$ | $\begin{gathered} \text { Mean } \\ \text { annual } \\ \text { air } \\ \text { tem- } \\ \text { pera- } \\ \text { ture. } \\ \left({ }^{\circ} \mathrm{FF} .\right) . \end{gathered}$ | Teri-perature increase (feet to $1^{\circ}$ ). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |
| Buchanan County: St. Joseph |  |  | 88 | - 52.2 | 60.5 | Mineral water. |
| brewery. | 1,805 |  | 82 |  |  |  |
| Camden County: c |  |  |  |  |  |  |
| Hahatonka. | 864 |  | 63 | 56.6 | 131 | Flow from 840 to 864 feet. |
| Clay County: ${ }^{\text {c }}$ |  |  |  |  |  |  |
| Excelsior Springs, 1 mile north of.$\qquad$ | 1,460 | Many. | 67 | e 54 | 112 | Pumped 150 feet. Source of water not given. |
|  | 1,327 | Many. | a 64.4 | 54 | 127 | Pumped 60 feet. |
| Dunklin County: Campbellf. | 910 |  | 78 | $g 60.5$ | 52 | Flow from 910 feet. |
| Franklin County: Sullivanc. | 1,550 | 8 | 63.5 | h 55.9 | 204 | Six-inch casing to bottom where main flow was found. |
| Greene County: Springfleld.. | 720 |  | 60 | i54.8 | 138 | Temperature at "bottom of well." ${ }^{j}$ |
| Fenry County: Clinton $k \ldots$ Howard County: | 913 | 400 | 64 | $l 56.6$ | 175 | "Cased to bottom." $m$ |
|  | 860 | Many. | 61.5 | - 53.9 | 113 | Saline water from 860 feet. |
| Boonslick. | 1,002 |  | 62 | 53.9 |  |  |
| Jasper County: Joplin | 908 | 16 | 65 | $q 56.2$ | 103 | Cased to bottom; pumped 1192 |
| Lewis County: |  |  |  |  |  |  |
| Canton $r$.. | 900 | 72 | 60 | s 52.5 | 116 | Flow from 870 to 900 feet. |
| La Grange | 850 | 60 | 61 | 52.5 | 94 | Flow from 800 to 850 feet. |
| Do ${ }^{\text {a }}$ | 850 |  | 60 | 52.5 |  |  |
| Livingston County: Utica $k$. | 421 | Few. | 59 | u 52.4 | 64 | Cased 400 fect. |
| Marion County: Hannibal $k$-- | 950 600 | 208 40 | 60 58 | - 53.1 |  | Flow is from considerable depth. No data. |
| Pike County: Louisiana $w . .$. | 1,275 | Many. | $x 64.2$ | $y 53.1$ | 115 | Cased 910 feet. Source of water: |
| Randolph County: Moberly..... | 508 | 100 | 60 | 253.5 | 78 | Pumps; water probably from |
|  |  |  |  |  |  | bottom. |
| Randolph Springs a $\ldots$.... | 969 330 | 120 | y 59 | ${ }_{b s} 53.5$ | ${ }^{176}$ | Flows; source of water not given. |
| Ralls County: Rensselaer station, 5 miles south of. | 330 | 7 | 57 | bb 53.1 | 77 | Flows from 300 to 330 feet. |

a Average for 28 s.
$b$ Oregon average plus $0.5^{\circ}$.
c U. S. Geol. Survey Water-Supply Paper 195, p. 156, 1907. Schweitzer, Paul, Missouri Geol. Survey, vol. 3, p. 99, 1892.
d Warsaw average for 13 years.
e Liberty average for 27 years.
${ }_{f}$ U. S. Geol. Survey Water-Supply Paper 102, p. 392, 1904.
$g$ New Madrid average.
a Oakfield average for 25 years.
$i$ A verage for 29 years.
j U. S. Geol. Survey Water-Supply Paper 195, p. 132, 1907.
$k$ Missouri Geol. Survey, vol. 3, pp. 120-121, 1892.
$l$ Warsaw average.
m U. S. Geol. Survey Water-Supply Paper 195, p. 118, 1907.
$n$ Missouri Geol. Survey, vol. 3, pp. 141-143, 1892.

- Average for 31 years.
$p$ Idem, p. 138.
$q$ Mount Vernon average.
r U. S. Geol. Survey Water-Supply Paper 195, p. 45, 1907.
8 Keokuk average plus $0.1^{\circ}$.
${ }_{t}^{8}$ U.S. Gepol. Survey Water-Supply Paper 195, p. 46, 1907.
$u$ Kidder average.
- Fulton average.
w Missouri Geol. Surver, vol. 3, pp. 94-95, 1892.
$x 63.5^{\circ}$ on July 15, 1903 (U. S. Geol. Survey Water-Supply Paper 195, p. 52, 1907).
y Average for 23 years.
$z$ Brunswick average for 28 years.
aa Missouri Geol. Survey, vol. 3, p. 73, 1892.
bb Hannibal avèrage for 25 years.

Temperature in wells in Missouri-Continued.

| Location. | Depth (feet) | $\begin{aligned} & \text { Flow } \\ & \text { (gal- } \\ & \text { lons } \\ & \text { per } \\ & \text { nin- } \\ & \text { ute). } \end{aligned}$ | $\begin{aligned} & \text { Tem- } \\ & \text { pera- } \\ & \text { ture } \\ & \left({ }^{\circ} \mathrm{F} .\right) . \end{aligned}$ | Mean annual air tem- per:- ture ( ${ }^{\circ} \mathrm{F}$.). | Tem-perature in. crease (feet to $1^{\circ}$ ). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| St. Louis County: <br> St. Lonis: |  |  |  |  |  |  |
| Asylum... | 3,8432 |  | a 105 | ${ }^{\text {b }} 55.8$ | 78.1 | Granite, etc., 3,558 to 3,843 feet. |
| Belcher well c | 2,200 | 75 | 73.4 | ${ }^{1} 55.8$ | c 85.3 | Water from 1,515 feet; much $\mathrm{H}_{2} \mathrm{~S}$; water-bearing sandstone |
| * $\begin{array}{r}\text { Dof.... } \\ \text { Union jowe } \\ \text { North St. L }\end{array}$ |  |  |  |  |  | from 1,502 to 1,640 feet. |
|  | 1,590 | Many. | 69 | 55.8 | 102 | Water from 1,351 feet. |
|  | 1,470 | Many. | 70 62 | 55.8 55 | 103 | Water pumped 70 feet. ${ }^{\text {Water }}$, |
|  |  |  | 62 | 55.8 | 11. | Water pumped 45 feet. Sonrco of water not given. |
| Vernon County:g Nevada, near | $\left\{\begin{array}{r} 800 \\ 869 \\ 1,001 \\ 650 \end{array}\right.$ | 120$\left\{\begin{array}{r}10 n y . \\ \text { Few. }\end{array}\right.$ | 63 | ${ }^{4} 55$ | 100 | Flow from below 785 feet. |
| Nevada... |  |  | 67 | 55 | 72, ${ }_{1}$ | F Flow from below 785 feet; pump- |
| Richards |  |  | 62 | 55 | 93 | Cassy. |

a Broadhead. G. C. (St. Louis Acad. Sci. Trans., vol. 3, p. 216, 1878), who used a registering thermom eter and gave temperatures of $106^{\circ}$ at 3,127 feet; $107^{\circ}$ at 3,129 feet; $106^{\circ}$ at 3,264 feet; $105^{\circ}$ at 3,473 feet; $105^{\circ}$ at 3,533 feet; $105^{\circ}$ at 3,604 feet; $104 \frac{1}{2}^{\circ}$ at 3,641 feet; $105 \frac{1}{2}^{\circ}$ at 3,728 feet; and $105^{\circ}$ at 3,800 feet, observed with thermometer reading to one-fifth of $1^{\circ}$.
$b$ Average for 44 years.
c Litten, A., Belcher \& Bros. artesian well (St. Louis Acad. Sci. Trans., vol. 1, pp. 80-86, 1856-1860).
Another authority gives 2,199 feet (U. S. Geol. Survey Water-Supply Paper 195, p. 159, 1907).
d Litten gives $55.22^{\circ}$ as deduced from observations by Engleman for 22 years, indicating a gradient of $1^{\circ}$ to 83.3 feet.
$e$ Calculated to a depth of 1,502 feet.
$f$ Am. Jour. Sci., 2 d ser., vol. 15, p. 460, 1853.
$g$ U. S. Geol. Survey Water-Supply P'aper 195, p. 104, 1907.
$h$ Average for 22 years.

## SUMMARY.

The data from the Missouri wells listed above appear to indicate considerable diversity in gradient, but doubtless some of this diversity is due to error, especially as to source of water. The tests by Broadhead and Litten in the deep wells of St. Louis are most valuable, giving mean rates of increase of $1^{\circ}$ in about 78 and 85 feet. The apparent decrease of $2^{\circ}$ from 3,127 to 3,843 feet and some other inconsistent variations are difficult to explain. Some of the shallower wells appear to indicate a diminishing rate toward the surface, but uncertainty as to source of flow and influence of pumping may affect the observations. The Clinton well, which is said to be cased to the bottom, indicates the surprisingly low rate of $1^{\circ}$ in 175 feet, and the Sullivan well shows a rate of $1^{\circ}$ in 153 feet. The rates indicated for the Springfield well, $1^{\circ}$ in 138 feet; the Canton well, $1^{\circ}$ in 132 feet; the Hahatonka well, $1^{\circ}$ in 131 feet; and the Fayette well, $1^{\circ}$ in 113 feet, are valid if, as is reported, all these wells are cased to or nearly to the bottom. The data from Mexico, Excelsior Springs, Hannibal, Louisiana, Montgomery, Moberly, and Randolph Springs may be defective. The relatively high rates of $1^{\circ}$ in 52 feet indicated by the Campbell well and $1^{\circ}$ in 53 feet by the Utica well
are exceptional but almost surely correct. The rate indicated by the St. Joseph brewery well, $1^{\circ}$.in $60 \frac{1}{2}$ feet, is also regarded as reliable, although the source of the water is not given precisely.

GEOLOGIC RELATIONS.
The southeastern third of Missouri consists of a low dome in the center of which the pre-Cambrian granites and other rocks are exposed in the Iron Mountain region. West of this dome the strata dip west, and north of it they slope gently to the north and northwest. At Kansas City and northward the pre-Cambrian rocks are far below the surface. The wells included in the above list penetrate a great variety of strata, all on the north or west slopes of the dome. The 3,843 -foot hole at St. Louis is the only one that reached the granite, but its water supply is derived from sandstones of Ordovician age. The wells at Moberly, Hannibal, Canton, Lagrange, Louisiana, Rensselaer, Gunters, Hahatonka, Fayette, Sullivan, Montgomery, Springfield, Nevada, and Richards all draw from sandstones in the Ordovician or Cambrian. The deep well near St. Joseph entered Mississippian limestones at 1,250 feet and possibly reached the St. Peter sandstone. The deep borings of Excelsior Springs are believed to draw their mineralized waters from the "Jefferson City limestone and base of the Pennsylvanian." ${ }^{1}$ The well at Campbell, in the southeast corner of the State, penetrates a thick deposit of clay of the Lagrange formation (Tertiary) and draws from the sand of the upper part of the Ripley formation (Cretaceous). The well at Utica draws water from the base of the Cherokee shale (Pennsylvanian).
There is no evidence that the variations in rate of increase in temperature are related to the horizons from which the waters are derived, to structural features, or to proximity to the pre-Cambrian rocks, with the possible exception of the high rate at the Campbell well, which is in the Tertiary and Cretaceous. The St. Joseph and Utica wells which show high rates are higher in the Paleozoic succession than the wells farther east and south.

## MONTANA.

## ANACONDA MINE.

In 1899 H . V. Winchell made a series of observations in the Anaconda mine, near Butte. He placed Darton thermometers in drill holes at various depths and left them for several days. The following results were obtained:

Temperature at different depths in the Anaconda mine, near Butte, Mont.

a Based on mean annual air temperature of $42^{\circ}$, average for 21 years at Butte.
The reading at a depth of 800 feet is abnormally high, probably on account of some local condition. The rates of increase indicated between various depths are as follows: 200 to 1,600 feet, $1^{\circ}$ in 49.8 feet; 400 to 1,000 feet, $1^{\circ}$ in 48.2 feet; 400 to 1,600 feet, $1^{\circ}$ in 48.5 feet; 600 to 1,600 feet, $1^{\circ}$ in 43.6 feet; 1,000 to 1,600 feet, $1^{\circ}$ in 48.8 feet. S. F. Emmons made an additional observation in the Gagnon mine 1,550 feet below the surface. The Darton thermometer was placed in a 3 -foot drill hole and read $79^{\circ}$ after 48 hours and $79 \frac{1}{4}^{\circ}$ after 7 days. As the mean annual temperature here is $42^{\circ}$ the rate indicated is $1^{\circ}$ in 41.7 feet. This mine is the nearest one to the latest rhyolite eruption.

## MILES CITY.

The flow from a 456 -foot well in Miles City, Custer County, is reported ${ }^{1}$ to have a temperature of $57^{\circ}$. Its volume is 5 gallons a minute and its pressure 7 pounds. Flows were also found at 300 and 390 feet, but presumably they are cased off. The mean annual air temperature at this place is $44.1^{\circ}$ (average for 39 years), or $12.9^{\circ}$ less than the temperature of the flow, a difference indicating a rate of increase of $1^{\circ}$ in 35.3 feet.

## LITTLE BITTERROOT VALLEY.

Meinzer ${ }^{2}$ has given some data on temperature in artesian wells in Little Bitterroot Valley.

[^24]Temperatures in artesian wells in Little Bitterroot Valley, T. 22 N., R. 23 W., Mont.

| Location. | Depth (feet). | $\begin{gathered} \text { Flow } \\ \text { (gallons } \\ \text { per } \\ \text { minute). } \end{gathered}$ | $\begin{gathered} \text { Tempera- } \\ \text { ture } \\ \left({ }^{\circ} \mathrm{F} .\right) . \end{gathered}$ | Depth per degree of increase in temperature (feet). $a$ |
| :---: | :---: | :---: | :---: | :---: |
| NW. $\frac{1}{4}$ sec. 18. | 267 | 14 | 91 |  |
| SE. $\frac{1}{4} \mathrm{sec} .20$. | 232 | 365 | $71 \frac{1}{2}$ | 8.3 |
| SE. $\frac{1}{4} \mathrm{sec} .20$. | 232 | 400 | 78 | 6.7 |
| NW. $\frac{1}{4}$ sec. 29. | 244 | 245 | ${ }^{\circ} 120$ | 3.2 |
| SW. ${ }^{1}$ Sec. 29. | 260 | 85 | 99 | 9.7 |
| SW. $\frac{1}{4}$ sec. 29. | 274 | 25 | 921 | 5.6 |

a Calculated from average of mean annual air temperature at Plains and Kalispell, $43.4^{\circ}$ plus $0.2^{\circ}$.
$b$ Reported to have been $126^{\circ}$ three years before; the volume of flow has also diminished.
The temperatures of flow from wells 52,82 , and 90 feet deep were $57 \frac{1}{2}^{\circ}, 53^{\circ}$, and $59 \frac{1}{2}^{\circ}$, respectively, and flows from two of the deeper wells had low temperatures. Water in a 94 -foot dug well at Lone Pine post office had a temperature of $52 \frac{1}{2}^{\circ}$. The water occurs in a bed of sand and gravel underlying a thick mass of silt supposed to have been deposited by a lake. Meinzer suggests that the variations in temperature in these wells, which are all near together, probably indicates that the warmer water comes in part from some deeper source. The Camas Hot Springs are in the northwest cormer of sec. 3, T. 21 N., R. 24 W .

## NEBRASKA.

## TEMPERATURES.

The well temperatures in Nebraska listed below, except the observations at Lincoln and Omaha, were obtained by G. E. Condra in a reconnaissance of the counties in the northeast comer of the State. ${ }^{1}$ These data are not only very important in themselves but they extend the area of close observation in eastern South Dakota described in subsequent pages.

[^25]Temperatures in wells in Nebraska.

a Yankton, S. Dak., average for 43 years.
o Vermilion, S. Dah., average for 16 years.

