## Solubility of gases in water: Henry's Law

• concentration dissolved  $\alpha$  partial pressure of the gas

$$K_H \text{ (units mol L}^{-1} \text{ atm}^{-1}) = c_X/p_X$$

Large  $K_H$  means high solubility;  $K_H$  always decreases with T; gases less soluble at higher T (all gases, all solvents)

**Henry's law constants at 298 K**: K<sub>H</sub> in mol L<sup>-1</sup> atm<sup>-1</sup> from Seinfeld and Pandis, *Atmospheric Chemistry and Physics*, Wiley, 1998 p. 341; values do not include subsequent reactions of the dissolved species, such as acid dissociation.

substance	$\mathbf{K}_{\mathbf{H}}$	substance	$\mathbf{K}_{\mathbf{H}}$
$\mathrm{O}_2$	$1.3 \times 10^{-3}$	NO	$1.9 \times 10^{-3}$
$NO_2$	$1.2 \times 10^{-2}$	$O_3$	$1.13\times10^{-2}$
$N_2O$	$2.5 \times 10^{-2}$	$\mathrm{CO}_2$	$3.4 \times 10^{-2}$
$H_2S$	0.12	$\mathrm{SO}_2$	1.23
$CH_3ONO_2$	2.6	$CH_3O_2$	6
OH	25	$HNO_2$	49
$NH_3$	62	CH <sub>3</sub> OH	220
CH <sub>3</sub> OOH	230	HCl	730
$\mathrm{HO}_2$	2000	CH <sub>3</sub> COOH	8800
$H_2O_2$	75,000	$HNO_3$	200,000

Note: Environment Canada quotes  $K_H$  in the reverse direction (escape from water): units Pa m<sup>3</sup> mol<sup>-1</sup>, hence large  $K_H$ —> low water solubility.

Solubility of  $O_2$  in water – context is whether water will support aquatic life

$$K_{\rm H} = 1.3 \times 10^{-3} \text{ mol L}^{-1} \text{ atm}^{-1} \text{ at equilibrium, at } 25^{\circ}\text{C}$$
 $\longrightarrow c(O_2) = 2.7 \times 10^{-4} \text{ mol /L}$ 
 $\longrightarrow 8.7 \text{ mg/L (8.7 ppm)}$ 

Note definition of ppm for solids and solutions (by mass)

- $c(O_2) < 8.7 \text{ ppm}$ :
  - at higher temperatures (thermal pollution)
  - if decaying or oxidizable material consumes  $O_2$  -> concept of biochemical oxygen demand (BOD)
  - water is stagnant (reduced air exchange)

## Measures of the oxygen status of water

- BOD; incubate with microorganisms for 5 days in closed container, measure  $c(O_2)$  before and after
- Chemical Oxygen Demand (COD) titrate the sample vs excess  $Na_2Cr_2O_7/H^+$ ; easily oxidized substances consume  $Na_2Cr_2O_7$ ; determine the amount of  $Na_2Cr_2O_7$  left over; 1 mol  $Na_2Cr_2O_7 \equiv 1.5$  mol  $O_2$
- Total Organic Carbon (TOC) oxidize the organic compounds to CO<sub>2</sub> by combustion; analyze CO<sub>2</sub> produced
- Dissolved Oxygen (DO) often done by titration:

$$Mn^{2+} + 2OH^{-} + \frac{1}{2}O_{2} \longrightarrow MnO_{2}(s) + H_{2}O$$
  
 $MnO_{2} + 4H^{+} + 2I^{-} \longrightarrow I_{2} + Mn^{2+} + 2H_{2}O$   
 $I_{2} + Na_{2}S_{2}O_{3} \longrightarrow Na_{2}S_{4}O_{6} + 2NaI$ 

## CO<sub>2</sub> solubility in water

• More complex than  $O_2$  because  $CO_2(aq) \equiv H_2CO_3(aq)$ , which can dissociate through acid-base equilibria

$$CO_{2}(g) + H_{2}O(l) \implies H_{2}CO_{3}(aq)$$

$$K_{H} = 3.4 \times 10^{-2} \text{ mol } L^{-1} \text{ atm}^{-1}$$

$$H_{2}CO_{3}(aq) \implies H^{+}(aq) + HCO_{3}^{-}(aq)$$

$$K_{a} = 4.2 \times 10^{-7} \text{ mol } L^{-1}$$

- Note that in carrying out calculations, the concentrations of CO<sub>2</sub>(g) and H<sub>2</sub>CO<sub>3</sub>(aq) do not change, because the atmosphere is an inexhaustible reservoir
- Total "dissolved carbonate" =  $\{H_2CO_3(aq) + HCO_3^-(aq) + CO_3^{2-}(aq)\}$ : increases with increasing pH