Temperatures in wells in Nebraska-Continued.

| Location. | Depth (feet). | Flow (gallons per minute). | Tem-perature ( ${ }^{\circ} \mathrm{F}$.). | Mean <br> air tem-perature ( ${ }^{\circ}$ F.). | Depth per degree of increase in tem-perature (feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Douglas County: Omaha. | 650-800 | Many. | 58-62 | a 50 |  | Many wells; large flows. |
| Do. | 664 | Many. | 58 | 50 | 83 |  |
| Jefferson County: Sec. 4, T. 2 N., R. 2 E. | 500 | Severai. | 70 | b 51.9 | 27.7 |  |
| Knox County: <br> Sec. 18, T. 33, R. 8 W.... | 770 | Several. | 70 | c 48.1 | 33.8 | Third flow at 740 feet. |
| Sec. 29, T. 31, R. 6 W... | 770 | 14 | $74 \frac{1}{2}$ | 48.1 | 29.2 | Third fow at 740 fect. |
| Niobrara, mill. . . | 656 | 1,900 | 68 | 4. 1 | 32.9 |  |
| Niobrara packing house. . | 600 | 280 | 66 | c 48.1 | 33.5 |  |
| Sec. 12, T. 32, R. 6 W | 548 | 240 | 65 | 48.1 | 32.4 | Main flow at 520 feet. |
| Sec. 13, T. 33, R. 5 W. | 740 | 900 | 64 | 48.1 | ${ }^{\text {d }} 37.5$ | Water at 600 feet. |
| Sec. 16. T. 33, R. 3 W.... | 400 | 8 | 62 | 48.1 | 28.8 |  |
| Sec. 15, T. 33, R. 3 W.... | 400 | Few. | 61 | 48.1 | 31.0 |  |
| Sec. 13, T. 33, R. 3 W.... | 504 | 2 | 63 | 48. 1 | 33.9 |  |
| Sec. 19, T. 33, R. 3 W.... | 550 | 10 | 63.3 | 48. 1 | 35.7 |  |
| Sec. 20, T. 33, R. 3 W | 650 | 30 | 62 | 48.1 | 46.8 |  |
| Sec. 17, T. 33, R. 2 W | 400 | $25 \frac{1}{2}$ | 622 | 48.1 | 27.8 |  |
| Sec. 20, T. 33, R. 2 W | 600 | Few | 62 | 48.1 | 43.2 | Flow probably from near 400 feet. |
| Sec. 16, T. 33, R. 2 W.... | 400 | 11 | 62 | 48.1 | 28.8 |  |
| Sec. 15, T. 33, R. 2 W.... | 403 | 14 | 591 | 48.1 | 35.4 |  |
| Sec. 15, T. 33, R. 2 W... | 425 | 14 | 61. | 48.1 | +32.9 |  |
| Lancaster County: Lincoln Salt Lake. | 2,463 | Many. | 60.7 | e 50.1 | $f 57$ | Main flow from 600 feet. Saline water. Smaller flow from 828 feet. Test by C. A. Fisher, 1898. |
| a Average for 46 years. <br> $b$ Fairbury average for 21 years. <br> c Santee Agency average. |  |  |  |  |  |  |
|  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |

## SUMMARY.

The rates of increase indicated at Lincoln and Omaha are moderate and apparently are fairly close approximations. The many records of flowing wells in the lowlands along Missouri River in Boyd, Cedar, Dixon, and Knox counties all show high rates and in general are closely accordant. From these records lines in Plate I, showing relations to the South Dakota area, have been drawn. Most of the records show rates of increase of $1^{\circ}$ in 22 to 30 feet, and there is a notably warm spot in the area southwest of Vermilion. Possibly some higher temperatures in some of the shallower wells are due to waters rising from lower beds, and it may be that all the high temperatures are due to this cause, if this area is a zone of permeable strata in which such a rise would be possible.

## GEOLOGIC RELATIONS.

The many artesian wells in Cedar, Knox, Dixon, and Boyd counties draw their water from the Dakota sandstone, which underlies shales of Benton age and lies on Paleozoic limestone and possibly in part also in pre-Cambrian rocks. The beds are nearly horizontal
and apparently have no local disturbances or notable stratigraphic differences, so that no suggestions can be offered to account for the variations of gradient in relation to the geology. The Omaha wells penetrate deeply into Paleozoic limestone that lie nearly horizontal. The deep boring at Lincoln passed through very slightly tilted Paleozoic limestones and sandstones and reached quartzite that is supposed to be of pre-Cambrian age.

## NEVADA.

COMSTOCK LODE.
The high temperatures in the deep workings of the Comstock lode near Virginia City are well known through the extended investigations of G. F. Becker. ${ }^{1}$ The temperature at a depth of 3,100 feet was found to be $170^{\circ}$, indicating a rate of increase of $1^{\circ}$ in 28 feet, and an elaborate series of temperature determinations in drill holes at various depths gave an average rate of $1^{\circ}$ in 33 feet. The rate decreases horizontally away from the lode. It is believed that the high rate in the Comstock lode indicates the presence of hot volcanic material a few miles underground, and that ascending water is the mportant factor in the transportation of the heat. As a large number of data are presented, the reader interested in details should consult the monograph cited, also a critical review by Locke. ${ }^{2}$

## TONOPAH.

Leon Dominian made measurements of underground temperatures in the Mizpah Extension, Ohio Tonopah, and Montana Tonopah shafts at Tonopah. ${ }^{3}$ These shafts had very limited side workings and no complete ventilation. The thermometers were placed in shallow drill holes. The following results were obtained:

Temperatures ( ${ }^{\circ} \mathrm{F}$.) in mines at Tonopal, Nev.

| Depth (feet). | Mizpah extension. | Ohio Tonopah. | $\begin{gathered} \text { Montana } \\ \text { Tono- } \\ \text { pah. } \end{gathered}$ | Depth (feet). | Mizpah extension. | Ohio Tonopah. | Montana 'I'onopah. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 100. | 601 | 60 |  | 500. | 69 | 664 |  |
| 200. | 613 | 61 |  | 600. | 701 | 69 | 702 |
| 300. | 64 | 62. |  | 700. | 72 | 74 |  |
| 317. |  |  | 64 | 766 |  | 78 |  |
| $\begin{aligned} & 400 . \\ & 460 . \end{aligned}$ | $66 \frac{1}{2}$ | 64 | 68 | 780. | 731 |  |  |

The rate of increase from 100 to 780 feet in the Mizpah extension is $1^{\circ}$ in 51.3 feet; from 100 to 760 feet in the Ohio Tonopah, $1^{\circ}$ in 37 feet; from 317 to 600 feet in the Montana Tonopah $1^{\circ}$ in 43.5 feet.

[^26]From 400 to 780 feet in the Mizpah it is $1^{\circ}$ in 54.3 feet, and from 400 to 766 feet in the Ohio it is $1^{\circ}$ in 26.1 feet. Spurr in comparing the Tonopah and Comstock data points out the fact that the temperature at 766 feet in the Ohio Tonopah was found in the Comstock at 900 feet, and he suggests that the increased rate toward the bottom may be due to proximity to a "local heat focus, such as a hot spring." The $l_{\text {ack }}$ of a record of mean annual air temperature at Tonopah makes it difficult to compute the gradient. The average given by the Weather Bureau for 1916 is $48.9^{\circ}$. The temperature of $60^{\circ}$ at a depth of 100 feet therefore shows a rapid increase in the upper part of the mine.

## NEW JERSEY.

## TEMPERATURES.

Records have been made of temperatures of flows of some of the many artesian wells along the seashore in southern New Jersey. The more useful ones are given in the following table:

Temperature in wells on coast of southern New Jersey.

| Location. | Total depth (feet). | Flow (gallons per minute). | Tem-perature ( ${ }^{\circ} \mathrm{F}$.). | Mean annual <br> air tem-perature ( ${ }^{\circ} \mathrm{F}$.). | Depth per degree of <br> increase <br> ture <br> (feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Asbury Park $a$. | 448 | 20 | 60 | ${ }^{6} 51.9$ | 55 |  |
| Atlantic City ${ }^{\text {a }}$ | 813 | 105 | 66 | c 52.4 | 59 |  |
| Longport.... | 803 | 180 | 66 | c 52.4 | 59 |  |
| Ocean Grove. | 420 | 40 | 60 | d 51.9 | 52 |  |
| Seven Islands. | 408 |  | 60 | 52.1 | 51 | Water also at 570 and 694 feet; probably cased off. |
| Sea Girt. | 755 | 50 | 65 | e 52 | 58 |  |
| Wildwood $f$. | 655 | 300 | 63 | g 53.6 | 69 |  |
| Dof.... | 1,244 | 10 | 67 | $g 53.6$ | $h 88$ | Small flows at 625, 750, and 843 feet and salt water at 1,185 feet; probably some mixing of flows. |

a New Jersey Geol. Survey Ann. Rept. for 1895, pp. 72-74, 1896.
$b$ Average for 27 years.
c Average for 43 years.
d Asbury Park average.
e Asbury Park average plus $0.1^{\circ}$.
$f$ New Jersey Geol. Survey Ann. Rept. for 1894, pp. 155-170, 1895.
$g$ Cape May 32 -year average minus $0.1^{\circ}$.
$h$ Calculated to 1,185 feet.
All but the deep well at Wildwood are believed to get their flow from the bottom, so that they give a reliable indication of the rate of increase if the temperatures are correct.

Water from a 615 -foot pumped well at Newark, yielding 50 gallons a minute, is reported to have a temperature of $55 \frac{1}{2}^{\circ},{ }^{1}$ but the temperature may be modified by pumping. Mudge ${ }^{2}$ mentions a 394 -foot well at New Brunswick in which the rate of increase was found to be $1^{\circ}$ in 72 feet.

[^27]
## SUMMARY.

The wells on the coast of New Jersey from Asbury Park to Wildwood show rates of increase that seem closely accordant when the possible plus and minus is considered. In general there is an increase to the south from $1^{\circ}$ in 55.3 and 52 feet at Asbury Park and Ocean Grove to $1^{\circ}$ in 59 feet at Atlantic City and Longport and $1^{\circ}$ in 69 feet in the 655 -foot well at Wildwood.

GEOLOGIC RELATIONS.
All these wells draw their water from sands in the Coastal Plain sedimentary series, ${ }^{1}$. which dips gently to the east and lies on a platform of pre-Cambrian crystalline rocks. The Asbury Park, Ocean City, and Sea Girt wells draw from sands in the Miocene. It would appear from these facts that there may be a lower rate in the higher beds, but the data are not sufficiently numerous or reliable to prove it.

## NEW MEXICO.

Temperature determinations were made by W. T. Lee in September, 1905, in the well at Sandia siding, on the Atchison, Topeka \& Santa Fe Railway, $11 \frac{1}{2}$ miles southwest of Isleta. A deep-well thermometer was used, with precautions to obtain accurate data. The water stood at a depth of 445 feet in the boring, which was 843 feet deep and penetrated Cretaceous shale and sandstone, overlain by sand of the Tertiary Santa Fe formation at the edge of the sheet of recent lava which covers the plateau to the north and east. The following. results were obtained:

Temperatures at different depths in well at Sandia siding, N. Mex.


843 feet.............................. 78.1 543 feet....................................... 72.1
743 feet............................. 76.7 . 443 feet................................. 71.7
643 feet
75.6

If $54.5^{\circ}$, the mean annual air temperature at Los Lunas (average for 27 years), which is at.about the same latitude, although 400 feet lower, is taken as a standard, the rates of increase indicated by these results are as follows:

Depth per degree of increase in temperature in well at Sandia siding, N. Mex. Feet. Feet.

|  | Feet. |  | Fee |
| :---: | :---: | :---: | :---: |
| Surface to 843 feet | $35 \frac{1}{2}$ | Surface to 443 |  |
| Surface to 743 feet | 35 | 443 to 843 feet. | $62 \frac{1}{2}$ |
| Surface to 643 feet | $30 \frac{1}{2}$ | 643 to 843 feet. | 80 |
| Surface to 543 feet | 31 | 443 to 643 feet. | 51 |

The variability, especially in the amount of change from depth to depth as taken, may be due to convection, but the rate at the bottom is doubtless near the true one.

[^28]
## NORTH CAROLINA.

No significant well temperatures have been reported from North Carolina. The water flowing from the 985 -foot well at Wilmington has a temperature of $70^{\circ}$, or about $7^{\circ}$ higher than the mean annual air temperature, but the flows mingle from depths of $379,496,578$, 608, 734, and 985 feet.

## NORTH DAKOTA.

Nettleton ${ }^{1}$ has given temperatures of flows of several wells in North Dakota. The determinations were probably made with care, and as the flows come from the bottoms of the wells and most of them are of large volume the data give a close approximation of the rate of increase.

Temperatures in deep wells in North Dakota.


These borings are in Cretaceous shale and obtain water from the Dakota sandstone, which lies on or not far above the pre-Cambrian basement. The data from the Hamilton well are doubtless inaccurate as to the depth from which the water comes, or there is a chilling of the small flow near the surface.

## OHIO.

## FINDLAY.

The most important data on underground temperature in Ohio are the measurements in the 2,980 -foot hole at Findlay by John Johnson. ${ }^{2}$ The hole, which was a test for oil or gas, began in the Niagara limestone and continued through older Silurian, Ordovician, and Cam-

[^29]brian shale, limestone, and sandstone to pre-Cambrian granite, which was penetrated 210 feet. ${ }^{1}$ The strata have a very low dip. Johnson used six thermometers, accurately calibrated and sunk in groups of three in a special cage to prevent jarring. The principal results were as follows (averages):

## Temperatures at different depths in hole at Findlay, Ohio.

|  | ${ }^{\circ} \mathrm{F}$. |  | ${ }^{\circ} \mathrm{F}$. |
| :---: | :---: | :---: | :---: |
| 550 feet. | 55.6 | 2,650 feet. | 77.9 |
| 770 feet | 58.1 | 2,800 feet. | 79.7 |
| 1,165 feet | 67.0 | 3,000 feet. | 82.1 |
| 1,900 feet | 72.5 |  |  |

Some gas coming from beds at a depth of 770 feet probably vitiated the readings in that vicinity. Readings after 18 hours continuously at 100 feet gave $51.8^{\circ}$, which is $1.4^{\circ}$ higher than the mean average air temperature from 27 years' observations by the United States Weather Bareau. The figures given by Johnson indicate rates of increase as follows:

Depth per degree of increase in temperature in hole at Findlay, Ohio.

|  | Feet. |  | Feet. |
| :---: | :---: | :---: | :---: |
| 100 to 3,000 feet | 95 | 1,165 to 3,000 feet | 12 L |
| 550 to 3,000 feet: | 92 | 1,900 to 3,000 feet | $114 \frac{1}{2}$ |

## MISCELLANEOUS WELLS.

Several other wells in Ohio give the following data as to rate of increase:

Temperatures in wells in Ohio.

| Location. | Depth (feet). | $\begin{gathered} \text { Tempera- } \\ \text { ture } \\ \left({ }^{\circ} \mathrm{F} .\right) . \end{gathered}$ | Mean annual air temperature ( ${ }^{\circ} \mathrm{F}$.). | Depth per degree ofincrease in temperature (feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Cuyahoga County: Cleveland.. Franklin County: | 535 | 54 | a 49.3 | 114 |  |
| Columbus................... | b 2, 775 ${ }^{2}$ | 91 | c 52.1 | 71 |  |
| Do.d. | 2,575 | 88 | 52.1 | $71 \frac{1}{2}$ | With Walferdin thermometer down 25 hours. |
| Hamilton County: |  |  |  |  |  |
| Cincinnati | 1,360 | 60 | 54.3 | 87 |  |
| Do | 1,960 | 63 | 54.3 | 8 | No flow. |
| Do | 2,408 | 62 | 54.3 |  |  |
| Richland County: |  |  |  |  |  |
| Plymouth 9............... | 3,020 2,800 | 92.8 90.6 | f 49.4 | ${ }_{68} 69$ | Cased to 916 feet. |
|  | 1,400 | 71.2 | ........... | 64 | Temperature taken by Dr. R. D. Sykes. |

[^30][^31]The Cleveland well is pumped and the water in whole or part may come from higher beds. Temperatures were taken in the bottom of the old Columbus well and at $3,020,2,800$, and 1,400 feet in the Plymouth well. In the Plymouth well the rate of increase from 1,400 to 2,800 feet is $1^{\circ}$ in 76.2 feet, and from 1,400 to 3,020 feet $1^{\circ}$ in 70.5 feet. The strata in all these holes are limestones, sandstones, and shales dipping ar low angles. The Columbus and Plymouth wells penetrated to a horizon low in the Ordovician. The Cleveland hole was sunk for water in sandstone of Devonian age.

## OKLAHOMA.

## TEMPERATURES.

A series of careful observations in wells and borings at several points in Oklahoma were made in 1906 by E. G. Woodruff, who used maximum thermometers lent by the Weather Bureau and lowered them to the depths given in the table. The following are the results:

Temperatures in deep wells in Oklahoma.

| Location. | Depth (feet). | Tempera- ture ( ${ }^{\circ} \mathrm{F}$.). | Mean annual air temperature ( ${ }^{\circ} \mathrm{F}$.). | Depth per degree orincrease in tem- perature (feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Cleveland Countr: <br> Norman | 407 | 65.0 | 60.5 | 901 | Old well, plugged. |
|  | 240 | 62.5 | 60.5 | 120 | Pumping well. |
| Kay County: Newkirk, 9 miles southeast of. | 760 | 76.0 | 58.7 | 44 | Gas well, 970 feet deep. Salt water. |
| Lincoln County: Fallis........ | 525 | 63.0 | a 60.4 | 202 | Old hole, full of water. |
| Logan County: Meridian | 2,200 | 82.0 | b 60.3 | 1012 | Boring in progress. |
| Muskogee.... | 1,020 | 84.0 | c 60.4 | 43 | Old oil well. |
| Do. | 410 | 67.0 | 60.4 | 62 | Do. |
| Noble County: Perry, 1 mile east of. | 500 | 68.0 | ${ }^{\text {d } 59.0}$ | $55 \frac{1}{2}$ | Old boring. |
| Pawnee County: |  |  |  |  |  |
| Cleveland. | 1,650 | 107 | - 59.2 | $34 \frac{1}{2}$ |  |
| Ralston, 1 mile wes | 1,625 <br> 2,350 | 107 | 59.2 $f 59.2$ | 34 34 | Well shot the day before. In progress. |
| Payne County: Yale. | 2,500 | 124 | 59.1 | 38.5 | Well abandoned. |
| Pittsburg County: McAlester, 2 miles northwest of. | 540 | 73 | $g 63$ | 54 | In coal, diamond drill, stopped 17 hours before test. |
| Tulsa County: <br> Red Fork, 2 miles northeast of, sec. 14. | 705 | 78 | $h^{6} 60.3$ | 40 | Abandoned hole full of salt water; some gas. |
| Red Fork, sec. 15.......... | 1,170 | 98 | 60.3 | 31 | In shale. Drilling suspended only for test. |
| Wagoner County: Wagoner.... | 1,010 | 78 | <60.3 | 57 | Small salty flow. Temperature |
| Washington County: |  |  |  |  |  |
| Bartlesville. | 1,100 | 86 | ${ }^{j} 59.2$ | 41 | Oil well; some gas. |
| Owen, 2 miles northeast of. | 1,275 |  |  |  | Oil well; tubing just with drawn. |
| Ramona, 5 miles west of... | 1,700 | 103 | 59.4 | 39 | Do. |
| a Guthrie average for 24 years plus $0.1^{\circ}$. <br> b Guthrie average. <br> c A verage for 18 years. <br> a Average for 10 years. <br> $e$ Average for 14 years. <br> $f$ Stillwater average for 24 years. <br> o A verage for 12 years. <br> h Tulsa average for 13 year <br> Average for 20 years. <br> ; Pawhuska average plus <br> ${ }^{2}$ Pawhuska average minu <br> l Pawhuska average plus |  |  |  |  |  |
|  |  |  |  |  |  |
|  |  |  |  |  |  |
|  |  |  |  |  |  |
|  |  |  |  |  |  |

## SUMMARY.

All the borings tested are in the northeastern part of the State except the one at McAlester, which is considerably south of the others. The results show a considerable range in the geothermal gradient, but exclusive of the pumping well at Norman, the old hole at Fallis, and the deep boring at Meridian the rate of increase is $1^{\circ} \mathrm{in}$ 34 to 62 feet. The boring at Meridian was in progress and may have been chilled by the pumping of cool surface water to remove the drillings. All the conditions in the Fallis well are not known, but they appeared to be favorable for a reliable test. For many of the localities there is some uncertainty as to the mean annual air temperature, which had to be approximated from stations at some distance. The holes in Pawnee and Tulsa counties and at Yale and Ramona, which show rates of $1^{\circ}$ in 30 to 40 feet, appear to be consistent in indicating a high rate of increase for at least a portion of the central-northeastern part of the State.

## GEOLOGIC RELATIONS.

The borings all penetrate limestone and shale of Carboniferous age, which dip at a low angle to the west and southwest. There are no notable local disturbances and no igneous rocks, and the pre-Cambrian rocks lie far below the surface.

## PENNSYLVANIA.

## ANTHRACITE REGION.

In 1909 to 1911 I made a series of careful determinations of underground temperatures in the anthracite coal fields. Thermometers were sunk in bore holes and in several places buried for a while in holes in remote chambers in coal mines. The following results were obtained:

Temperatures in coal mines in anthracite region of Pennsylvania.

| Location. | Depth (feet). | $\begin{gathered} \text { Tempera } \\ \text { ture } \\ \text { ( } \left.{ }^{\circ} \mathrm{F} .\right) . \end{gathered}$ | $\begin{gathered} \text { Mean } \\ \text { annual } \\ \text { tempera- } \\ \text { ture } \\ \left({ }^{\circ} \mathrm{F} .\right) . \end{gathered}$ | Depth per degree of increase in temperature (feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Scranton, Pine Brook mine.... | 960 | 59.5 | a 49.3 | 94 | Bore hole 670 feet deep in mine, 290 feet below surface. Water flows. Boring suspended 18 hours. |
| Throop, 1 mile northeast of. | 540 | 53.5 | 49.3 | 129 | Drilling finished 24 hours. |
| Storrs mine, gangway, 1.C Big bed. | 776 | 60.0 | 49.3 | 724 | Old boring 405 feet deep in mine 371 fect below ground. |
| Sloan mine, lowest Dunmore bed. | 580 | 64.5 | 49.3 | 38 | In new chamber in a 6 -foot hole bored in coal the day before. |
| Woodward mine, gangway No. 5 E . | 1,200 | 61.0 | ${ }^{6} 49.8$ | 107 | 6 -foot hole in coal; no ventilation. |
| South Wilkes-Barre, mine, 9 slope, Baltimore bed. | 1,300 | 64.0 | 49.8 | 913 | 6-foot hole in coal; some ventilation. |

Temperature in coal mines ins the anthracite region of Pennsylvania-Continued.

| Location. | Depth (feet). | $\begin{gathered} \text { Tempera- } \\ \text { ture } \\ \left({ }^{\circ} \mathrm{F} .\right) . \end{gathered}$ | $\begin{gathered} \text { Mean } \\ \text { annual } \\ \text { tempera- } \\ \text { ture } \\ \left({ }^{\circ} \mathrm{F} .\right) . \end{gathered}$ | Depth per degree of increase in tem(feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Dorrance mine. | 479 | 63.5 | 49.8 | 35 | 360 -foot hole in old Baltimore workings, 119 feet below ground. |
| Auchincloss mine, foot of No. 3 slope, No. 1 shaft. | 1,283 | 64.2 | a 49.9 | 90 | Old 115-foot hole from Ross bed to Red Ash bed. |
| Lackawanna \& Western Coal Co., southwest of WilkesBarre: |  |  |  |  |  |
| Bore hole 172. | 880 | 56.5 | ${ }^{6} 50$ | $135 \frac{1}{2}$ | Full of water. |
| Bore hole....i....... | 843 1,030 | 56.0 61.5 | 50 50 | 1403 | Do. |
| Glen Lyon drill hole near.. | 1,598 | 61.5 | 50 | 130 | Full of water. Thermometer down on Jan. 12, 1911, for 24 hours. Old hole. |
| Pottsville, Pottsville shaft, 2 miles north of. | 1,596 | 61.8 | c 48.3 | 118 | Shatt full of water. |
| Newkirk, $\frac{1}{2}$ mile north of. ${ }^{\text {a }}$. ${ }^{\text {a }}$ | 1,090 | 59.5 | 48.3 | 90 | Old drill hole in coal measures. |
| Minersville, near, Lehigh Valley Co. | 620 | 51.0 | 48.3 | 109 | Old borehole with flow of water at temperature $51^{\circ} \mathrm{F}$. |
|  | 630 | 54.0 | 48.3 | 110 | Old bore hole containing water. |
| Mount Carmel, east o | 1,125 | 56.3 | 48.3 | 141 | Boring in progress; stopped 12 |
| Mount Carmel. | 1,125 | 56.5 | 48.3 | 137 | Same hole; another thermome- |
| Mount Carmel, flowing well. | 1,125 | ${ }^{\text {d }} 54$ | 48.3 | 197 | ter. Main flow from 1,125feet. Some at 650 feet (end of casing). Air lift to depth of 300 feet but not working at time of test. |
| Mount Carmel, 2 miles east of.. | 1,300 | 55.0 | 48.3 | 194 | Old hole, 2,060 feet deep but blocked at 1300 feet |
| Lofty, $\frac{1}{2}$ mile east of. . | 680 | ${ }^{\text {d }} 51$ | e 47.0 | 170 | 1,100 -foot boring blocked at 680 |
| East Brookside. | 1,850 | 60.5 | $f .49 .3$ | 165 | 3 -foot hole in coal in mine |
| Tower City. | 415 | 53.0 | 49.3 | 113 | Old drill hole, full of water. |

a Scranton average plus $0.60^{\circ}$.
${ }^{b}$ Scranton average plus $0.7^{\circ}$.
c Gordon average for 13 years.
${ }^{d}$ By two thermometers.
a Approximate altitude is 1,500 feet. Mean annual may be much lower than $47^{\circ}$.
$f$ Gordon average plus $0.1^{\circ}$.
An 1,800-foot diamond-drill hole bored horizontally in the coal measures by the Philadelphia \& Reading Coal \& Iron Co. in the Huntet tunnel, near Ashland, in Schuylkill County, found a 150 -gallon flow of water with a temperature of $54^{\circ}$, or about $4^{\circ}$ above the mean annual air temperature. A similar hole on Potts Run near Mahanoy City, in the same county, found a 200 -gallon flow with a temperature of $51^{\circ}$, but the depth below the surface is not stated.

Except those in the holes in coal in the Dorrance and Sloan mines the rates of increase are remarkably low. The reasons for the differences in the others, however, are not apparent. In the holes near Scranton and Wilkes-Barre the strata are in a wide syncline with gentle dips and with a few local crenulations. In the Auchincloss mine the beds are highly tilted and faulted. The test in this mine was made in a 115 -foot hole in one of the lower chambers, through which passes a ventilating current. The Glen Lyon hole and the Lehigh \& Wilkes-Barre Co. bore holes are in steeply tilted strata.

The Pottsville shaft was an abandoned shaft at a mine that had been full of water for several years. This and the remaining observations in the table were made in the southern anthracite basin, where the strata lie in a deep syncline with rather steep dips on the sides.

## WEST ELIZABETH.

A well 5,386 feet deep has been put down on Peters Creek about $2 \frac{1}{2}$ miles west of West Elizabeth and about 12 miles south-southeast of Pittsburgh. The rocks are nearly horizontal and comprise strata from those not far below the Pittsburgh coal bed to those near the base of the Devonian. Temperatures in this well were taken by William Hallock ${ }^{1}$ in 1897, by lowering four thermometers to a depth of 5,000 feet and leaving them suspended for 16 hours. The reading was $120.9^{\circ}$, an increase of $68.2^{\circ}$ over the mean annual air temperature at Pittsburgh for 46 years ( $52.7^{\circ}$ ), which indicates a rate of $1^{\circ}$ to 73.3 feet. A reading at 2,350 feet gave about $78^{\circ}$, indicating a somewhat higher rate, but this observation was not considered satisfactory.

## McDONALD.

Some preliminary information regarding a 6,975 -foot hole 4 miles northwest of McDonald, 15 miles southwest of Pittsburgh, has been given by I. C. White and C. E. Van Orstrand. ${ }^{2}$ Until recently this was the deepest hole in the United States. Precise determinations of temperature have been made by C. E. Van Orstrand, some of which are as follows:

|  | ${ }^{\circ} \mathrm{F}$. |  | F. |
| :---: | :---: | :---: | :---: |
| 100 feet. | 57.0 | 4,600 feet. | 110. 5 |
| 350 feet. | 59.2 | 5,100 feet. | 121. 1 |
| 1,100 feet. | 64.4 | 5,600 feet. | 133. 3 |
| 2,100 feet. | 76.3 | 6,000 feet. | 140.0 |
| 2,600 feet. | 82.4 | 6,600 feet. | 144. 1 |
| 3,100 feet. | 88. 2 | 6,775 feet. | 145.8 |
| 3,600 feet. | 96.8 | 6,975 feet. | 144.9 |
| 4,100 feet. | 104. 4 |  |  |

The rates of increase indicated by some of these figures are as follows:

|  | Feet per degree F. |  | Feet per degree F. |
| :---: | :---: | :---: | :---: |
| 100 to 6,775 feet. | . 75.2 | 3,100. to 6,775 feet. | - 63.8 |
| 100 to 6,600 feet. | 75.8 | 4,100 to 6,775 feet. | 64.6 |
| 100 to 5,600 feet. | . 72.1 | 5,100 to 6,775 feet. | 67.7 |
| 1,100 to 6,775 feet | ..... 69.7 |  |  |

[^32]White ${ }^{1}$ has given five readings in this well accredited to Dr. Hallock:

$$
\text { Temperatures at different depths in hole near Mc } \dot{\text { Donald, }} \text { Pa. }
$$

|  | ${ }^{\circ} \mathrm{F}$. |  | ${ }^{\circ} \mathrm{F}$. |
| :---: | :---: | :---: | :---: |
| 525 feet. | 57 | 5,010 feet. | 120 |
| 2,252 feet | 64 | 5,380 feet. | 127 |

The rate indicated by the temperature at 5,380 feet is $1^{\circ}$ in $72 \frac{1}{2}$ feet, calculated from the mean annual air temperature at Pittsburgh, $52.7^{\circ}$; the difference from 525 to 5,380 feet is $1^{\circ}$ in about 69 feet.
Mr. C. E. Van Orstrand ${ }^{2}$ found that the temperature was $142^{\circ}$ at a depth of 6,100 feet. With a temperature of about $55^{\circ}$ at 100 feet the rate of increase is $1^{\circ}$ in about 70 feet.

## PITTSBURGH.

An 1,826 -foot hole at the American Iron \& Steel Works in the 24th ward in Pittsburgh found a strong flow of water at 1,535 to 1,606 feet with a temperature of $68^{\circ},{ }^{3}$ which indicates a rate of $1^{\circ}$ in about 100 feet. The water was somewhat gassy.

## HOMEWOOD.

In the Dillworth well at Homewood a series of six tests was made with the following results: ${ }^{4}$

Temperatures in Dillworth well, Homewood, Pa.

| Depth (feet) | Temperature ( $\left.{ }^{\circ} \mathrm{F}.\right)$. | Depth per degree of increase in temperature (feet). |
| :---: | :---: | :---: |
| 3, 600 | ${ }_{89}^{96}$ | 83 |
| 3,920 | 102 |  |
| 4,002 | 108 | $72 \frac{1}{2}$ |
| 4, 215 | 111 | 72 |
| 4,295 | 114 | 70 |

The tests were of only five to ten minutes duration, but exclusive of the reading at 3,710 feet, where the temperature was probably . lowered by gas, the rate appears to be reasonable. The increasing rate downward is a peculiar feature.

[^33]
## SOUTH CAROLINA.

## TEMPERATURES.

Among the few records of temperatures of flows from deep wells in South Carolina the following appear to be reliable:

Temperatures in wells in South Carolina.


a South Carolina resources, etc., State Board of Agriculture, p. 674, 1883.
b Average for 46 years.
c Stephenson, L. W., A deep well at Charleston, S. C.: U. S. Geol. Survey Prof. Paper 90, p. 70, 1914. Temperature by I. N. Knapp, engineer.
d Calculated to 1,971 feet.
e Am. Jour. Sci., 2 d ser., vol. 47, p. 357, 1869.
$f$ U. S. Geol. Survey Water-Supply Paper 114, p. 151, 1805.
o U. S. Geol. Survey Bull. 298, p. 265, 1906.
${ }^{h}$ Calculated to 1,820 feet.
© Yemassee average for 23 years.
SUMMARY.
As the Charleston wells are cased to or nearly to the bottom, presumably the others are also. The temperatures are probably reliable, especially those of the Charleston wells. The new well at the electric-light plant is cased to a depth of 1,869 feet, and the flow comes from $\cdot$ a bed extending from 1,971 to 1,999 feet. The results of the Charleston and Fort Moultrie wells are fairly accordant and give an average rate of $1^{\circ}$ in about 60 feet. It is possible that the lower rates indicated by the wells at Greenpond and Johns Island are due to ingress of water above the bottoms of the borings.

## GEOLOGIC RELATIONS.

These wells all penetrate clays, sands, and gravels of the thick wedge of Tertiary and Cretaceous strata constituting the Coastal Plain. These strata dip very gently eastward and lie on an eastward sloping floor of old crystalline rocks. The Charleston wells reach the basal beds of the Black Creek formation, which is low in the Upper Cretaceous; the other holes in the list reach sands of Eocene age.

## SOUTH DAKOTA.

## TEMPERATURES.

In the course of detailed investigations of the geology and water resources of parts of South Dakota I and my associates have observed many underground temperatures, mainly of the flow of artesian wells. Some carefully determined temperatures of representative wells have been recorded by Sheppard, ${ }^{1}$ and several tests were made by Nettleton. ${ }^{2}$ Records of wells at Belle Fourche, Cheyenne Agency, and Edgemont were obtained from correspondents. A large number of data were obtained by J. E. Todd ${ }^{3}$ in field work for geologic folios and reports on the artesian area in Beadle, Brown, Edmunds, Faulk, Hand, and Spink counties, in the eastern part of the State. His observations were made with instruments that had been verified in Washington. No flow was tested unless $\cdot$ it was believed to be of sufficient volume to give a closely approximate indication of the temperature at the bottom.

Temperatures in wells in South Dakota.


[^34]Temperatures in wells in South Dakota-Continued.

a J. E. Todd, observer.
o J. H. Sheppard, observer.
c Average for 20 years.
d E. S. Nettleton, observer.
e Yankton average plus $0.5^{\circ}$.
f A berdeen average for 27 years minus $0.9^{\circ}$.
04 berdeen average minus $0.4^{\circ}$.
$n$ Aberdeen average minus $0.3^{\circ}$.
Aberdeen average plus $0.1^{\circ}$.
$\$$ Mean for 27 years.
${ }_{3}$ Aherdeen average plus $0.1^{\circ}$.
$t$ Doubtful.
$m$ A herdeen average plus $0.2^{\circ}$.
n Approximated

- F. A. Durst, observer.
p U.. S. Geol. Survey Water-Supply Papar 227, p. 77, 1909.