## Calculation of the solubility of CO<sub>2</sub> in pure water

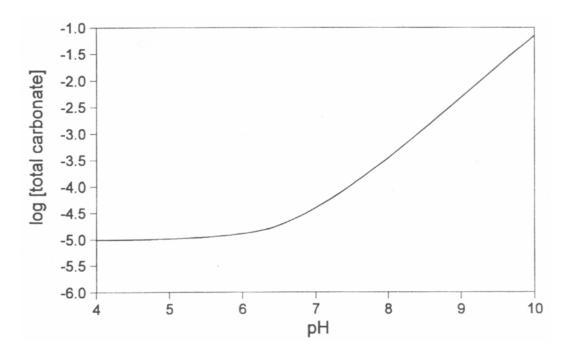
- $p(CO_2, g) = 375 \text{ ppmv} \longrightarrow c(CO_2, aq) = 1.3 \times 10^{-5} \text{ mol/L}$
- for  $K_a = [H^+][HCO_3^-]/[H_2CO_3] = 4.2 \times 10^{-7} \text{ mol/L } (25^{\circ}\text{C})$

[H<sup>+</sup>][HCO<sub>3</sub><sup>-</sup>] = 
$$x^2$$
 = K<sub>a</sub> [H<sub>2</sub>CO<sub>3</sub>]  
=  $(1.3 \times 10^{-5} \text{ mol/L})(4.2 \times 10^{-7} \text{ mol/L})$   
 $x = 2.3 \times 10^{-6} \text{ mol/L}$   
pH = 5.63;  
total "carbonate" =  $\{1.3 \times 10^{-5} + 2.3 \times 10^{-6} \text{ mol/L}\}$   
=  $1.5 \times 10^{-5} \text{ mol/L}$ 

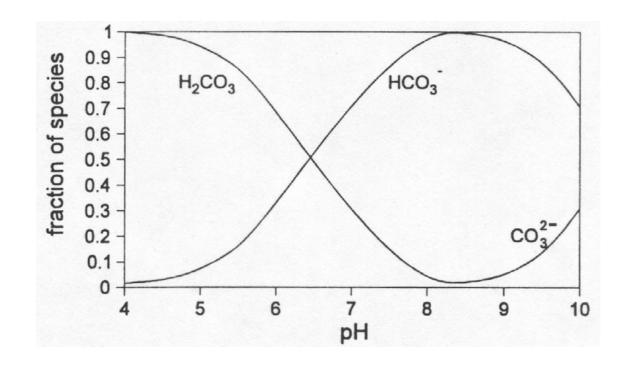
• Even completely clean water in equilibrium with atmospheric  $CO_2$  does not have pH = 7!! Keep this thought for discussion of acid rain.

When the pH of the water is fixed by the presence of other solutes:

• total dissolved carbonate increases as pH rises



• Note the speciation of carbonate

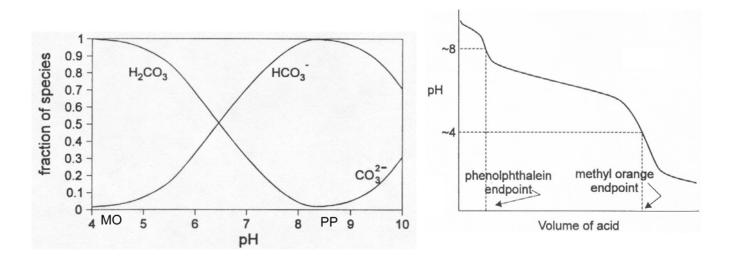


**Alkalinity** of water is a measure of the concentration of all bases in the water, **not** its pH, which is determined largely by the strongest base present: text pp. 140-142

- Alkalinity is measured by titrating the water against standard acid ≡ moles/concentration of H<sup>+</sup> needed to neutralize the bases
- Phenolphthalein alkalinity is the amount of acid needed to reach the phenolphthalein endpoint (pH 8.5)

# remembering that titration is from high to low pH

- Total alkalinity is the amount of acid needed to reach the methyl orange endpoint (pH 4)
- If there are no other bases present (as in *e.g.*, industrial waste water), the phenolphthalein endpoint measures mostly  $CO_3^{2-}$ ; the methyl orange endpoint measures  $CO_3^{2-} + HCO_3^{-}$



- Two measurements to determine both  $CO_3^{2-}$  and  $HCO_3^{-}$ :
  - both total and phenolphthalein alkalinity or
  - one of the above plus pH  $\rightarrow$  ratio  $[CO_3^{2-}]/[HCO_3^{-}]$

**Hardness** of water is a measure of the concentration of "hardness ions" (mainly Ca<sup>2+</sup> and Mg<sup>2+</sup>) that form insoluble salts, especially carbonates: text, pp. 142-146.