Temperatures in wells in South Dakota-Continued.

| I ocation. |  | $\begin{aligned} & \text { Depth } \\ & \text { of main } \\ & \text { flow } \\ & \text { (feet). } \end{aligned}$ | $\begin{gathered} \text { Flow } \\ \text { (gallons } \\ \text { per } \\ \text { minute). } \end{gathered}$ | Temperature ( ${ }^{\circ} \mathrm{F}$.). | Mean annual temper( ${ }^{\circ} \mathrm{F}$.). | Depth per of increase in temperature (feet). |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Charles Mix County: |  |  |  |  |  |  |
| Greenwood. |  | 651 | 3,000 | 70 | 48.9 | 30.9 |
| Lake Andes |  | 725-773 | 1,500 | 70 | 46 | 30.2 |
| Do |  | 802 | Many. | 70 | 46 | 33.4 |
| Sec. 17, T. 98, R. 66 |  | ${ }_{8} 80$ | Many. | 70 | a 45 | 34.4 |
| Sec. 33, T. 99; R. 66 |  | 810 | 20 | 70 | 45 | 33.6 |
| Custer County: |  |  |  |  |  |  |
| Davison County: |  |  |  |  |  |  |
| Mitchell. |  | 530 | 600 | d 56 | e 45.5 | 50.5 |
| Sec. 3, T. 101, R. 62 |  | 565 | 60 | $f 56$ | $g 45.5$ | 54.8 |
| Sec. 35, T. 104, R. 6 |  | 507 | 40 | 56 | e 45.5 | 48.3 |
| Day County: Andover....... |  | 1,070 | 300 | n 71.6 | a 42 | 39.3 |
|  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |
| Armour, 3 miles south of |  | 703-736 |  | 65 | 46.4 | 31.8 |
| Edmunds County: |  |  |  |  |  |  |
| Ipswich ... |  | 1,265 | Many. | ${ }^{\text {h }} 71.6$ | $k 42$ | 42.7 |
| SE. 7 sec. 22, T. 123, R. 68 |  | 1,195 | 97 | 73 | 42 | 38.6 |
| SE. $\frac{1}{\text { sec. }} 23$, T. 123, R. 68 |  | 1,125 |  | f 70.5 | $m 42.2$ | 39.8 |
| SE. $\frac{1}{2}$ sec. 8, T. 123, R. 67 |  | 1,140 |  | ${ }^{f} 63$ | ${ }^{m} 42.2$ | 54.8 |
| NW. $\frac{1}{2}$ sec. 10, T. $123, \mathrm{R} .67$ |  | 1,185 |  | ${ }^{6} 69$ | $m 42.2$ | 44.2 |
| SW. $\frac{1}{4}$ sec. $10, \mathrm{~T} .123, \mathrm{R} .67$. |  | 1,115 |  | ${ }^{\prime} 69$ | ${ }^{m} 42.2$ | 41.6 |
| SWW d sec. 11, T. 123, R. 67. |  | 1,135 |  | $f 68$ | $m 42.2$ | 44. |
| NW. $\frac{1}{2}$ sec. 20 , T. 123, R. 67. |  | 1,140-1, 160 |  | ${ }_{f} 69.5$ | $m 42.2$ | 41.8 |
| SW. ${ }_{\text {NE }} \frac{1}{2}$ sec. $\mathrm{sec} .23, \mathrm{~T}, \mathrm{~T} .123, \mathrm{~T}, 123, \mathrm{R} .67$ |  | 1,140 1,078 | ....... | $f 67$ $f 67$ | ${ }_{m}^{m} 42.2$ | ${ }_{43.5}^{46 .}$ |
| Center of sec. 8 , T. $123, \mathrm{R} .66$ |  | 1,100 |  | ${ }^{7} 66$ | ${ }_{m} 42.2$ | 46.5 46.2 |
| SE. $\frac{1}{4} \mathrm{sec} .20, \mathrm{~T} .123, \mathrm{R} .66$ |  | 1,050-1, 100 |  | $f 69.5$ | $m 42.2$ | 38.5 |
| NW. $\frac{1}{4}$ sec. $16, \mathrm{~T} .122, \mathrm{R} .6$ |  | 1,140 |  | $f 65.5$ | ${ }^{m} 42.2$ | 48.9 |
| SE. ${ }_{\text {St }}$ sec. $26, \mathrm{~T} .122, \mathrm{R} .67$ |  | 1,000-1,060 |  | ${ }^{\prime} 65$ | ${ }^{m} 42.2$ | 43.7 |
| SE. ${ }^{1}$ sec. $4, T .122, \mathrm{R} .66$. |  | 1,080 |  | ${ }^{\prime} 65.5$ | ${ }^{m} 42.2$ | 46.8 |
| SW. ${ }^{\text {s sec. }} 10, \mathrm{~T} .122, \mathrm{R} .66$. |  | 1,025-1, 070 |  | ${ }^{\prime} 65.5$ | ${ }^{m} 42.2$ | 45. |
| NE. ${ }^{1}$ sec. 19, T. 122, R. 66. |  | 1,038 |  | ${ }_{f} 66.5$ | ${ }^{m} 4.2$ | 44.5 |
| NE. $\frac{1}{4}$ sec. 29, T. 122, R. 66 |  | 1,007-1,045 |  | ${ }^{\prime} 655$ | ${ }^{m} 42.2$ | 43.7 |
|  |  | 1,050 1,060 |  | f 65.5 $f 675$ | ${ }_{74}{ }_{4} 42.2$ | 45.1 |
| SE. $\frac{1}{2}$ sec. $33, \mathrm{~T}$. 122 R R 66 |  | 1,060 |  | ${ }_{f}^{666}$ | ${ }^{n} 42.2$ | 41.9 |
| SE. $\frac{1}{3}$ sec. 8, T. $121, \mathrm{R} .67$. |  | 1,050+ |  | f 64 | ${ }_{n}{ }_{42}$ | 42.9 47.8 |
| NWं. $\frac{1}{} \mathrm{sec}, 32$ T. $121, \mathrm{R} .67$ |  | 1,125 |  | f 66 | ${ }^{1} 42$ | 46.8 |
| SE. $\frac{1}{4}$ sec. $26, \mathrm{~T} .121, \mathrm{R} .67$. |  | 1,155 |  | $f 65$ | $n 42$ | 50.2 |
| SW. $\frac{1}{}$ sec. 10, T. 121, R. 66 |  | 1,000 |  | $f 65$ | ${ }^{\square} 42$ | 43.5 |
| SE $\frac{1}{2}$ sec. $29, \mathrm{~T}, 121, \mathrm{R} .66$. |  | 1,000 |  | ${ }_{f} 64.5$ | ${ }^{n} 42$ | 44.4 |
| NE. $\frac{7}{4}$ sec. 32, T.121, R. 66 |  | 967 |  | $f 64$ | $n 42$ | 44. |
| Fall River County: Edgemont. |  | 2,920-2,965 | 350 | $\bigcirc 122$ | $p 46$ | 39 |
| Faulk County: |  |  |  |  |  |  |
| SE. $\frac{1}{4} \mathrm{sec} .13, \mathrm{~T} .120, \mathrm{R} .68$. |  | 1,125 | 25 | $f 66$ | $q 42.7$ | 48.2 |
| NE. $\frac{1}{4}$ Sec. $19, \mathrm{~T} .120, \mathrm{R} .67$. |  | 1,140 | Many. | $f 68$ | $q 42.7$ |  |
| SE. $\frac{1}{\text { sec. }} 21, \mathrm{~T} .120, \mathrm{R} .67$ |  | $1,040 \pm$ | Many. | $f 67$ | $q 42.7$ | 42.8 |
| SW. $\frac{1}{4}$ sec. 1, T. 120, R. 67. |  | 1, 085-1, 120 | Many. | $f 66$ | $q 42.7$ | 46.5 |
| SE. $\frac{1}{4}$ sec. $26, \mathrm{~T} .120, \mathrm{R} .67$ |  | 1,050 |  | $f 69.5$ | $q 42.7$ | 39.1 |
| NW. $\frac{1}{\text { sec. } 24, \mathrm{~T} .120, \mathrm{R} .67}$ |  | 1,038 |  | $f 67+$ | q42.7 | 42.7 |
| NW. $\frac{1}{4}$ sec. 6, T. $120, \mathrm{R} .66$ |  | 1,066 | 100 | $f 67.5$ | $q 42.7$ | 42.9 |
| SW. $\frac{1}{4}$ sec. 2, T. $120, \mathrm{R} .66$. |  | 1,001 | 65 | $f 66.3$ | $q 42.7$ | 42.4 |
| NE. $\frac{1}{4}$ sec. 26, T. 120, R. 66 |  | 965 | Many. | $f 65.5$ | $q 42.7$ | 42.3 |

${ }_{b}$ Approximated.
${ }^{b}$ N.H. Darton, observer, with deep-well thermometer.
${ }^{c}$ Oelrichs average for 25 years minus $0.2^{\circ}$.
d E. Ss Nettleton, olserver.
$e$ Average for 23 years.
$f$ J. E. Todd, observer.
$o$ Mitchell a verage plus $0.2^{\circ}$.
${ }^{h}$ J. H. Sheppard, ohserver.
8 Pierre average for 26 years minus $1.7^{\circ}$.
$j$ A verage for 22 years.
${ }_{l}^{\boldsymbol{k}} \boldsymbol{l}$ Bowdile average for 14 years minus $0.2^{\circ}$.
$t$ Flow from 1, 195 feet.
$m$ Bowdle average.
$n$ Bowdle average minus $0.2^{\circ}$
$o$ Taken by engineer with Darton thermometer; varied from $121^{\circ}$ to $122^{\circ}$.
$p$ Oelrichs average, minus $0.2^{\circ}$.
${ }_{q}$ Faulkton average for 22 years minus $0.2^{\circ}$.

Temperatures in wells in South Dakota-Continued.

| Location. | $\begin{gathered} \text { Depth } \\ \text { of main } \\ \text { flow } \\ \text { (feet). } \end{gathered}$ | $\begin{array}{\|c\|} \text { Flow } \\ \text { (gallons } \\ \text { per } \\ \text { mirute). } \end{array}$ | Temper ature ( ${ }^{\circ} \mathrm{F}$.) | Mean annual ature ( ${ }^{\circ} \mathrm{F}$.). | Depth per degree crease in temperature (feet). |
| :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |  |
| SW. $\frac{1}{1}$ sec. $1, \mathrm{~T} .119, \mathrm{R} .67$ | 980 | Many. |  | 642.8 | 37.4 |
|  | 1,050 1,060 | Many. | a $a$ 698 | 642.8 842.8 8 | 39.3 42.1 |
| SW. $\frac{1}{\text { sec. 18, T. 119, }}$ R. 67. | 1,257 | Many. | a 71 | b 42.8 | 44.5 |
| SW. it sec. 17, T. 119, R. 67 | 1,120 | Many. | a 69 | b 42.8 | 42.7 |
|  | 1,050 | Many. | a 65 | ${ }^{6} 42.8$ | 47.3 |
| SW. $\frac{1}{2}$ sce. 24, T. 119, R. 67 | 1,162 | Many. | a $69+$ | $\bigcirc 42.8$ | 44.3 |
| SE. ${ }^{\text {s sec. } 25, \mathrm{~T} .119, \mathrm{R} .67}$ | 1,031 | 20 | a 69 | ${ }^{6} 42.8$ | 39.3 |
| SW. ${ }^{\text {d sec. } 25, \mathrm{~T}}$. 119. R. 67 | 1,035 |  | ${ }^{\text {a }} 68$ | ${ }^{6} 42.8$ | 41 |
| NW. ${ }^{\text {a sec. } 30, \mathrm{~T} .119 \mathrm{R} .67}$ | 1,000 |  | a 70 | b 42.8 | 36.7 |
| NE. $\frac{1}{4}$ sec. 31, T. 119, R. 67 | 910 |  | a 70 | ${ }^{6} 42.8$ | 33.4 |
|  | 950 |  | ${ }^{6} 67$ | ${ }^{6} 42.8$ | 39.2 |
| SE. $\frac{1}{}$ scc. 9, T. 119, R. 66 | 988 | $50 \pm$ | $\stackrel{\text { a }}{ } 6$ | ${ }^{6} 42.8$ | 42.6 |
| NW. $\frac{1}{4}$ sec. 15, T. 119, R. 66 | 1,010 | 50 | a 66 | ${ }^{6} 42.8$ | 43.5 |
| NW. $\frac{1}{8}$ sec. 20, T. 119, R. 66 | 1,002 | 50 | $a 67$ | ${ }^{6} 42.8$ | 41.4 |
| SE. ${ }^{\text {d sec. } 23, \mathrm{~T} .119, ~ R . ~} 66$ | 1,020 | 50 | ${ }^{6} 64$ | ${ }^{6} 42$ | 48.1 |
| Faulkton. | 1,032土 | 100 | c 74.5 | ${ }^{\text {d }} 42.9$ | 32.6 |
| SE. $\frac{1}{\text { sec. } 24, ~ T . ~ 118, ~ R . ~} 68$ | 1,050+ | 80 | a 69 | ${ }^{\text {d }} 42.9$ | 40.2 |
| SE. $\frac{1}{4}$ sec. 18, T. 118, R. 67 | 1,067 | 75 | a 71 | d 42.9 |  |
| NW. $\frac{1}{4}$ sec. $20, \mathrm{~T} .118, \mathrm{R} .67$ | 1,020 |  | ${ }^{\text {a } 66.5}$ | d 42.9 | 42.2 |
| NW. $\frac{1}{4}$ sec. 28, T. 118, R. 67 | 1,029 | 75 | ${ }^{\text {a }} 69$ | ${ }^{\text {d }} 42.9$ | 39.4 |
| West center of sec. 22, T. 118, | 1,010 | 70 | ${ }^{\text {a }} 69$ | ${ }^{\text {d }} 42.9$ | 38.7 |
| SW. d sec. 13, T. 118, R. 67. | 1,002 | 60 | a 69 | d 42.9 | 38.4 |
| SW. ${ }^{\text {d }}$ sec. 34, T. 118, R. 67 | . 980 | Many. | a 67.8 | ${ }^{\text {d }} 42.9$ | 39.3 |
| NW. $\frac{1}{4}$ sec. 15, T. 118, R. 66 | 1,029 | Many. | ${ }^{\text {a } 68.5}$ | d 42.9 | 40.1 |
| NW. $\frac{1}{1}$ sec. 25, T. 118, R. 66 | 1972 | Many. | ${ }^{a} 678.3$ | d 42.9 | 39.8 |
| SW. $\frac{1}{4}$ sec. 31, T. 117, T. 68 | 1,095 | 90 | ${ }^{\text {a } 68.8}$ | C 43 | 42.4 |
| Miranda. | 1,095 | Many. | $a 68.3$ | e 43 | 43.3 |
| Center of sec. 31, T. 117, R. | 1,010 | Many. | ${ }^{\text {a } 69.6}$ | e 43 | 38 |
| SW. $\frac{1}{4}$ sec. 6, T. 117 R. 67. | 1,095 | Many. | ${ }^{\text {a } 68.8}$ | -43 | 42.4 |
| NE. $\frac{1}{\text { sec. }}$ 31, T. 117, R. 66 | 1,030 | Many. | ${ }^{\boldsymbol{a}} 69.6$ | -43 | 38.7 |
| Rockham.... | 1,010 |  | ${ }^{a} 69.6$ | e 43 |  |
| SE. ${ }_{\text {Tin }}$ Sec. $11, \mathrm{~T}$ |  | 45 | ${ }^{a} 66.7$ | ${ }^{-43}$ | 38.5 |
| Orient...... | 1, 215 | 90 | a 75 | C43 | 48 |
| Gregory County: Fort Randall.................$~$Hand County: |  |  |  |  |  |
|  |  |  |  |  |  |
| Miller...... | 1,115-1,139 | 360 | c79. 8-f78 | 944.0 | 31.2 |
| Miller, near | 1,105 |  | a 76 | $g 44.0$ | 34.6 |
| Do. | 1,100 |  | a 78.9 | $g 44.0$ | 31.5 |
| Do. | 1,100 | 200 | a 78.5 | g 44.0 | 31.9 |
| NE. $\frac{1}{\text { sec. } 10, \mathrm{~T} .116, \mathrm{R} .66}$ | ${ }^{1} 151$ |  | a 73 | h 43.1 | 31.8 |
| NE. $\frac{1}{4}$ sec. 7, T. 115, R. 68. | 1,150 | Many. | ${ }^{\text {a }} 74$ | ${ }^{\text {h }} 43.1$ | 37.2 |
| SE. $\frac{1}{4}$ sec. $22, \mathrm{~T} .115, \mathrm{R} .68$ | 1,065 | Many. | a 74.5 | h 43.1 | 33.9 |
| SE. sec. 12, T. 115, R. 67 | 955 |  | a 71 | ${ }^{\text {h }} 43.1$ | 34.2 |
| NE. $\frac{1}{\text { sec. 22, T. 114, R. } 66}$ | 880 | 42 | ${ }^{\text {a } 68.6}$ | 43.3 | 34.4 |
| NE. $\frac{1}{}$ sec. 12, T. 113, R. 68 | 1,040 | Many. | a 78.9 | j 43.5 | 29.4 |
| SE. $\frac{1}{6}$ sec. 3, T. 113, R. 67 | 1,008 | 60 | a 71.5 | ${ }^{3} 43.8$ | 36.4 |
| SE. $\frac{1}{2}$ sec. 18, T. $113, \mathrm{R} .66$ | 1,105 | 40 | a 76.5 | j 43.8 | 33.8 |
| SE. $\frac{1}{4}$ sec. $25, \mathrm{~T}, 113, \mathrm{R} .67$ | 1,100-1, 129 | 400 | ${ }^{\text {a } 78.5}$ | ${ }^{j} 4338$ | 31.7 |
| $\mathrm{NW} .1{ }^{\frac{1}{2} \text { sec. } 22 .}$, 'T. 113, R. 66 |  |  | a 73.5 | f 43.8 | 32.6 |
|  | 1,099-1,133 |  | ${ }^{\text {a } 78.5}$ | ${ }^{4} 43.8$ | 31.7 |
|  | 1,100 | 110 | a 78.9 | ${ }^{k} 43.8$ | 31.6 |
|  | 1,165 | 14 | ${ }^{\circ} 78$ | ${ }^{h} 44$ | ${ }^{36.3}$ |
|  | 1,165 .892 | 50 | a 76.6 | $h 44.0$ | 33.1 |
|  |  |  |  |  |  |
| Pierre. | $\left\{{ }^{11,215-}\right.$ | 200 | a 89 | m 45.5 | 27.9 |
| East Pier | 1,150- | 900 | c 91.8 | 45.5 | 24.8 |
|  | 1,170 |  |  |  |  |
| Harold. | 1,451 | \} 84 | c 94.9 | 45 | 28.7 |

a J. E. Todd, observer.
b Faulkton average for 22 years minus $0.1^{\circ}$.
c J. H. Sheppard, observer.
d Average for 22 years.

- Faulkton average plus $0.1^{\circ}$.
$f$ E. S. Nettleton, observer.
- Highmore average for 26 years minus $0.6^{\circ}$.
h Howell average for 15 years.
- Howell average plus $0.2^{\circ}$.
$j$ Howell average plus $0.4^{\circ}$.
$t$ Howell average plus $0.7^{\circ}$.
1 Other flows 1,140 and 1,185 feet; much gas.
$m$ A verage for 26 years.

Temperatures in wells in South Dakota-Continued.


[^35]Temperatures in wells in South Dakota-Continued.

| Location. | Depth of main flow (feet). | $\left\|\begin{array}{c\|} \text { Flow } \\ \text { (gallons } \\ \text { per } \\ \text { minute). } \end{array}\right\|$ | Temperature ( ${ }^{\circ} \mathrm{F}$.). | Mean annual temperature ( ${ }^{\circ} \mathrm{F}$.). | $\underset{\text { Depth }}{\text { per }}$ degree of increase in temperature (feet). |
| :---: | :---: | :---: | :---: | :---: | :---: |
| SW. $\frac{1}{4}$ sec. ${ }^{\circ} 31$, T. 117, R. 63 | 866 | 45 | $a 66.7$ | b 42.9 | 36.4 |
| East center of sec. 7, T. 117, R. 63 | 972-1,020 |  | a 68.7 | 42.9 | 37.3 |
| NW. $\frac{1}{}$ sec. 32, T. 117 , R. 62. | - 810 | 300 | $a 64.4$ | 42.9 | 37.7 |
| SE. i sce. 27, T. 117, R. 61. | 900 | 110 | ${ }^{6} 62.2$ | 42.9 | 46.6 |
| Doland....'. | 895 | 370 | c 69 | 42.9 | 34.2 |
| NW. $\frac{1}{2}$ sec. 33, T. 116, R. 65 | 900 | 120 | a 67.2 | 42.9 | 37 |
| SE. $\frac{1}{\text { sec. } 13, \mathrm{~T} .116, \mathrm{R} .65 .}$ | 890 |  | a 66.7 | 42.9 | 37.4 |
| Redfield...... | 964 | 1,260 | c 70.1 | 42.9 | 35.4 |
| Redfield, ${ }^{\text {a }}$ | 920 |  | a 70.1 | 42.9 | ${ }_{32}^{33.8}$ |
| Redficld... | 944 |  | a 71.7 | 42.9 | 32.8 |
| NE. $\frac{1}{4} \mathrm{sec} .35, \mathrm{~T} .116, \mathrm{R} .64$ | 949 |  | a 64.4 | 42.9 | 44.1 |
| NW. $\frac{1}{\text { sec. 31, T. 116, R. } 64}$ | 912 |  | ${ }^{a} 66.7$ | 42.9 | 38.3 |
| SE. $\frac{1}{2}$ sec. 9 , T. 116, R. 63. | 851 | 150 | ${ }^{a} 62.2$ | 42.9 | 44.1 |
| East center of sec. 35, T. 116, R. 63 | 900 |  | ${ }^{a} 64.2$ | 42.9 | 42.3 |
| East center of sec. 25, T. 116, R. 63 | 880 |  | ${ }^{6} 62.2$ | 42.9 | 45.6 |
| SW. $\frac{1}{4} \mathrm{sec} .4, \mathrm{~T} .116, \mathrm{R} .62$. | 875 | 50 | ${ }^{a} 64.5$ | 42.9 | 40.5 |
| South center of sec. 34, T. 116, R. 61 | 915 |  | ${ }^{6} 62.5$ | 42.9 | 46.7 |
| SE. $\frac{1}{\text { d }}$ Sec. 13, T. 115, R. 64 | 942 |  | ${ }^{a} 64.4$ | ${ }^{\text {d }} 43$ | 45 |
| NW. 4 sec. 9, T. 115 , R. 64 | 930 |  | ${ }^{6} 6.5$ | c 42.9 | 41.7 |
| SW. ${ }^{\text {sec. } 10, \mathrm{~T} .115, \mathrm{R} .64}$ | 935 |  | ${ }^{\text {a } 65.4}$ | 42.9 | 41.5 |
| SW. $\frac{1}{2} \mathrm{sec} .33, \mathrm{~T} .115, \mathrm{R} .63$ | 867 |  | ${ }^{a} 65.4$ | 42.9 | 38.5 |
| SE. $\frac{1}{2}$ sec. 1, T. $115, \mathrm{R} .63$ | 988 |  | ${ }^{a} 62.2$ | 42.9 | 51.2 |
| NW. $\frac{1}{2} \mathrm{sec} .29$, T. $115, \mathrm{R} .62$ | 936 |  | a64. 5 | 42.9 | 43.3 |
| NE. $\frac{1}{\text { sec. }}$ 22, T. 115, R. 61 | 885 | 75 | ${ }^{\text {a } 62.5}$ | 42.9 | 45.2 |
| NE. $\frac{1}{2 s e c . ~ 23, ~ T . ~ 115, ~ R . ~} 61$ | 885 | 75 | ${ }^{6} 65.3$ | 42.9 | 39.6 |
| NW. d sec. 27, T. 115, R. 61 | 450 |  | ${ }^{5} 56.2$ | 42.9 | 33.8 |
| West, center of sec. 6, T. 115,R. 60 | 940 | 20 | ${ }^{a} 66.7$ | 42.9 | 39.5 |
| SW. $\frac{1}{\text { sec. } 9, \text { T. } 115 \text {,'R. } 60 . . . . . . . ~}$ | 886 |  | ${ }^{a} 67.7$ | 42.9 | 35.7 |
| SE. $\frac{1}{4} \mathrm{sec} .3, \mathrm{~T} .115, \mathrm{R} .60$ | 900 |  | ${ }^{a} 67.0$ | 42.9 | 37.3 |
| Glidden, sec. 32, T. 114, R. 6 | 1,973 | 550 | a 70: 1 | 42.9 | 35.7 |
| NE. $\frac{1}{4}$ sce. 29, T' 114, R. 65 | 1,029 | 35 | ${ }^{a} 71.5$ | 42.9 |  |
| NE. $\frac{1}{\text { sec. } 30, \text { T. 114, R. } 62}$ | 840 | 200 | ${ }^{\text {a } 67.0}$ | 42.9 | 34.8 |
| SW. $\frac{1}{2}$ sec. 5, T. 114, R. 62 | 835 |  | ${ }^{6} 64.5$ | 42.9 | 38.6 |
| SW. sec. 7, T. 114, R. 61 | 770 | 25 | $a 62.2$ | 42.9 | 40.2 |
| NW. $\frac{1}{4} \mathrm{sec} .19, \mathrm{~T} .114, \mathrm{R} .61$ | 800 |  | ${ }^{a} 66.7$ | 42.9 | 33.6 |
| SW. $\frac{1}{}$ sec. 27, T. 114, R. 61 | 744 |  | ${ }^{a} 64.4$ | 42.9 | 34.6 |
| NE. $\frac{1}{2}$ scc. 23, T. 114, R. 61. | 790 |  | ${ }^{a} 65.4$ | 42.9 | 35.1 |
| West center of sce. 28, T. 114, R. 61 | 735 |  | ${ }^{a} 65.3$ | 42.9 | 32.8 |
| NW. $\frac{1}{4} \mathrm{scc} .111, \mathrm{~T} .114, \mathrm{R} .60$ | 884 |  | ${ }^{6} 699.6$ | 42.9 | 33.1 |
| SE. $\frac{1}{\text { sec. } 18, \mathrm{~T} .114, \mathrm{R} .60 .}$ | 860 |  | ${ }^{a} 63.7$ | 42.9 | 33.3 |
| SE. $\frac{1}{4}$ sec. 17, T. 114, R. 60 | 848 |  | ${ }^{6} 67$ | 42.9 | 35.2 |
| NE. $\frac{1}{\text { sec. } 23, \mathrm{~T}} .114, \mathrm{R} .60$ | 835 |  | a 67.5 | 42.9 | 33.9 |
| SW. $\frac{1}{2}$ sec. 27, T. 114, R. $60 \ldots . .$. | 850 | 30 | ${ }^{6} 62.2$ | 42.9 | 44 |
| Todd County: Rosebud Agency, 18 mi east of $f$ | 2,500 | No. | 153좆ㄱ | $g 46.9$ | 23.5 |
| Yankton County: $h$ |  |  |  |  |  |
| Yankton. | 489-595 | 1,450 | ${ }^{1} 62$ | j46.5 | 31.5 |
| Do. | $432-455$ $610-615$ | 330 880 | ${ }_{6} 62$ | 46.5 | 39.3 |
| Yankton Asylum | $600-672$ | 165 | 164 | 46. 5 | 34.3 |
| Yankton cement works | 450-500 | 1,300 | ${ }^{6} 64$ | 46.5 | ${ }^{24.3}$ |
| Sec. 12, T. $93, \mathrm{R} .56$. | 375 380 | 8 | $a$ $\square$ 60 | 46.5 |  |
| Sec. 11, T. 93, R. 56. | 380 | 12 | $\square$ $\square$ $a$ 60 | ${ }_{46.5}^{46.5}$ | 28.2 29.6 |
| Sec. 19, T. 93, R. 56. | 400 | 25 | a 60 | 46.5 | 29.6 |

a J. E. Todd, observer.
b Redińeld average for 19 years
c J. II. Sheppard, observer. Nettleton gives $64^{\circ}$ for flows from both wells.
d Redficld average plus $0.1^{\circ}$.
e Redfield average
$f$ Darton thermometer sunk to bottom.
$o$ Average at Cutmeat for 23 years
h 51sti Cong., 1st sess., S. Ex. Doc. 222, p. 115, 1890; Artesian and underfiow investigation: 52d Cong.,
1st sess., S. Ex. Doc. 41, pt. 2, p. 65, 1892
E. S. Nettieton, observer.
$f$ Average for 27 years.

## SUMMARY.

Of the 281 wells given in the above list all but 8 are in the region east of the Missouri. A large proportion of these are in Beadle, Brown, Faulk, Edmunds, Hand, and Spink counties, where many are close together. Records from representative wells in Aurora, Bon Homme, Buffalo, Charles Mix, Davison, Day, Douglas, Gregory, Hughes, Hutchinson, Hyde, Jerauld, Kingsbury, Marshall, .Sanborn, and Yankton counties indicate the geothermal gradient in an area of wide extent. The deep artesian wells at Cheyenne Agency, with a rate of increase of $1^{\circ}$ in about 35 feet; at the head of Oak Creek, northeast of Rosebud, with a rate of $1^{\circ}$ in $23 \frac{1}{2}$ feet; and at Kennebec, with a rate of $1^{\circ}$ in nearly 26 feet, extend the area. The flow from the Highmore well is too small to give a reliable indication of the temperature of the water at the source of the flow. In Plate I are shown the principal data for the eastern part of the State. The figures indicate feet to the degree of increase in representative wells. The lines connect wells of the same rates for $20,25,30,35,40,45$, and 50 feet to the degree and delimit the zones of intermediate rates. They are not isogeotherms, but if the well temperatures were calculated to some uniform depth, such as 500 to 1,000 feet, and these connected by isogeothermal lines, these lines would be parallel to the rate lines on the map.

In general, it will be seen from Plate I that the rate is nearly uniform in groups of wells, but it varies from place to place with most remarkable regularity. In the James River valley from Ashton to Aberdeen it is $1^{\circ}$ in 40 to 45 feet. To the west and to the south through Redfield and in the east-central part of Beadle County it is $1^{\circ}$ in less than 35 to 40 feet. At Wolsey and Alpena it is $1^{\circ}$ in about 23 feet, and at Pierre, Crow Creek, and Chamberlain it is $1^{\circ}$ in less than 30 feet. This zone of high rates extends down the Missouri Valley through Springfield and Yankton counties, S. Dak., and the northern part of Boyd, Knox, Cedar, and Dixon counties, Nebr. At Fort Randall the rate reaches a maximum of $1^{\circ}$ in $17 \frac{1}{2}$ feet. The rate decreases rapidly to the north and east, as shown by scattered wells in Brule, Aurora, Davison, Douglas, Hutchinson, Sanborn, and Bon Homme counties. The rapid change from Springfield to Tyndall and the regular bat less rapid diminution from Alpena to Mitchell are notable features in the "cooler" area.

The records in the western part of the State are diverse. The flow from a 515 -foot well near Belle Fourche had a temperature of $54^{\circ}$, indicating a rate of $1^{\circ}$ in $51 \frac{1}{2}$ feet, but a flow at 323 feet gave a less rate. The flow from the Orman well, 1,417 feet deep, with a temperature of $94^{\circ}$, indicated a rate of $1^{\circ}$ in 28.3 feet, and the large flow from the 2,985-foot well at Edgemont indicated a rate of $1^{\circ}$ in


## EXPLANATION



Underground contours (Under artesian basin) Contour interval 100 ft . broken line approximate


Temperature curves Figures show depth in feet per degree of increase in temperature as indicated by flows of representative deep

39 feet, or nearly the same as the rate indicated by the record of a thermometer I sunk in an old 725 -foot hole at Buffalo Gap, 30 miles. northeast of Edgemont

## EOLOGIC RELATIONS.

Except the Edgemont wells, which draw from Cambrian sandstone, the many wells given on the list draw from the Dakota and associated sandstones. In the western part of the State these sandstones overlie "Red Beds" and a thick succession of Paleozoic limestones and sandstones, but east of the Missouri they lie directly on pre-Cambrian granite or quartzite. They are overlain by shale of Upper Cretaceous-age, which is 1,000 feet or more thick to the north but less to the south and along Missouri River from Chamberlain to Yankton. The strata are not flexed and present a gentle down slope to the west over a wide area. The sections in figure 3 show the relations.
On Plate I are shown the principal features of configuration of the granite and quartzite bedrock floor on which the water-bearing beds lie. This floor has been reached by numerous holes, so that its position is well known at many localities. A more detailed $137163^{\circ}-20-$ Bull. $701-6$
representation of it and the data used in constructing the map are set forth in my report on the geology of South Dakota. ${ }^{1}$

## TENNESSEE.

Only two well temperatures that appear to throw light on the geothermal gradient have been reported from Tennessee. One is at Covington, where, according to Glenn, ${ }^{2}$ the railroad well, 533 feet deep, has water within 31 feet of the surface, with a temperature of $66^{\circ}$. As the mean annual air temperature at this place is $60.1^{\circ}$, a rate of increase of $1^{\circ}$ in 90 feet is indicated. The water is in coarse sand under clay in the young deposits of the Mississippi embayment.

## TEXAS.

## TEMPERATURES.

No special determinations of underground temperature have been made in deep holes in Texas, but there are numerous records of the temperature of waters from artesian wells. Some of these records are reliable, and a few of them indicate the depth from which the flow is derived. The data are given in the following list. They are derived partly from published reports and partly from replies to circulars. Some late observations have been made by Mr. C. E. Van Orstrand, but the results are not yet published.