## **Analysis of hardness ions:**

- titration vs EDTA<sup>4-</sup> using Eriochrome Black T indicator (Ca only)
- atomic absorption spectroscopy

## Origin of hardness ions:

dissolution of gypsum

$$CaSO_4(s) \rightleftharpoons Ca^{2+}(aq) + SO_4^{2-}(aq)$$

• dissolution of limestone rocks: CaCO<sub>3</sub> (limestone); CaCO<sub>3</sub>.MgCO<sub>3</sub> (dolomite)

**NOT** 
$$MCO_3(s) \rightleftharpoons M^{2+}(aq) + CO_3^{2-}(aq)$$
  
**BUT**  $MCO_3(s) + H_2CO_3(aq) \rightleftharpoons M^{2+}(aq) + 2HCO_3^{-}(aq)$ 

- Note that underground,  $p(CO_2)$  is often much greater than 370 ppmv
- In what follows, note the text, footnote 8, p. 143 about  $K_{sp}$  calculations!!

$$\begin{split} \text{CaSO}_4 & \qquad \qquad K_{sp} = 4 \times 10^{-5} \; (\text{mol L}^{-1})^2 \\ \text{CaCO}_3 & \qquad \qquad K_{sp} = 6 \times 10^{-9} \; (\text{mol L}^{-1})^2 \\ \text{$^{1}\!\!\!/_{2}$CaCO}_3.\text{MgCO}_3 & \qquad K_{sp} = 5 \times 10^{-7} \; (\text{mol L}^{-1})^2 \end{split}$$

## **Dissolution of CaCO<sub>3</sub>**

$$\mathbf{K} =$$

$$\operatorname{CaCO}_{3}(s) \iff \operatorname{Ca}^{2+}(\operatorname{aq}) + \operatorname{CO}_{3}^{2-}(\operatorname{aq}) \qquad \mathbf{K}_{\operatorname{sp}}$$

$$H_{2}\operatorname{CO}_{3}(s) \iff H^{+}(\operatorname{aq}) + \operatorname{HCO}_{3}^{-}(\operatorname{aq}) \qquad \mathbf{K}_{\operatorname{al}}$$

$$H^{+}(\operatorname{aq}) + \operatorname{CO}_{3}^{2-}(\operatorname{aq}) \iff \operatorname{HCO}_{3}^{-}(\operatorname{aq}) \qquad 1/\mathbf{K}_{\operatorname{a2}}$$

Net: 
$$CaCO_3(s) + H_2CO_3(aq) \rightleftharpoons Ca^{2+}(aq) + 2HCO_3^{-}(aq)$$

or: 
$$CaCO_3(s) + H_2CO_3(aq) \rightleftharpoons Ca(HCO_3)_2(aq)$$

- K for net reaction =  $K_{sp} \times K_{a1}/K_{a2} = 5 \times 10^{-5} \text{ (mol L}^{-1})^2$
- when expressed as "ppm of CaCO<sub>3</sub>", values up to 300 ppm are obtained in hard water areas

**Hard water**: contains hardness ions: usually limestone areas *e.g.*, southern Ontario

**Soft water**: low concentrations of hardness ions: sandstone and granite areas e.g., northern and eastern Ontario

All water must have a **balance of cations and anions**; : hard water is usually well buffered against acidification —> relatively high concentrations of weak bases

Thus alkalinity is a measure of buffering capacity; high alkalinity usually correlates with high hardness

Water Softening: critical application for steam boilers due to deposition of salts

When hard water is heated:

$$Ca(HCO_3)_2 (aq) \rightleftharpoons CaCO_3(s) + H_2CO_3(aq) \longrightarrow CO_2(g)$$

Water softening is the process of removing hardness ions

1. Lime Softening (industrial use only): neutralize HCO<sub>3</sub><sup>-</sup> with OH<sup>-</sup>

$$Ca(OH)_2 (aq) + Ca(HCO_3)_2 (aq) \rightleftharpoons CaCO_3(s) + 2H_2O$$

2. Ion exchange resins, e.g., Na(A) where (A) = polymeric anion – example of Ca<sup>2+</sup> removal through cation exchange

$$\operatorname{Ca}^{2+}(\operatorname{aq}) + 2\operatorname{Na}(A)_{\operatorname{res}} \rightleftharpoons 2\operatorname{Na}^{+}(\operatorname{aq}) + \operatorname{Ca}(A_{2})_{\operatorname{res}}$$