Temperatures in deep wells in Texas.

| Location. | Depth of main flow (feet). | Volume of flow (gallons $\underset{\substack{\text { per min } \\ \text { ute }}}{ }$ ute). | $\begin{aligned} & \text { Tem- } \\ & \text { pera- } \\ & \text { ture } \\ & \left({ }^{\circ} \mathrm{F} .\right) . \end{aligned}$ | Mean annual tem-perature $\left({ }^{\circ} \mathrm{F}\right.$.). | Depth per degree of in tem-pera(feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Bell County: Temple, 1 mile west of.a | 1,85 | Many. | 91 | ${ }^{6} 65.8$ | $73 \frac{1}{2}$ | "Water from bottom." |
| Bexar County: San Antonio, 2 miles south of.c | 1,900 | 500 | 106 | d 68.0 | 50 | Sulphur water. Small flows at intervals below 1,535 feet, possibly not cased off. |
| Bosque County: Clifton, 3 miles north of.e | 662 | 220 | 84 | 67.1 | 39 | Water also at 182 and 610 feet, probably cased off. |
| $a$ Hill, R.T., Geography and geology of the Black and Grand prairies, Tex.: U. S. Geol. Survey Twentyfirst Ann. Rept., pt. 7, pp. 529-530, 1901. <br> $b$ Average for 27 years. <br> c Hill, R. T., and Vaughan, T. W., Geology of the Edwards Plateau and Rio Grande Plain adjacent to Austin and San Antonio, Tex.: U. S. Geol. Survey Eighteenth Ann. Rept., pt. 22, pp. 294-297, 1897. <br> ${ }^{d}$ Average for 32 years plus $0.1^{\circ}$. <br> e Hill, R. T., op. cit., p. 490. <br> $f$ Fenneman, N. M., Oil fields of Texas-Louisiana Gulf Coastal Plain: U. S. Geol. Survey Bull. 282, pp. 49, 56, 1906. <br> $\theta$ Idem, p. 87. |  |  |  |  |  |  |
|  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |
| ${ }^{1}$ Darton, N. H., U. S. Geol. Survey Water-Supply Paper 227, pl. 10, pp. 33-40, 1909. <br> ${ }^{2}$ Glenn, L. C., Underground waters of Tennessee and Kentucky west of Tennessee River and of adjacent rea in Illinois: U. S. Geol. Survey Water-Supply Paper 164, p. 117, 1906. |  |  |  |  |  |  |

Temperatures in deep wells in Texas-Continued.

| Location. | Depth of main flow (feet). | Volume of flow (gallons per minute). | $\begin{aligned} & \text { Tem- } \\ & \text { pers- } \\ & \text { ture } \\ & \left({ }^{\circ} \mathrm{F} .\right) . \end{aligned}$ | $\begin{gathered} \text { Mean } \\ \text { annual } \\ \text { tem- } \\ \text { pera- } \\ \text { ture } \\ \left(^{\circ} \mathrm{F} .\right) \end{gathered}$ | Depth per degree of increase in tem-perature (feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Burleson County: |  |  |  |  |  |  |
| Myers...... | 860 | 30 | 90 | $a 67.9$ | 38 |  |
| Bryan, 12 miles fr | 835 | 35 | 85 | $a 67.9$ | 49 |  |
| Do............ | 635 | 22 | 76 | $a 67.9$ | 78 |  |
| Do.. | 550 | 20 | 76 | $a 67.9$ | 63 |  |
| Whittaker, $3 \frac{1}{2}$ miles southeast of. $b$ | 850 | 50 | 77 | $a 67.9$ | 88 | Gas escapes also. |
| Denton County: Denton c.. | 550 |  | 72 | 64.4 | 70 |  |
| El Paso County: Fort Bliss. | 403 | Many. | 76 | d 62.9 | 31 | Water rises within 130 fect of surface. |
| Falls County: Marlin e. | 3,330 | 150 | 147 | $f 67.2$ | 4112 | Water sand at 3,310 to 3,330 feet. |
| Many other wells in county. | 1,750-1,850 | 150-300 | 103 | 67.2 | 50 |  |
| Galveston County: |  |  |  |  |  |  |
| Alta Loma, 30 wells. | 875-950 | 310 | 75-78 | $g 69.4$ | 101 | Water from 740 to 868 feet. |
| Fairwood $h_{\text {. }}$ | 575 | 50 | 78.5 | 69.4 | 63 | Water from 575 feet. |
| High Island $i . . . . . . . . . . .$. | 142-203 | Many. | 100+ |  |  | Strong brine; 7 wells. |
| Hitchcock f............. | 726 | . 70 | 77 | 69.4 | 89 | Water sand from 678 to 726 feet. |
| Hitchcock, $1 \frac{1}{2}$ miles northwest of. | 710 | 100 | 77 | 69.4 | 93 | 16 pounds; water from 710 feet. |
| Hitchcock, near......... | 768 | 45 | 78 | 69.4 | 89 |  |
| Galveston, 28th and Church streets. $k$ | 810 | 125 | 83 | 69.4 | 60 | Water from 810 feet; brackish. |
| Galveston, 36th and Church streets. $k$ | 1,365 | 243 | 84 |  |  | Water from 810 (probably) and 1,365 feet. |
| Galveston, 10 miles southwest of. $k$ | 827 | 37 | 80 | $70 \pm$ | 83 | Two wells; water from 827 feet. |
| Do................ | 797 |  | 83 | 69.4 | 58 |  |
| Grimes County: Navasota ${ }^{\text {l }}$. | 830 |  | 80 | m 68.1 | 70 | Water also at 250 fect, probably cased off. |
| Hardin County: |  |  |  |  |  |  |
| Batson ${ }^{n} . .$. | 1,200 | Many. | 125 | 069.1 | $21 \frac{1}{2}$ | Oil well. Salt water from 1,200 feet. |
| Batson, Higgins No. 4.. | 1,159 | Many. | 101 | 69.1 | 34 | Water 970-985 feet. Salt water from sands 1,080 to 1,130 feet. |
| Saratoga, 3 miles west of. $p$ | 958 |  | 100+ | 69.1 | 30 | From 938 to 950 feet. |
| Sourlake.............. | $880+$ | Many. | 100 | 69.1 | 27즐 | Flow from 850 to 880 feet in a well drilling for oil. |
| Sourlake | 985 |  | 101 | 69.1 | 31 | Saline flow at 985 feet from |
| Harris County: |  |  |  |  |  | a well drilling for oil. |
| Cedar Bayouq. | 727 | 5 | 76 | $r 68.9$ | 101 |  |
| Cedar Bayou, 3 miles northwest of. | 410 | 30 | 76 | 68.9 | 86 |  |
| Cedar Bayou. | 548 | 14 | 73 | 68.9 | 128 | Water at 525 feet; also at 320 feet. |
| Harrisburg, 즐 mile east of. | 640 | Flow. | 72 | 68.9 |  | Flow from? |
| Honston, Commerce and Fannin streets. | 900 | 25 | ${ }^{8} 82$ | 68.9 | 69 | Water also at about 480 feet. Pumped. |
| Houston, Y. M. C. A... | 500 | 200 | 75 | 68.9 | 82 | Water about 500 fect. Pumped. |

a Average at College Station for 27 years.
$b$ Taylor, T. U., Underground waters of Coastal Plain of Texas: U. S. Geol. Survey Water-Supply Paper 190, p. 60, 1907.
c Tifty-first Cong., 1st sess., S. Ex. Doc. 222, p. 264, 1890.
d El Paso average for 29 ycars.
e Hill, R. T., op. cit., p. 645.
$f$ Mean annual temperature of Waco for 28 years plus $0.1^{\circ}$.
$g$ Galveston avierage for 46 years.
$h_{\text {Texas Geol. Survey Fourth Rept., p. 101, } 1892 . ~}^{\text {R }}$
$i$ Deussen, Alexander, Geology and underground waters of the southeastern part of the Texas Coastal Plain: U. S. Geol. Survey Water-Supply Paper 335, pp. 139-141, 1914.
${ }_{k}$ See Singley, J. A., Report on artesian wells of Gulf coastal slope, p. 102, Austin, June, 1893.
$k$ Deussen, Alexander, op. cit., pp. 157, 160, 164.
$i$ Fifty-first Cong., 1st sess., S', Ex. Doc. 222, p. 341, 1890.
$m$ College Station average plus $0.2^{\circ}$.
${ }^{n}$ Fenneman, N. M., Oil fields of Texas-Louisiana Gulf Coastal Plain: U. S. Geol. Survey Bull. 282, pp.49, 56, 1906
o Beaumont average for 16 years.
$p$ Hayes, C. W., and Kennedy, William, Oil fields of Texas-Louisiana Gulf Coastal Plain: U. S. Geol. Survey Bull. 212,'pp. 59-60, 1903. Also in Fennernan, N. M., op. cit., p. 47.
$q$ Average for 16 years.
$r$ Houston average for 27 years.
${ }^{s}$ U. S. Geol. Survey Water-Supply Paper 335, p. 230, 1914.

Temperatures in deep wells in Texas-Continued.

| Location. | Depth of main flow (feet). | Volume of flow (gallons per minute). | Tem-perature ( ${ }^{\circ} \mathrm{F}$.). | Mean annual tem-perature ( ${ }^{\circ} \mathrm{F}$.). | Depth per degree of increase in tem-perature (feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Hays County: San Marcos. | 1,490 | 100 | 73 | 67.5 | 271 | Small flow at 1,345 feet, probably cased off. |
| Jefferson County: <br> Beaumont, 1 mile west of. | 1,034 | 8 | a 6 | b 69.1 | 204 | Flow from 1,000 feet; higher flows may mingle. |
| Port Arthur............ | 1,400 $\pm$ | Many. | a 80 | 69.1 | $128 \pm$ | Flows mineral water. Source not known. |
| La Salle County: Cotulla c. | 1,008 | Many. | 86 | d 71 | 67 | Flow supposed to be from 1,008 feet. |
| Matagorda County: Big Hill. | 722 | Many. | 99.6 | e 70 | $23 \frac{1}{2}$ | Flow from 700-730 feet. |
| McLennan County: <br> Bruceville. | 1,565 | 150 | 95 | $f 67.1$ | 56 | Water at 500 feet probably cased off. |
| China Springs, 12 miles northwest of Waco. | 1,380 | Many. | 102 | 67.1 | 37 | Flows at $500,600,800$, and 1,100 feet, probably cased off. |
| Hermosa, near Waco... | 1,730 | 150 | 103 | 67.1 | 48 | Flows at 670 and 1,400 feet. <br> (Hill.) |
| McGregorg.............. | 1,030 | Many. | 82 。 | 67.1 | 69 | Water also at 502 feet. (Hill.) |
| McGregor. . . . . . . . . . . . | 991 | 350 | 82 | 67.1 | $66 \frac{1}{2}$ |  |
| Waco: <br> 5 miles west of. $\qquad$ | 1,470 | 350 | 90 | 67.1 | 64 | Cased to 735 feet. |
| Moore's well $\boldsymbol{h}$. | 1,852. | Many. | 103 | 67.1 | $51 \frac{1}{2}$ | First water at 1,000 feet. (Hill.) |
| 11 wells of Water Co. ${ }^{i}$ | 1, 828-1, 860 | Many. | 103 | 67.1 | 51 | Flows at 700-800 and 1,100 feet also. |
|  | 1,607 | Many. | 97 | 67.1 | $53 \frac{1}{2}$ | Do. |
|  | 1,812 | Many. | 100 | 67.1 | 55 | Do. |
|  | 1,820 | Many. | 1021 | 67.1 | 51 | ${ }^{\text {Do. }}$ |
|  | 1,860 | Many. | 103 | 67.1 | 52 | Flow from 1,800 to 1,815 feet. (Hill.) |
|  | 1,776 | Many. | 103 | 67.1 | $49 \frac{1}{2}$ |  |
|  | 1,866 | Many. | 1032 | 67.1 | 51 | Sandstone, 1,760-1,861 feet. <br> (Hill.). $j$ |
| Nacogdoches County: Nacogdoches. | 500 | 500 | k 74 | $l 65.4$ | $39 \frac{1}{2}$ | Flow from 340-500 feet. |
| Navarro County: <br> Corsicana No. 1 | 2,483 | 200 | 126 | $m 66$ | 41 | Flow at bottom. |
| Corsicana No. 2. | 2,515 | 200 | 126 | 66 | 40 | Water sand, 2,400 to 2,460 feet. $n$ |
| Corsicana No. 3. | 2,500 | 200 | 126 | 66 | $41 \frac{1}{2}$ |  |
| Tarrant County: <br> Fort Worth............. | 3,250 | Many. | - 126 | p 65 | 187 | Cased to main flow at 1,127 feet. |
| Marine, $\frac{1}{2}$ mile northeast | 1,200 | 545 | 78 | 65 | 74 | Water from 960-1,000 feet. 7 |
| Travis County: <br> Austin Natatorium r... | 1,875 | 300 | 100 | 867.4 | $t 51$ | Cased to 1,100 feet. Small flow at 1,215 feet; not cased off. Mainflow 1,875 feet from sandstone (1,675 to 1,875 feet). |
| Manor ${ }^{\text {. }}$. | 2,560 | 700 | 93 | 67.4 | $\bigcirc 481$ | Water-bearing rock 1,250 to 1,400 feet. |
| Waller County: Hempstead | 1,131 | 100 | 80 | ${ }^{w} 68.6$ | $97 \frac{1}{2}$ | Flow from 1,110 feet. |
| Williamson County: Roundrock. | 1,400 | Many. | 95 | $x 67.4$ | 51 | Pumped. |
| Zavalla County: Carrizo Springs, 22 miles north. | 910 | Several. | ע 90 | 272 | 40 | Cased 720 feet to main flow. |

[^36]$m$ Hill, R. T., op. cit., p. 311
$n$ Hill, R. T., op. cit., p. 641.

- Idem, p. 576.
$p$ Average for 22 years.
$q$ Hill, R. T., op. cit., p. 577.
r U.S. Geol. Survey Eighteenth Ann. Rept., pt. 2, p. 284.
${ }_{8}$ Average for 61 years.
$t$ Calculated to 1,675 feet.
u U. S. Geol. Survey Eighteenth Ann. Rept., pt.
2, pp. 284-286, 1898; Twenty-first Ann. Rept., pt. 7,
pp. 511-514, 1901; Water-Supply Paper 335, p. 355 ,

1914. 

$v$ Calculated to 1,250 feet.
whouston average minus $0.3^{\circ}$.
$x$ Austin average.
$y$ Eagle Pass average plus $0.1^{\circ}$.
2 U. S. Geol. Survey Bull. 298, p. 286, 1906.

SUMMARY.
It is difficult to judge the relative values of the Texas temperature records. Some may not be correct, and others may not represent the temperature at the depth stated owing to the mingling of flows. The flow from the deep well at Marlin is supposed to come from a depth of 3,350 feet, and if so it indicates a rate of $1^{\circ}$ in 42 feet. The several wells at Waco, not far southeast of Marlin, give rates of $1^{\circ}$ in $49 \frac{1}{2}$ to 53 feet. Other wells in the same county (McLennan) indicate $1^{\circ}$ in $48,64,66 \frac{1}{2}, 69$, and 56 feet, but the apparent lower rate is possibly due to the ingress of higher waters. The warm flow from a depth of 1,380 feet at China Springs, 12 miles northwest of Waco, indicates the exceptionally high rate of $1^{\circ}$ in nearly 37 feet. A similar rate is indicated by the Clifton well in Bosque County, about 25 miles northwest of China Springs. The three deep wells at Corsicana, still farther north, give a rate of $1^{\circ}$ in 40 to $41 \frac{1}{2}$ feet. Fort Worth has a rate of $1^{\circ}$ in $18 \frac{1}{2}$ feet if all the flow is derived from a depth of 1,127 feet, as stated. The rate at Nacogdoches may be as high as $1^{\circ}$ in $39 \frac{1}{2}$ feet if the flow comes from a depth of 340 feet. The Austin wells indicate rates of $1^{\circ}$ in $48 \frac{1}{2}$ and 51 feet, but the true rate here may be considerably less if the flow is from a deeper source than the depth used in the calculations. The Round Rock figure is uncertain, but the rates for Hempstead and Navasota appear to be valid. The data for Harris County, Port Arthur, and Beaumont are not satisfactory, and the wells 2 miles south of San Antonio and at Cotulla and Temple may not give reliable data. The temperature given for the San Marcos well is doubtless unreliable. The five wells in Hardin County appear to have rates of $1^{\circ}$ in $21 \frac{1}{2}, 27 \frac{1}{2}, 30,31$, and 34 feet, without likelihood of much error. The Big Hill 722-foot well has a similar rate, and the High Island well evidently indicates a steep rate. The rate indicated by the Galveston County wells varies, but the figures 60 and 63 feet to the degree appear the most accurate, there being uncertainty as to the source of the water in the others. As the water at Fort Bliss is pumped the temperature given may not be as high as that in the bottom of the well, but the figure given $\left(76^{\circ}\right)$ indicates that the rate is $1^{\circ}$ in at least 31 feet.

## GEOLOGIC RELATIONS.

The wells in the above list penetrate a great variety of formations, which are mostly, however, of Cretaceous and Tertiary age. In most parts of the area the shales are tilted at low angles and in places traversed by faults. No igneous rocks are reported near any of the wells listed, and the granites or other old bedrocks are far beneath the surface. The wells in the southeastern part of the State and along the coast draw from sands of Tertiary age. In the oil fields of Hardin County there are sharp local uplifts or domes. The Nacogdoches well draws from lower beds in the Eocene.

The group of wells extending from north to south across the center of the State draw their waters from sands in the Trinity group, which is low in the Cretaceous. The deep boring at Fort Worth passed through the base of these sands and penetrated the Paleozoic rocks for more than 2,000 feet. At Marlin, in Falls County, where the lowest sand lies deep, it was necessary to sink 3,330 feet to reach it. The Corsicana wells draw from the Woodbine sand, in the Upper Cretaceous, considerably above the highest beds of the Trinity. The 1,900 -foot well 2 miles south of San Antonio is believed to have not quite reached the base of the Edwards limestone. ${ }^{1}$ The San Marcos and San Antonio wells are situated near a great fault. The Fort Bliss well draws from the Quaternary valley fill.

## UTAH.

A few well temperatures in the Milford-Beaver region, in southwestern Utah, have been reported by Lee. ${ }^{2}$

Temperatures in wells in Utah.

| Location. | Depth (feet). | Flow (gallons per minute). | Temperature ( ${ }^{\circ} \mathrm{F}$.). | Depth per degree of increase in temperature (feet). $a$ | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Milford. | 425 | 25 | 72 | 18 |  |
| Do | 750 | 37 | 80 | 24 |  |
| Do. | 381 | 13 | 65 | 23 |  |
| Do. | 280 | 7 | 68 | 14 |  |
| Milford, railroad well. | 310 | 30 | 71 | 14 |  |
| Sec. 28, T. 29 S., R. 7 W | 213 |  | 54 | $36 ?$ | Pumped. |
| Sec. 36, T. 29 S., R. 8 W | 314 | 7 | 74 | 12 | Flows at 200 and 314 feet. |
| Sec. 28, T. 29 S., R. 8 W | 385 |  | 53 | - $77 ?$ |  |
| Sec. 5, T. 28 S., R. 10 W | 286 | 8 | 71 | 12 |  |
| Greenville School....... | +244 | 11 | 68 | 12 |  |
| Neels. | 1,998 | ...... | Hot. |  | Hot water 1,205 feet and below. Reached granite. |

$a$ Calculated from Milford mean for 1916, 48.2 ${ }^{\circ}$. Water in wells 60 to 66 feet deep is stated to have a temperature of $51^{\circ}$ to $55^{\circ}$.

The waters are decidedly thermal considering the slight depths of the wells. The cause of the heat is not apparent. Probably the two wells in the list showing temperatures of $53^{\circ}$ and $54^{\circ}$ are cooled by pumping or by the ingress of shallow water.

All these wells are in the Quaternary clay, sand, and gravel deposits of the Lake Bonneville basin. The boring at Neels reached granite.

[^37]
## VIRGINIA.

## RICHMOND COAL FIELD.

In 1836 to 1842 Rogers ${ }^{1}$ determined temperatures in the coal mines in the Newark group west of Richmond, Va., the results of which he discussed at considerable length. The coal measures consist mostly of sandstone lying in a basin underlain by granitic rocks; the coal is in the basal beds. The strata have steep dips and are greatly broken, and owing to this condition a large amount of water penetrates into the mines. The following results were obtained:

Temperatures in coal mines near Midlothian, Va.

| Location. | Depth (feet). | $\begin{gathered} \text { Tempera- } \\ \text { ture } \\ \left({ }^{\circ} \mathrm{F} .\right) . \end{gathered}$ | $\begin{gathered} \text { Depth } \\ \text { per de- } \\ \text { gree of in- } \\ \text { crease in } \\ \text { tempera- } \\ \text { ture } \\ \text { (feet).a } \end{gathered}$ | Remarks. |
| :---: | :---: | :---: | :---: | :---: |
| Greenhole pit. | 100 | 58.0 |  | In pool of water in bottom. |
| Mills \& Reed's pit | 318 | 59.5 |  | In large flow of water from galleries. |
| Do. | 375 | 61.0 | 125 | Do. |
| Do. | 420 | 63.0 | 84 | Do. |
| Black Heth engine pit | 412 | 63.0 | 83 | Pool of water in gallery not working, also in floor. |
| Wills \& Michael's pit | 386 | 62.0 | 96 | Poolin bottom main shaft. |
| Black Heth pit.. | 570 | 65.5 | 76 | Pool in lowest level. |
| Black Heth new shaft. | 380 | 61.5 |  | Crevice in rock in bottom cooled by drippings from above. |
| Midlothian 1842 shaft. | 330 | 61.75 | 88 | Spring in wall. |
| Do. | 600 | 66.25 | 74 | Small streams at bottom some distance above coal. |
| Midlothian 1839 shaft. | 780 | 68.75 | 72 | Water mixed with some drippings from higher up. |

$a$ Calculated from mean annual air temperature at Richmond, $58.5^{\circ}$ (37-year average to 1916) minus $0.5^{\circ}$ for difference in the altitude and iatitude. Rogers used $56.75^{\circ}$.

The first observation was made in a mine in operation and the last four in shafts just completed or in progress. Exclusive of the three first and one in the new Black Heth shaft, the average rate indicated by the others is $1^{\circ}$ in 82 feet.

The figures indicate a less rate in the old workings than in the new shafts, a difference which may be due to the cooling by seepage and ventilation. Rogers made the following comments on the results:

Comparing the three last observations, the two former of which were made in the same shaft and the last in one only a few hundred feet removed, there is ground for inferring that the rate at which the temperature increases grows less as the depth augments. In descending from 330 feet to 600 -that is, through 270 feet-the rise of temperature is $4.5^{\circ}$; while in descending from 600 to 780 , or through 280 feet, the rise is only $2.5^{\circ}$. This difference would, I think, have been less could I have obtained the temperature at 780 feet free from the cooling influence of the copious drippings from above. Yet even with the most liberal allowance, there would still remain evidence of a diminishing rate of increase with the depth.

[^38]Rogers considered the observations at 330 and 600 feet.in the Midlothian shaft the most reliable, and these give a rate of $1^{\circ}$ in 88 and 74 feet, respectively; the difference between them is $1^{\circ}$ in 60 feet. The rate for the 780 -foot reading ( $1^{\circ}$ in 72 feet) is accordant, although, as suggested by Rogers, the water may be cooled slightly by drippings from above.

## FORT MONROE.

In 1902 a 2,130 -foot boring was made at Fort Monroe, and Capt. C. P. Townsley kindly lowered a Darton thermometer in it. After two hours the instrument recorded $92.5^{\circ}$. As the mean annual air temperature at Newport News is about $59.4^{\circ}$, a rate of $1^{\circ}$ in 64 feet is indicated. The boring passed through Tertiary and Cretaceous formations of the Coastal Plain to the basement of pre-Cambrian granite. The hole was full of salty water.

REEDVILLE.
At Reedville, in Northumberland County, the 85-gallon flow from a 680 -foot well is reported to have a temperature of $78^{\circ}$. As calculated from the mean annual air temperature at Warsaw, $56.4^{\circ}$, a rate of $1^{\circ}$ in $31 \frac{1}{2}$ feet is indicated.

## WASHINGTON.

## YAKIMA ARTESIAN BASIN.

There are numerous flowing wells in the vicinity of Yakima which afford data as to underground temperature. In most of them, however, several flows are intermixed, so that the observations do not indicate the temperature at any definite depth. There are three wells which reach the upper flow only and three others in which the casing is so set that only one flow can enter. They are in an area about $1 \frac{1}{2}$ miles long and within 120 feet elevation. The data for these six wells, as given by Geo. Otis Smith, ${ }^{1}$ are as follows:

Temperature at different depths in six wells in Yakima artesian basin, Wash

| Depth of flow (feet). | Flow (gallons per minute). | Temperature ( ${ }^{\circ} \mathrm{F}$. .). | Depth per degree of increase in temperature(feet).a |
| :---: | :---: | :---: | :---: |
| 515 | Many. | 70.7 | 25.0 |
| 637 | 120 | 72.2 | 28.9 |
| 835 | 60 | 73.2 | 36.1 |
| 625 | Many. | 73.2 | 27.0 |
| 752 | 300 | 72.2 | 34.0 |
| 820 | 360 | 72.7 | 36.2 |

[^39] Weather Bureau.

[^40]It was noted by Smith that the rate of increase is highest in the shallower wells.

Waring ${ }^{1}$ has reported that a large flow from a well on the Yakima Indian Reservation, in sec. 34, T. 11 N., R. 18 E., has a temperature of $68^{\circ}$. The main flow is from a depth of $507-512$ feet, possibly with a small admixture of water from 492 feet. The rate indicated by this observation is $1^{\circ}$ in about 28 feet.

Russell ${ }^{-2}$ has given temperatures of flows from several wells in the Yakima Valley, but the flows are small and the source of the water is not definitely located. One of these wells, in sec. 31, T. 13 N., R. 20 E., has a 2-gallon flow with a temperature of $73^{\circ}$ from a depth of 886 feet; another, in sec. 3, T. 12 N., R. 20 E., only 314 feet deep, has a flow with a temperature of $75^{\circ}$.

## WALLA WALLA.

In 1909 I noted the temperature of the large flows from several wells on the Blalock ranch west of Walla Walla. The flow from one well, 611 feet deep, had a temperature of $66.5^{\circ}$; that from another of similar depth, $67.25^{\circ}$; and that from a 563 -foot well, $67^{\circ}$. It is stated that the casings extend to a depth of 540 feet to a 20 -foot sheet of basalt under which is the water-bearing sand. As the mean annual air temperature of Walla Walla is $53.4^{\circ}$ (23-year average) a rate of $1^{\circ}$ in 41.4 feet is indicated, if the flow comes from a depth of 563 feet.

## WHITMAN COUNTY.

It is reported that the flow from a 110 -foot well at Pullman has a temperature of $60^{\circ}$, and that from a 176 -foot well at Tekoa a temperature of $76^{\circ}$, but these figures may not be reliable. As the mean annual air temperature at Colfax is $47.8^{\circ}$ and at Rosalia $46.5^{\circ}$, they indicate a very high rate. Doubtless heat from old lava flows is the cause.

## west virginia.

There are many deep borings in West Virginia, and temperatures in a number of them have been determined by the late William Hallock, ${ }^{3}$ by Johnson and Adams, ${ }^{4}$ and by Van Orstrand. ${ }^{5}$

[^41]
## WHEELING.

The well tested by Hallock was on Boggs Run, 4 miles southeast of Wheeling. It was 4,771 feet deep, and although water was found at moderate depths it was cased off by 1,570 feet of $4 \frac{7}{8}$-inch casing, so that the hole was dry when finished. The strata penetrated are nearly horizontal and begin high in the Carboniferous coal measures. Ordinary United States Signal Service self-registering thermometers were used. The following figures were obtained. Observations made two years later in water gave very closely accordant results. (See below.)

Temperatures at different depths in deep boring near Wheeling, W. Va.

|  | ${ }^{\circ} \mathrm{F}$. |  | ${ }^{\circ} \mathrm{F}$. |  | ${ }^{\circ} \mathrm{F}$. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 1,350 feet. | 68.75 | 2,625 feet. | 82. 20 | 3,625 feet. | 96.10 |
| 1,592 feet | 70.25 | 2,740 feet. | 83.65 | 3,730 feet. | 97. 55 |
| 1,745 feet | 71.70 | 2,875 feet. | 85.45 | 3,875 feet. | 100.05 |
| 1,835 feet. | 72.80 | 2,990 feet. | 86.60 | 3,980 feet. | 101.75 |
| 2,125 feet. | 76.25 | 3,125 feet. | 88.40 | 4,125 feet. | 104. 10 |
| 2,236 feet. | 77.40 | 3,232 feet. | 89.75 | 4,200 feet. | 105.55 |
| 2,375 feet | 79.20 | 3,375 feet. | 92. 10 | 4,375 feet. | 108. 40 |
| 2,486 feet. | 80.50 | 3,482 feet. | 93.60 | 4,462 feet. | 110.15 |

According to the Weather Bureau observations for 39 years the mean annual air temperature at Wheeling is $56.3^{\circ}$, but Hallock used $51.3^{\circ}$, the temperature found at a depth of 110 feet in a near-by coal mine. I believe $51.3^{\circ}$ is too low, for temperatures taken in mine galleries are misleading. Two years later water at a depth of 103 feet in the deep well was found to have a temperature of $52.5^{\circ}$. The rate calculated to a depth of 4,462 feet from a mean annual air temperature of $56^{\circ}$ is close to $1^{\circ}$ in 83 feet; the rate from 1,350 to 4,462 feet is $1^{\circ}$ in 75.2 feet; from 2,990 to 4,462 feet, $1^{\circ}$ in 62.5 feet.

After taking temperatures in the deep hole near Wheeling when it was dry, Hallock ${ }^{1}$ had opportunity to make an additional series two years later, when water stood within 40 feet of the top.

Comparison of temperatures in air and in water in deep well near Wheeling, W. Va.

| Depth <br> (feet). | Tempera- <br> ture in <br> water, <br> 1893 ( $\left.{ }^{\circ} \mathrm{F}.\right)$. | Tempera- <br> ture in <br> air, 1891, <br> inter- <br> polated <br> ( $\left.{ }^{\circ} \mathrm{F}.\right)$. | Water tem- <br> perature <br> minus air <br> tempera- <br> ture ( $\left.{ }^{\circ} \mathrm{F}.\right)$. |
| ---: | ---: | ---: | ---: |
|  |  |  |  |
| 1,586 | 70.12 | 70.12 | -0.03 |
| 1,921 | 73.95 | 73.82 | +.13 |
| 2,055 | 75.28 | 75.42 | -.14 |
| 2,276 | 78.13 | 77.93 | +.20 |
| 2,396 | 79.54 | 79.45 | +.09 |
| 2,539 | 81.21 | 81.15 | +.06 |
| 2,669 | $a 83.39$ | 82.75 | +.64 |
| 2,793 | 84.56 | 84.41 | +.15 |
| 2,937 | 86.12 | 86.07 | +.05 |
| 3,057 | 87.42 | 87.50 | -.08 |
| 3,196 | 89.27 | 89.30 | -.03 |
|  |  |  |  |

$a$ Doubtful.

[^42]These determinations show that there is no important difference between the air and water temperatures in a hole.

Hallock made a few observations at a deep boring at Radsford, W. Va., but as the hole was wet he did not complete the determinations.

## CHELYAN.

Johnson and Adams made many tests in the lower part of the 5,236foot hole at Chelyan, near Charleston. The results appear to indicate a temperature of about $129^{\circ}$ in the bottom. With a mean annual of $57.6^{\circ}$, average of 14 years at Charleston, this indicates a rate of $1^{\circ}$ in 73.3 feet.

## MANNINGTON.

The temperature determinations in the Hibbs No. 4 hole near Mannington were made with an electric resistance thermometer by Johnson and Adams and later with a mercury thermometer by Van Orstrand. The results were as follows, with rates of increase calculated from the 250 -foot reading subtracted from those at lower depths, all as indicated by the electric measurement.

Temperatures in Hibbs No. 4 hole near Mannington, W. Va.

| Depth <br> (feet). | Electric <br> thermom- <br> eter ( ${ }^{\circ} \mathrm{F}$ ). | Mercury <br> thermom. <br> eter ( $\left.{ }^{\circ} \mathrm{F}.\right)$ | Depth per <br> degree of <br> increase in <br> tempera- <br> ture (feet). |
| ---: | ---: | ---: | ---: |
| 100 | 56.0 | 53.6 | $\ldots \ldots \ldots \ldots .$. |
| 250 | 54.9 | 54.1 | $\cdots \cdots \cdots \cdots$ |
| 500 | 57.5 | 56.7 | 96 |
| 750 | 60.7 | 60.1 | 86 |
| 1,000 | 63.9 | 63.5 | 83 |
| 1,250 | 66.5 | 66.2 | 86 |
| 1,500 | 68.6 | 68.5 | 91 |
| 1,750 | 70.7 | 70.7 | 95 |
| 2,000 | 73.1 | 73.2 | 96 |
| 2,250 | 75.6 | 75.4 | 97 |
| 2,500 | 78.9 | 78.8 | 94 |

If calculated from the mean annual air temperature the temperature of $78.9^{\circ}$ in the bottom of the well indicates a gradient of $1^{\circ}$ in 93 feet.

Observations by Mr. Van Orstrand in a group of wells of the South Penn Oil Co. gave the following results:

Temperature measurements in borings of South Penn Oil Co., near Mannington, W. Va.


The rates of increase indicated by these observations are as follows: Well 14, 250 to 3,000 feet, $1^{\circ}$ in 104 feet; well 6, 100 to 1,980 feet, $1^{\circ}$ in 116 feet; well 2,100 to 2,650 feet, $1^{\circ}$ in 90 feet; well 4,100 to 3,225 feet, $1^{\circ}$ in 90 feet; well 7,100 to 2,888 feet; $1^{\circ}$ in 96 feet; well 9 , 100 to 2,500 feet, $1^{\circ}$ in 93 feet.

SPENCER.
The 3,403-foot hole bored by the Hope Gas Co. 22 miles east of Spencer, Clark County, was tested at 100 -foot intervals to 3,250 feet by Mr. Van Orstrand with the following results:

Temperature measurements in well 22 miles east of Spencer, W. Va.

|  | ${ }^{\circ} \mathrm{F}$. |  | ${ }^{\circ} \mathrm{F}$. |  | ${ }^{\circ} \mathrm{F}$. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 100 feet. | 54.1 | 1,000 feet. | 65.1 | 1,900 feet. | 74. 9 |
| 200 feet. | 55.0 | 1,100 feet. | 66.0 | 2,000 feet. | 76.2 |
| 300 feet. | 56.3 | 1,200 feet. | 67.0 | 2,100 feet. | 77.5 |
| 400 feet. | 57. 8 | 1,300 feet. | 68.5 | 2,200 feet. | 79.0 |
| 500 feet. | 59.4 | 1,400 feet. | 69.9 | 2,200 feet. | 80.3 |
| 600 feet | 60. 7 | 1,500 feet. | 70.9 | 2,400 feet. | 81. 4 |
| 700 feet | 61. 7 | 1,600 feet. | 72.6 | 2,500 feet. | 82.6 |
| 800 feet. | 62.7 | 1,700 feet. | 73.5 | 2,600 feet. | 83.4 |
| 900 feet. | 3. 66 | 1,800 feet. | 5. 47 | 3,250 feet. | 91. 9 |

The rates of increase indicated by these observations are as follows: 100 to 3,250 feet, $1^{\circ}$ in 83.3 feet; 1,000 to 3,250 feet, $1^{\circ}$ in 84 feet.