Resin regeneration with concentrated brine:

$$2\text{Na}^+(\text{aq}) + \text{Ca}(A_2)_{\text{res}} \rightleftharpoons \text{Ca}^{2+}(\text{aq}) + 2\text{Na}(A)_{\text{res}}$$

3. Deionized water: cation and anion exchangers in series, using H<sup>+</sup> form of the cation exchanger and OH<sup>-</sup> form of the anion exchanger – example of CaSO<sub>4</sub>

$$\operatorname{Ca}^{2+}(\operatorname{aq}) + 2\operatorname{H}(A)_{\operatorname{res}} \rightleftharpoons 2\operatorname{H}^{+}(\operatorname{aq}) + \operatorname{Ca}(A_{2})_{\operatorname{res}}$$

$$SO_4^{2-}(aq) + 2(C)OH_{res} \rightleftharpoons 2OH^-(aq) + (C_2)SO_{4res}$$

Hence:

$$2H^{+}(aq) + 2OH^{-}(aq) \longrightarrow 2H_{2}O$$

• Regeneration of the resin beds????

**Seawater**: a solution of high ionic strength. The main environment we will encounter where activities must be used rather than concentrations.

Ion	conc, mol/L	input, Tmol/yr	τ, Myr
$Na^+$	0.46	9.0	70
$K^+$	0.010	1.9	7
$Mg^{2+}$ $Ca^{2+}$	0.054	5.5	10
$Ca^{2+}$	0.010	12.2	1
$C1^{-}$	0.55	7.2	100
$SO_4^{2-}$	0.028	3.8	10
$HCO_3^-$	0.0023	32	0.1
$CO_3^{2-}$	0.0003	included wit	h HCO <sub>3</sub>

- Ocean water **approximately** in equilibrium with CaCO<sub>3</sub>, but  $Q_{sp} = [Ca^{2+}][CO_3^{2-}] >> K_{sp}$ : text, p. 150
- First reason:  $a(Ca^{2+})$  and  $a(CO_3^{2-}) < [Ca^{2+}][CO_3^{2-}]$ , *i.e.*,  $\gamma(Ca^{2+}) \sim 0.26$ ;  $\gamma(CO_3^{2-}) \sim 0.20$
- Second reason: complexation: formation of species such as:

(CaSO<sub>4</sub>): 8% of total Ca; (CaHCO<sub>3</sub>)<sup>+</sup>: 1% of total Ca

(MgCO<sub>3</sub>): 64% of total CO<sub>3</sub>; (NaCO<sub>3</sub>)<sup>-</sup>: 19% of total CO<sub>3</sub>; (CaCO<sub>3</sub>): 7% of total CO<sub>3</sub>

## Irrigation and water quality

- Read text pp. 147-149
- Read article from *The Economist*, link to internet =

http://www.economist.com/displaystory.cfm?story\_id=1906914

## **Properties of Water**

• Amounts on Earth: Oceans,  $\sim 10^{20}$  mol Rivers and lakes,  $\sim 10^{15}$  mol

## Freezing point depression

• Solutes depress the freezing point of water

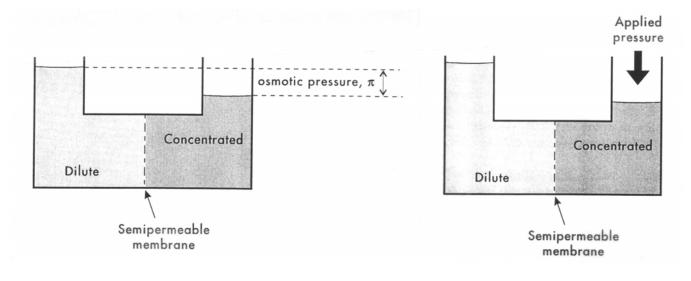
$$\Delta T = K_f \times m$$
 $K_f = \text{molal freezing point depresssion contant,}$ 
units K kg mol<sup>-1</sup>
 $m = \text{molal concentration of solute, mol kg}^{-1}$ 

- The freezing point depression is *independent* of the identity of the solute. For ionic solutes consider all the ions separately, *e.g.*, for NaCl there are *two* solutes to consider, Na<sup>+</sup> and Cl<sup>-</sup>
- Applications:

road salt trees in winter, fish in polar oceans (laboratory): determining molar mass

#### **Osmosis and Reverse Osmosis**

- osmotic pressure  $\pi = c \times RT$  c in mol  $L^{-1}$  R in L atm mol<sup>-1</sup>  $K^{-1}$  $\pi$  in atm
- osmotic pressure independent of the solute identity
- applications
   water rise in trees
   hypertonic and hypotonic solutions; impact on cells
   (laboratory): measuring molar mass of polymers and
   biopolymers
- reverse osmosis: a method of water purification



**Osmosis** 

**Reverse Osmosis**