## VOLCANO JUNCTION.

The 4,527-foot well at Sandhill, 4 miles north of Volcano Junction, Wood County, W. Va., was tested at different depths by Mr. Van Orstrand with a mercury thermometer for some tests and electrical resistance coil for others. The results were as follows:

Temperature measurements in well 4 miles north of Volcano Junction, W. Va.

|  | ${ }^{\circ} \mathrm{F}$ |  | ${ }^{\circ} \mathrm{F}$ |  | F. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 100 feet. | .$^{1} 54.2$ | 1,500 feet. | 68.0 | 3,250 feet. | ${ }^{195.0}$ |
| 500 feet. | 57.2 | 1,750 feet. | . ${ }^{170.9}$ | 3,500 feet. | ${ }^{199.0}$ |
| 700 feet. | 58.6 | 2,000 feet. | 74.8 | 3,750 feet. | ${ }^{1} 103.7$ |
| 800 feet. | 59. 6 | 2,250 feet. | ${ }^{179.1}$ | 4,000 feet. | ${ }^{1} 108.8$ |
| 900 feet. | 61.1 | 2,500 feet. | 82.7 | 4,100 feet. | ${ }^{1} 110.7$ |
| 1,000 feet | 62.8 | 2,750 feet. | ${ }^{186.2}$ | 4,250 feet. | ${ }^{1} 113.1$ |
| 1,250 feet | 65. 2 | 3,000 feet. | 91. |  |  |

The rates of increase indicated by these observations are as follows: 100 to 4,250 feet, $1^{\circ}$ in $70 \frac{1}{2}$ feet; 1,000 to 4,250 feet, $1^{\circ}$ in 64.6 feet.

## GRANTVILLE.

A 4,610-foot hole on the Poling farm, 5 miles southeast of Grantsville, W. Va., was tested by Mr. Van Orstrand with the following results:

Temperature measurements in well 5 miles southeast of Grantsville, W. Va.

|  | ${ }^{\circ} \mathrm{F}$. |  | ${ }^{\circ} \mathrm{F}$. |  | F. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 100 feeb. | .55. 3 | 1,200 feet. | .67.2 | 2, 400 feet. | 80.5 |
| 200 feet. | . 555 | 1,300 feet. | .68.3 | 2, 500 feet. | 81. 8 |
| 300 feet. | .56.8 | 1,400 feet. | .69.5 | 2,600 feet. | 82.9 |
| 400 feet. | . 57.9 | 1,500 feet. | 70.3 | 2,700 feet. | 84. 2 |
| 500 feet. | .59.2 | 1, 600 feet. | 71.7 | 2, 800 feet. | 85.4 |
| 600 feet. | .60.4 | 1, 700 feet. | 72.5 | 2, 900 feet | 86. 8 |
| 700 feet. | .61. 6 | 1,800 feet. | 73. 2 | 3, 000 feet. | 88. 1 |
| 800 feet. | .62.9 | 2,000 feet. | .75. 4 | 3, 500 feet. | 95. 2 |
| 900 feet. | .64.1 | 2, 100 feet. | .76.9 | 4, 000 feet | .101. 8 |
| 1,000 feet. | 65. 2 | 2, 200 feet. | 78. 1 | 4,500 feet. | .110.8 |
| 1, 100 feet. | . 65.9 | 2,300 feet. | .79.4 |  |  |

- Some of the rates of increase indicated by these figures are as follows: $100-4,500$ feet, $1^{\circ}$ in 80 feet; $100-2,800$ feet, $1^{\circ}$ in 90 feet; $1,000-4,500$ feet, $1^{\circ}$ in 76.7 feet; $2,000-4,500$ feet, $1^{\circ}$ in 70.3 feet; $3,000-4,500$ feet, $1^{\circ}$ in 66 feet.


## SHINNSTON.

The 4,920 -foot hole bored by the Hope Gas Co. on the farm of E. Robinson-at Wyatt, 5 miles northwest of Shinnston, was tested at every hundred feet to a depth of 2,000 feet by Mr. Van Orstrand with the following results:

Temperature measurements in a 4,290-foot hole at Wyatt, W. Va.

|  | ${ }^{\circ} \mathrm{F}$. |  | ${ }^{\circ} \mathrm{F}$. |  | ${ }^{\circ} \mathrm{F}$. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 100 feet | . 54.5 | 800 feet. | 62. 0 | 1,500 feet. | . 68.4 |
| 200 feet. | .55. 5 | 900 feet | 62.9 | 1,600 feet. | .$^{1} 68.0$ |
| 300 feet. | . 56.3 | 1,000 feet. | .64. 2 | 1,700 feet. | .70.7 |
| 400 feet. | . 57.5 | 1, 100 feet. | .65.0 | 1, 800 feet. | .71.9 |
| 500 feet. | . 58.9 | 1,200 feet. | .65. 6 | 1,900 feet. | .72. 9 |
| 600 feet. | .59.9 | 1,300 feet. | .66.5 | 2,000 feet. | .74.1 |
| 700 feet. | . 61.0 | 1, 400 feet. | . 67.7 |  |  |

The rate of increase indicated by the difference.in temperature between 100 and 2,000 feet is close to $1^{\circ}$ in 96 feet; from 1,000 to 2,000 feet the rate is $1^{\circ}$ in 101 feet.

## LAKE WELL.

This well reached a depth of 7,579 feet when drilling ceased in 1919. It is about 8 miles south of Fairmont, W. Va. Van Orstrand found the temperature at 7,500 feet to be $168.6^{\circ}$, indicating an average rate of temperature increase of $1^{\circ}$ in 66 feet.

Mr. Van Orstrand has also given a series of measurements from 100 to 5,400 feet as follows:

Temperature measurements in well on Lake Farm, 4 miles northeast of Valley Falls, W. Va.

|  | ${ }^{\circ} \mathrm{F}$. |  | ${ }^{\circ} \mathrm{F}$. |  | ${ }^{\circ} \mathrm{F}$. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 100 feet. | .53.7 | 1,300 feet. | .69.2 | 3,250 feet. | . 93.6 |
| 200 feet. | ..55. 8 | 1,400 feet | .70.6 | 3,500 feet. | 95. 4 |
| 300 feet. | ..57.5 | 1,500 feet | .71.8 | 3,750 feet | 100.9 |
| 400 feet | ..58.8 | 1,600 feet | .73. 1 | 4,000 feet. | 104.5 |
| 500 feet. | .59.9 | 1,700 feet. | .74. 3 | 4,250 feet. | 106. 0 |
| 600 feet. | . 61.0 | 1,800 feet. | .75. 3 | 4,500 feet. | .111. 4 |
| 700 feet. | ..62.2 | 1,900 feet. | .76. 7 | 4,750 feet. | 116. 2 |
| 800 feet. | .62.8 | 2,000 feet. | .78.4 | 5,000 feet. | 120.7 |
| 900 feet. | $\therefore 64.4$ | 2,250 feet. | 82. 4 | 5,250 feet. | .126. 1 |
| 1,000 feet. | .65. 6 | 2,500 feet. | .86. 0 | 5,400 feet. | .129.2 |
| 1,100 feet. | .66.8 | 2,750 feet. | .89.2 |  |  |
| 1,200 feet. | ..68. 1 | 3,000 feet. | . 90.5 |  |  |

The rate of increase from 100 to 5,400 feet is $1^{\circ}$ in 70 feet.
GOFF WELL.
The well on the Goff farm is about 8 miles northeast of Clarksburg in the northern part of West Virginia. Its depth is 7,386 feet.

Tests were made by Mr. Van Orstrand at various depths in the well with the following results:

Temperature measurements in Goff well 4 miles northeast of Bridgeport, W. Va. -E.

|  | ${ }^{\circ} \mathrm{E}$ |  | \%. |  | ${ }^{\circ} \mathrm{F}$ |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 100 feet | .55. 6 | 1,300 feet. | 66.5 | 2,500 feet. |  |
| 200 feet | . 56.3 | 1,400 feet. | 67. 2 | 3,000 feet. | 87. 6 |
| 300 feet. | . 57.5 | 1,500 feet. | 67.8 | 3,500 feet. | 33. |
| 400 feet. | .58. 8 | 1,600 feet. | . 68.0 | 4,000 feet. | 0. |
| 500 feet | .60. 2 | 1,700 feet. | 67.8 | 4,500 feet. | 107. |
| 600 feet. | .61. 6 | 1,800 feet. | .70.8 | 5,000 feet. | 114. |
| 700 feet. | .62. 5 | 1,900 feet. | . 73.6 | 5,500 feet. | 122.3 |
| 800 feet. | 63. 5 | 2,000 feet. | 74.9 | 6,000 feet. | 132. |
| 900 feet. | . 64.5 | 2,100 feet. | 76.1 | 6,500 feet. | 143. |
| 1,000 feet | .65. 3 | 2,200 feet. | 77. 3 | 7,000 feet. | 153. |
| 1,100 feet | . 65.8 | 2,300 feet. | 78.3 | 7,250 feet. | 157. |
| 1,200 feet | 65.5 | 2,400 feet. | 79.3 | 7,310 feet. | 158. |

Some of the rates of increase indicated by these figures are as follows: 100 to 7,310 feet, $1^{\circ}$ in 70.2 feet; 100 to 7,000 feet, $1^{\circ}$ in 70.7 feet; 100 to 5,000 feet, $1^{\circ}$ in 83.6 feet; 1,000 to 7,310 feet, $1^{\circ}$ in 67.8 feet; 2,000 to 7,310 feet, $1^{\circ}$ in 63.7 feet; 4,000 to 7,250 feet, $1^{\circ}$ in 56.3 feet.

## WISCONSIN.

## TEMPERATURES.

Temperatures of flows of certain Wisconsin wells are given in the reports of the State Survey ${ }^{2}$ and water-supply papers of the United States Geological Survey, but for the most part the depth from which the flows are derived is not given, and some of the wells are pumped. The following selected list may afford a few data of value:

[^43]Temperatures in wells in Wisconsin.

| Location. | $\begin{aligned} & \text { Depth } \\ & \text { (feet). } \end{aligned}$ | $\begin{aligned} & \text { Flow } \\ & \text { (gan- } \\ & \text { lons } \\ & \text { per } \\ & \text { min- } \\ & \text { ute). } \end{aligned}$ | $\begin{aligned} & \text { Tem- } \\ & \text { pera- } \\ & \text { ture } \\ & \text { ( }{ }^{\circ} \mathrm{F} \text {.). } \end{aligned}$ | Mean annual tem- pera- ture ( ${ }^{\circ} \mathrm{F}$.). | Depth per degree of increase in tem-pera(feet). | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Crawford County: <br> Prairie du Chien a |  |  |  |  | 54 |  |
| Pro........... | 960-1,004 | 540 | 57 | 47.5 |  |  |
| Green Lake County: Berlin. |  | Many. | 52 | 845.2 | 63 | Flow from below 324 feet.c |
| Marinette County: Marinette. | 716 |  | 49 | d 43.3 | 71 | Cased to 333 feet; water at 405-410 feet; not any below 415 feet. |
| Oconto County: Oconto.. | 400 | 450 | 50 | e 44.1 | 68 | Cased to 225 feet; source of |
| Pepin County: Durand $f \ldots$ | 200 | Many. | 51 | 044.2 | 30 |  |
| Richland County: Richland Center. | 748 | 2,400 | 53 | ${ }^{\circ} 45.3$ | 96 | Pumped 11 feet; water probably from considerably above bottom. |
| Sheboygan County: Sheboygan.f | 1,475 | 225 | i 59.1 | ${ }^{\text {j }} 45.7$ | 100 | Pumped 104 feet; main supply probably from 1,340 feet. |
| Vernon County: Genoa...... | 450 | 200 | 52 | ${ }^{k} 46.7$ | 87 | Pumped 30 feet. |
| We Soto............il. |  | (126 | 52 58 | 146.7 46 | 84 88 | Flow. ${ }_{\text {A }}$ (ift, pumped 155 feet. |
| Walworth County: Elk- horn. | 1,050 | Many. | 58 | 46 | 88 | Air lift, pumped 155 feet. |

a Calculated to depth of 1,340 feet.
$b$ Oshkosh average for 28 years.
c Wisconsin Geol. and Nat. Hist. Survey Bull. 35, p. 364, 1914.
d Average at Menominee, Mich., for 18 years.
e Average for 26 years
f. Wisconsin Geol. and Nat. Hist. Survey Bull. 35, p. 502, 1914.

Idem, p. 284.
h Viroqua average for 27 years plus $0.5^{\circ}$.
$i$ Wisconsin Geol. and Nat. Hist. Survey, p. 569, 1914.
$j$ Average for 18 years.
$k$ Mean of La Crosse and Prairie du Chien.
$l$ Delavan average for 24 years.

## SUMMARY.

Exclusive of Sheboygan, Elkhorn, and Richland wells, of which the true underground temperatures are uncertain, there is a fair degree of accordance in the rates indicated by the Prairie du Chien, Berlin, Marinette, Genoa, De Soto, and Oconto wells. Owing, however, to some uncertainty as to the precise source of water, the figures can be regarded only as approximations. The temperature of the flow from relatively shallow wells at Durand indicates a high rate.

## GEOLOGIC RELATIONS.

The wells in the above list are all in the Paleozoic rocks on the slope of the low structural dome whose summit is in the broad area of crystalline rocks that occupies the northern part of the State. There are no local disturbances of any note. All the wells draw from sandstone, either the St. Peter sandstone, as at Oconto and Sheboygan, or the Upper Cambrian, some distance below the St. Peter horizon. At Marinette the granite "bedrock" is reached at a depth of 712 feet, and at Richland at 665 feet. At Durand it is
not far below the bottom of the 200 -foot well. At Sheboygan a hole 1,782 feet deep did not reach the bedrock. There is no apparent relation between the gradients and geologic conditions.

## WYOMING.

The deep-well temperatures in Wyoming given in the following table were carefully determined with Darton thermometers sunk to the bottom of the holes by my assistants or myself.

Temperatures in wells in Wyoming.

|  | Depth (feet). | Temperature ( ${ }^{\circ} \mathrm{F}$.). | Mean annual air temperature ( ${ }^{\circ} \mathrm{F}$.). | $\begin{aligned} & \text { Depth } \\ & \text { per } \\ & \text { degree of } \\ & \text { increase } \\ & \text { in tem. } \\ & \text { perature } \\ & \text { (feet). } \end{aligned}$ | Remarks. |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Big Horn County: <br> SW. $\frac{1}{4}$ SE. $\frac{1}{2} \sec .27$, T. 102 <br> R. 48,10 miles southwest of Meteetse. <br> Cambria. $\qquad$ <br> Newcastle, 2 miles west.... | $\begin{array}{r} 1,400 \\ 1,475 \\ 374 \end{array}$ | 66 59 63 | $\begin{array}{r} a 40 \\ \\ b 45.2 \\ 45.2 \end{array}$ | 54 107 21 | Dry hole; thermometer sunk by C. A. Fisher. <br> Thermometer sunk by G. B. Richardson, 1899. Thermometer sunk by N. H. Darton, 1900. |

a Mean annual air temperature at Four Bear, 10 miles west, plus $0.3^{\circ}$ for lower altitude. $b$ Newcastle average.

All these figures are correct, and no vitiating conditions are indicated. The Cambria and Newcastle wells were full of water. The steep gradient of the Newcastle well is inexplicable.



[^0]:    ${ }^{1}$ Prestwich, J., On underground temperatures: Roy. Soc. London Proc., vol. 41, pp. 1-116, 1886.
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    ${ }^{6}$ Climatological data for the United States by sections, vol. 3, No. 13, Washington, 1917.
    ${ }^{6}$ Slichter found while investigating the underflow of Arkansas River in western Kansas in 1301 that the temperature of water in shallow wells ranged from $56^{\circ}$ to $58^{\circ} \mathrm{F}$., or $2^{\circ}$ to $4^{\circ}$ higher than the mean annual air temperature of the region. (See U. S. Geol. Survey Water-Supply Paper 153, p. 12, 1906.)

[^1]:    ${ }^{1}$ Schwennesen, A. T., Ground water in San Simon Valley, Ariz. and N. Mex.: U. S. Geol. Survey Water-Supply Paper 425, pp. 1-35, 1917.
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[^3]:    $a$ Mean annual temperature for Riverside for 35 years
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    c San Bernardino average for 25 years.

[^4]:    ${ }^{1}$ Hydrology of San Bernardino Valley, Calif.: U. S. Geol. Survey Water-Supply Paper 142, 124 pp., 1905.

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    d Pasadena average for 27 years.

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[^8]:    a Bakersfield average for 28: years; wells 27 to 100 feet, $68^{\circ}$ to $71^{\circ}$ mostly.
    6 Hanford average for 17 years.
    c Merced average for 43 years.
    d Stockton average for 46 years.
    e Well cased to bettom; flows also at 980 and 1,030 feet, which may mingle.

[^9]:    a Visalia average for 30 years.
    ${ }^{b}$ Mean of Visalia and Porterville.
    c Porterville average for 28 years.

[^10]:    ${ }^{1}$ Mendenhall, W. C., Ground waters of the Indio region, Calif., with a sketch of the Colorado Desert: U. S. Geol. Survey Water-Supply Paper 225, 1909.

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    ${ }^{b}$ Canon City average for 29 years.
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    © Rocky Ford average plus 0.50 . $^{\circ}$

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    ${ }^{2}$ Florida Geol. Survey Third Rept., pp. 114-115, 151, 1910.

[^15]:    a Malabar average for 25 years.

    - Orlando average minus $3^{\circ}$.
    c 21 years average.
    d Middleburg average.
    e Jacksonville average for 46 years.
    $f$ A verage for 48 years.
    9 Average for 33 years.
    $h$ Braidentown average plus $2^{\circ}$.

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    $d$ J. H. Sheppard, observer.
    ${ }^{e}$ E.S. Nettleton gives $72.0^{\circ}$.
    $f$ Huron average.
    A verage for 10 years.
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